SPECIAL PAPERS IN PALAEONTOLOGY · 53

Mid-Dinantian ammonoids from the Craven Basin, north-west England



THE PALAEONTOLOGICAL ASSOCIATION

MID-DINANTIAN AMMONOIDS FROM THE CRAVEN BASIN, NORTH-WEST ENGLAND

BY
NICHOLAS J. RILEY

with 8 plates and 51 text-figures

THE PALAEONTOLOGICAL ASSOCIATION LONDON

JUNE 1996

© The Palaeontological Association, 1996

CONTENTS

	page
Abstract	5
Introduction	5
STRATIGRAPHY OF THE WORSTON SHALE GROUP	5
Ammonoid Zonation	7
Fascipericyclus-Ammonellipsites (FA) Zone	9
Bollandites-Bollandoceras (BB) Zone	9
Beyrichoceras (B) Zone	9
Ammonoid Distribution	9
Clitheroe Limestone Formation	9
Hodder Mudstone Formation	10
Hodderense Limestone Formation	11
Pendleside Limestone Formation	11
Locality Index	11
Clitheroe Limestone Formation	11
Hodder Mudstone Formation	11
Hodderense Limestone Formation	12
Pendleside Limestone Formation	12
Systematic Palaeontology	12
Order Agoniatida Ruzhencev, 1957	12
Suborder Prolecanitina Miller and Furnish, 1954	12
Superfamily Prolecanitacea Hyatt, 1884	12
Family Prolecanitidae Hyatt, 1884	12
Subfamily Prolecanitinae Hyatt, 1884	13
Genus Michiganites Ruzhencev, 1962	13
Subfamily Protocanitinae Weyer, 1972a	14
Genus Merocanites Schindewolf, 1922	14
Order Goniatitida Hyatt, 1884	21
Suborder Goniatitina Hyatt, 1884	21
Superfamily Dimorphocerataceae Hyatt, 1884	21
Family Dimorphoceratidae Hyatt, 1884	22
Subfamily Dimorphoceratinae Hyatt, 1884	22
Genus Dimorphoceras Hyatt, 1884	22
Superfamily Girtyocerataceae superfam. nov.	25
Family Girtyoceratidae Wedekind, 1918	25
Genus Beyrichoceras Foord, 1903	26
Superfamily Nomismocerataceae Librovitch, 1957	2.7

Family Nomismoceratidae Librovitch, 1957	28
Genus Eonomismoceras Kusina, 1974	29
Genus Nomismoceras Hyatt, 1884	30
Superfamily Pericyclaceae Hyatt, 1900	31
Family Muensteroceratidae Librovitch, 1957	31
Genus Beyrichoceratoides Bisat, 1924	32
Genus Bollandoceras Bisat, 1952	40
Genus Dzhaprakoceras Popov, 1965	48
Family Pericyclidae Hyatt, 1900	69
Subfamily Ammonellipsitinae subfam. nov.	69
Genus Ammonellipsites Parkinson, 1822	70
Genus Helicocyclus Schindewolf, 1951 a	73
Subfamily Pericyclinae Hyatt, 1900	74
Genus Bollandites Bisat, 1952	75
Genus Parahammatocyclus gen. nov.	78
Genus Rotopericyclus Turner, 1948	80
DISCUSSION AND CONCLUSIONS	82
ACKNOWLEDGEMENTS	83
References	83

- X

ABSTRACT. A major new ammonoid-bearing sequence is described from the Worston Shale Group, which spans the late Tournaisian (early Chadian) to mid-Viséan (early Asbian) interval in the Craven Basin of north-west England. The faunas are referable to the Fasciperic vclus-Ammonellipsites (FA), Bollandites-Bollandoceras (BB) and lower part of the Beyrichoceras Zone (B1) ammonoid zones. Particularly significant are the early Viséan assemblages from the FA Zone, which are the most diverse known, apart from faunas described previously from the northern Urals of the former Soviet Union. The described assemblages comprise: Ammonellipsites clitheroensis sp. nov., Beyrichoceras aff. obtusum, Beyrichoceratoides inflatus sp. nov., B. sp. aff. komiorum, B. sp. 1, B. sp. 2, Bollandites varians sp. nov., Bo.? sp., Bollandoceras hodderense Bol. micronotum, Bol. sp. 1, Bol. sp. 2, Dimorphoceras leagramense sp. nov., Dzhaprakoceras bellmanense sp. nov., D. catena sp. nov., D. deflexum, D. cf. djaprakense, D. flexum sp. nov., D. hispanicum, D. levis sp. nov., D. paracatena sp. nov., D. subglobosum, D. undulatum sp. nov., D. westheadi sp. nov., Economismoceras planum sp. nov., Helicocyclus divergens sp. nov., Merocanites cf. applanatus, M. quadrilobus sp. nov., Michiganites sp., Nomismoceras rotiforme, Parahammatocyclus chaigleyensis gen. et sp. nov. and Rotopericyclus postremus sp. nov. Existing suprageneric classification is modified; the new superfamily Girtyocerataceae is erected together with two new subfamilies: the Ammonellipsitinae and Pericyclinae. Several genera are reclassified, including Beyrichoceras, transferred from the Anthracoceratidae to the Girtyoceratidae, and Bollandites, transferred from the Muensteroceratidae to the Pericyclinae.

THE mid-Dinantian ammonoid assemblages described in this account, from the Worston Shale Group in the Craven Basin of north-west England, are the most prolific known, apart from those described by Kusina (1973, 1974, 1980, 1983) from the northern Urals in that part of the Commonwealth of Independent States (CIS, formerly the USSR) now known as the Russian Federation. They are also the most accurately dated mid-Dinantian assemblages known, being well constrained in terms of both foraminiferal and conodont biostratigraphies (Riley 1990b).

Dinantian ammonoids were first described from the Craven Basin of north-west England in the classic publication of Phillips (1836). Subsequent publications include those of Bisat (1924, 1934, 1950, 1952) and Moore (1930, 1936, 1939, 1946, 1950, 1952); both authors dealt with late Viséan faunas.

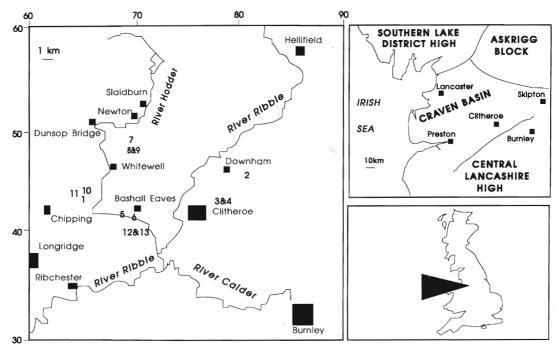
Mid-Dinantian assemblages were first referred to in Earp et al. (1961) but these were not described in detail and were only referred to a few species; the only described material was *Merocanites similis* (Crick, 1895), from the late Tournaisian, Chatburn Limestone Group. Further brief reference to mid-Dinantian assemblages from this district was made by Riley (1990a, 1990b, 1991).

STRATIGRAPHY OF THE WORSTON SHALE GROUP

This group, with its component formations and members, has been defined and described recently by Riley (1990b). The location and lithostratigraphy is summarized in Text-figures 1–2. The Dinantian sequence in the Craven Basin is very thick (over 6 km) and accumulated in an actively extending asymmetrical rift basin, trending east-north-east to west-south-west. Marine sedimentation in the early Dinantian began on a carbonate ramp which subsequently fractured and rifted into a series of intra-basinal highs and lows. The rift basin is bounded by the Askrigg Block and Southern Lake District High to the north, and the Central Lancashire High to the south. Because the base of the sequence is unexposed, the age and nature of the oldest Dinantian is unknown; however, seismic evidence (Aitkenhead et al. 1992) suggests that around 3 km are concealed.

The oldest strata exposed underlie the Worston Shale Group and comprise the Chatburn Limestone Group, of late Tournaisian age. These strata are developed in a shallow marine, clastic-influenced, carbonate ramp facies. About 1·2 km of Chatburn Limestone has been proved from outcrop and boreholes. At the basin margin there is lateral passage into red beds and evaporites (Stockdale Farm Formation).

The onset of deposition of the Worston Shale Group (Clitheroe Limestone Formation) coincided with the progressive steepening of the ramp. A variable and complex stratigraphy resulted, reflected

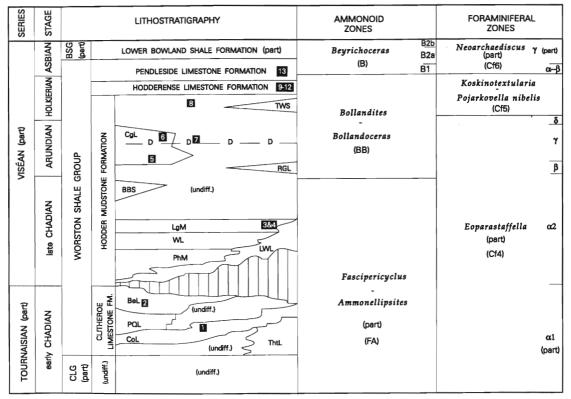


TEXT-FIG. 1. Sketch map of ammonoid localities; for explanation of locality numbers see text.

initially by the development of deeper water Waulsortian limestones and associated facies (Coplow and Bellman Limestone members) in the late Tournaisian. These biohermal limestones, of possible microbial origin (Lees and Miller 1985), display remarkable lateral thickness changes, ranging from a few metres to c. 400 m.

By the earliest Viséan (late Chadian) the steepened ramp had fractured, by tectonic rifting, into a series of intra-basinal highs and lows, resulting in a complete reorganization of depositional style. The base of the Viséan (Hodder Mudstone Formation) is therefore associated with widespread submarine erosion and unconformity, marking the switch from a carbonate ramp to a hemipelagic depositional regime. Boulder beds, derived from the underlying formation, accumulated locally, as did thin Waulsortian limestones (Limekiln Wood Limestone Member). Carbonate production within the basin soon ceased and the background dysaerobic accumulation of terrigenous mudstone was interrupted by repeated limestone turbidite influx (Whitemore, Rain Gill, Embsay and Chaigley limestone members) until early Holkerian times. These turbidites were derived from the surrounding carbonate ramps and platforms which now developed on the newly submerged Askrigg Block, Southern Lake District High and Central Lancashire High. Minor sandstone turbidite influx also occurred (Buckbanks and Twiston sandstone members) as the surrounding carbonate source areas were by-passed during lowstands. The Hodder Mudstone Formation varies in thickness from less than 100 m up to 900 m, depending on the relief of the basal surface and the relative proportion of turbidites to hemipelagic mudstones.

During the Holkerian the basin became progressively starved of sediment, culminating in the widespread deposition of pelagic cephalopod limestones (Hodderense Limestone Formation) of 4–15 m thickness. By early Asbian times limestone turbidite deposition resumed; the Pendleside Limestone Formation has a local thickness of up to 300 m. These turbidites represent the youngest Worston Shale Group stata. By late Asbian times the basin changed from a predominantly dysoxic environment to a euxinic one, represented by the black shales of the overlying Bowland Shale Group. This change is thought to have resulted from stratification of the basin water column as the



TEXT-FIG. 2. Stratigraphical distribution of ammonoids in the Worston Shale Group around Clitheroe; for explanation of locality numbers (in the lithostratigraphy column) see text; lithostratigraphical abbreviations as follows: BBS, Buckbanks Sandstone Member; BeL, Bellman Limestone Member; BSG, Bowland Shale Group; CgL, Chaigley Limestone Member; CLG, Chatburn Limestone Group; CoL, Coplow Limestone Member; LgM, Leagram Mudstone Member; LWL, Limekiln Wood Limestone Member; PhM, Phynis Mudstone Member; PQL, Peach Quarry Limestone Member; RGL, Rain Gill Limestone Member; TWS, Twiston Sandstone Member; ThtL, Thornton Limestone Member; WL, Whitemore Limestone Member. Vertical hatching denotes unconformity; (undiff.) = undifferentiated.

surrounding carbonate platforms became fringed by bioherms, interrupting circulation between the basin and the platform interiors.

AMMONOID ZONATION

Mid-Dinantian ammonoid zonation has recently been revised by Riley (1990a, 1991) and is summarized in Text-figure 3. Riley (1993) has also reviewed Dinantian classification in the British Isles. Problems with the definition and correlation of the Chadian Stage, which has its stratotype in the Craven Basin, have been investigated in detail by Riley (1995). Riley (1990a, 1991) redefined the Fascipericyclus-Ammonellipsites (FA) and Beyrichoceras (B) ammonoid zones; and proposed a new intervening zone, the Bollandites-Bollandoceras Zone. These zones are represented on all continents except Antarctica and South America. A range chart showing the distribution of ammonoid genera through the mid-Dinantian is given in Text-figure 3. As with late Viséan and Namurian ammonoid zones, international correlation is best achieved at the generic level. Few species are distributed internationally, and the sporadic nature of individual ammonoid occurrences in the mid-Dinantian renders the absolute stratigraphical range of individual species difficult to assess.

SERIES	TOUR	IAI	SIAN (part)		-	VI	SÉAN (part)		1 -
FORAM. ZONES GENERA	Cf3 (p	art)	Cf4 al	Cf4 α	2	Cf4 β–δ	Cf5	Cf6 α-β	Cf6 γ
Protocanites	\leftarrow	-					i		
Pericyclus									
Rotopericyclus	`								
Muensteroceras	`								
Eurites	`								
Muensteroceratoides									
Winchelloceras	\leftarrow				_				
	\leftarrow								
Irinoceras	`								
Merocanites		-							-
Fascipericyclus		-							
Neopericyclus									
Ammonellipsites									
Beyrichoceratoides									
Polaricyclus				-					
Stenocyclus		_			_				
Quasintoceras			_						
Aquilonites			_						
Intoceras				_					
Michiganites									
Eonomismoceras									
Helicocyclus							 		
Kozhimites									
Rhiphaeocyclus					_			 	
Hammatocyclus									
Bollandites				-					
				-	_				-
Bollandoceras		-			_				
Simmonoceras									
Parahammatocyclus Oxintocores					_		-	-	
Oxintoceras Dimorphoceras		-							
		-			_				- 3
Nomismoceras				-					
Beyrichoceras		-		-					L
Entogonites		_		-					
Girtyoceras		_		-				-	
Goniatites Prolecanites				-					
									\equiv
Metadimorphoceras	-			-					
Pronorites				├					
Praedarellites Arnsbergites	-	<u> </u>			_		-		
				1		D.D.	<u> </u>		,
AMMONOID ZONES	Pe		FA			ВВ	5	B	G

Fascipericyclus-Ammonellipsites (FA) Zone

This is an assemblage zone and is defined at the first entry of *Merocanites*. It overlies the *Pericyclus* Zone. The base lies within the late Tournaisian; however, its precise position in relation to conodont biostratigraphy remains to be proven. *Fascipericylus* enters low in the zone, followed by *Ammonellipsites* in the latest Tournaisian. The zone extends across the Tournaisian-Viséan boundary. *Fascipericyclus* becomes extinct in the early Viséan prior to the entry of *Helicocylus*, *Michiganites* and *Eonomismoceras* in the uppermost part of the zone. The top of the zone corresponds closely with the base of the Arundian Stage, but still lies within the late Chadian. It is marked by the extinction of *Ammonellipsites* and *Helicocyclus*.

Bollandites-Bollandoceras (BB) Zone

The base of this assemblage zone directly overlies the FA Zone and coincides with the first entry of *Bollandites*. Near the base, *Bollandoceras* enters and these two genera persist into the overlying *Beyrichoceras* Zone. *Parahammatocyclus* gen. nov. occurs in the middle part, with *Dimorphoceras* and *Nomismoceras* entering in the upper part of the zone. The top of the zone lies in the latest Holkerian and corresponds with the base of the overlying *Beyrichoceras* Zone.

Beyrichoceras (B) Zone

This is a genus zone based on the total range of *Beyrichoceras*. The zone is subdivided into three subzones (Riley 1990a), termed successively the *Entogonites* (B1), *Goniatites hudsoni* (B2a) and *Goniatites globostriatus* (B2b) subzones. The B2a and B2b subzones lie within the late Asbian. The *Beyrichoceras* Zone is succeeded directly by the *Goniatites* Zone.

AMMONOID DISTRIBUTION

This section summarizes the stratigraphical distribution of ammonoids within the local succession. The areal distribution of localities is illustrated in Text-figure 1, and the stratigraphical occurrence of the assemblages is summarized in Text-figure 2. Where relevant, BGS specimen numbers are quoted in parentheses. Locality numbers are listed in the locality index at the end of this section.

Unlike the late Viséan and Namurian ammonoids in northern England, Worston Shale Group assemblages are not restricted to discrete marine black shale bands, known as marine bands, but occur in a variety of lithologies in a totally marine sequence. Ammonoids are rare in the Clitheroe Limestone Formation and are always associated with Waulsortian limestones. Ammonoids become abundant at the base of the Hodder Mudstone Formation, thus complementing the change from a carbonate ramp to the dysoxic hemipelagic environment associated with this formation. Despite their abundance, material preserved well enough for identification is sparse and limited to only a few localities where early diagenesis favoured uncrushed preservation. Such occurrences are associated with early pyrite replacement, and with carbonate-rich environments such as in boulder beds and limestone turbidites. The most widespread horizon where uncrushed faunas abound is in the cephalopod limestones represented by the succeeding Hodderense Limestone Formation. The following account documents the occurrence of identifiable faunas in the Worston Shale Group.

Clitheroe Limestone Formation

The ammonoid assemblages in this formation lie in the middle part of the FA Zone, the early part of the zone being indicated in the upper part of the underlying Chatburn Limestone Group by the

TEXT-FIG. 3. Composite range chart for mid-Dinantian ammonoid genera adapted from Riley (1991). The genera *Atratoceras*, *Terektytes* and *Torulites* are omitted since their stratigraphical ranges are poorly known. Abbreviations; *Gon. = Goniatites* Zone, *Peri. = Pericyclus* Zone. Arrows indicate extension of ranges; broken horizontal lines denote uncertainty.

occurrence of *Merocanites similis*. Several genera have not yet been recorded, including *Ammonellipsites*, *Fascipericyclus* and *Merocanites*. Their absence is attributed to the rarity of ammonoids in this formation.

Coplow Limestone Member. This is the oldest fauna known in the western part of the Craven Basin. Archival material in the BGS collection includes a badly corroded example of *Dzhaprakoceras* sp. (BGS 86246) from Coplow Quarry, Clitheroe and a fragmentary body chamber of a gerontic *Rotopericyclus* cf. rotuliformis (BGS W 551) collected during the last century from the River Hodder at Knowlemere, near Newton. *Dzhaprakoceras* cf. djaprakense was recovered during the present study from Leagram Knott (locality 1) near Chipping.

Bellman Limestone Member. The strata between the Coplow and Bellman Limestone members have not yielded ammonoids. Ammonoids were noted from Sykes Knoll (locality 2), near Downham (Ramsbottom in Earp et al. 1961); the revised fauna comprises Beyrichoceratoides aff. komiorum, Dzhaprakoceras hispanicum and D. subglobosum. Muensteroceratoides cf. aksuensis has been found from The Ridge (SD 7690 4330), near Worston, and at Whitemore Knott (SD 6442 4755), near Whitewell, by A. Cronshaw (Keele University, pers. comm. 1981) and J. Saville (Airton, pers. comm. 1983), respectively.

Hodder Mudstone Formation

This formation encompasses the upper part of the FA Zone through to the upper part of the BB Zone. Crushed merocanitids and muensteroceratids *sensu lato* are common throughout the formation; the following account is restricted to identifiable faunas.

Phynis Mudstone Member. Ammonoids are poorly preserved in this member, but include a specimen of *Fascipericyclus* sp. from a temporary exposure (1977) at Downham Mill (SD 7895 4530) collected by R. F. Grayson, P. W. Phillips and the late S. Westhead (pers. comm. 1977).

Whitemore Limestone Member. As in the underlying Phynis Mudstone Member ammonoids are common but poorly preserved. Faunas have been recovered from several localities including Barn Gill (SD 7210 5355), Slaidburn and an unnamed stream (SD 6480 4793) near Whitemore Knott. Ammonellipsites sp. ex gr. kochi is encountered for the first time and is accompanied by indeterminate muensteroceratids, Merocanites sp. and Eonomismoceras sp.

Limekiln Wood Limestone Member. Ramsbottom in Earp et al. (1961, p. 179) recorded Muensteroceras cf. euryomphalum from Bellman Quarry, near Clitheroe; this specimen (BGS 86247) is too corroded for specific identification, but may possibly represent a species of Eurites or Muensteroceratoides. Ammonellipsites cf. clitheroensis sp. nov., large fragments of Dzhaprakoceras sp., D. deflexum and Merocanites sp. occur in the adjacent Salthill Quarry (locality 3a).

Leagram Mudstone Member. A temporary exposure (1978–79) of a boulder bed at the base of the Leagram Mudstones at Bellman Quarry (locality 4a) yielded the following ammonoids: Ammonellipsites clitheroensis sp. nov., Beyrichoceratoides inflatus sp. nov., Dzhaprakoceras bellmanense sp. nov., D. deflexum, D. levis sp. nov., D. paracatena sp. nov., D. undulatum sp. nov., D. westheadi sp. nov. Merocanites sp. and Rotopericyclus postremus sp. nov. The overlying 3 m of mudstone at Bellman and Salthill quarries (localities 4b and 3b respectively) yielded the following components, in addition to those above; Dzhaprakoceras catena sp. nov., Eonomismoceras planum sp. nov., Helicocyclus divergens sp. nov., and Michiganites sp. Higher beds in the Leagram Mudstone are often crowded with poorly preserved merocanitides ('compressus beds' of Parkinson 1926) and indeterminate muensteroceratids. Poorly preserved specimens of Ammonellipsites sp. ex gr. kochi have been recovered from the Leagram Mudstones in Standen Brook, at Primrose Bridge (SD 7390 4070) near Clitheroe, by R. F. Grayson (pers. comm. 1977), and several specimens, possibly of Helicocyclus sp. have been recorded from Worston Brook (Parkinson 1926, p. 211; later

identified as *Pericyclus djaprakensis* and *P. trapezoidalis* by Ramsbottom *in* Earp *et al.* 1961, p. 178). Crushed, poorly preserved *Helicocyclus* sp. occur also in a stream section (SD 6945 4835) at Ashnott, 25 m below the first entry of primitive archaediscid foraminiferans; this section is thought to overlie directly the Leagram Mudstone Member.

Chaigley Limestone Member and equivalent strata. Well preserved ammonoid faunas have been recovered from two levels at approximately 30 m and 100 m above the base of the Chaigley Limestone Member. Both these occurrences coincide with Arundian foraminiferal assemblages dominated by Uralodiscus spp. The lower fauna, from an exposure in the River Hodder at Papermill Wood (locality 5), consists of Parahammatocyclus chaigleyensis sp. nov. and Dzhaprakoceras hispanicum. The higher fauna, from an exposure in the River Hodder near Buck Hill (locality 6), contains Parahammatocyclus chaigleyensis sp. nov. and Bollandoceras sp. 2., Beyrichoceratoides sp. 2 and Bollandites? sp. At Bonstone Brook (locality 7). P. chaigleyensis sp. nov. occurs also in the Dunbarella Bed. In Leagram Brook (SD 6333 4512), near Chipping, indeterminate nomismoceratids (with arched venters) occur some 5 m below the Dunbarella Bed. In the upper part of the Hodder Mudstone Formation, identifiable faunas have been recovered from Crag Beck at Ashnott (locality 8), and include Beyrichoceratoides sp. 1, Bollandoceras sp. 1 and Merocanites sp. At Lower Hodder Bridge (SD 7045 3920), near Stonyhurst, crushed Bollandoceras sp. and Dimorphoceras sp. are scattered through the upper 50 m of the formation.

Hodderense Limestone Formation

This formation lies in the upper part of the BB Zone. Ammonoids are abundant and widespread; the most diverse fauna so far recovered has come from Leagram Brook (locality 11), near Chipping, and includes *Bollandites varians* sp. nov., *Bollandoceras hodderense*, *Dimorphoceras leagramense* sp. nov., *Nomismoceras rotiforme* and *Merocanites applanatus*.

Pendleside Limestone Formation

Much of this formation corresponds to the B1 Subzone, with only the basal part lying within the underlying BB Zone. The only identifiable fauna comes from the Mearley Limestone Member at locality 13; it includes poorly preserved *Merocanites* sp., together with *Bollandoceras micronotum* and *Nomismoceras rotiforme*, but the most significant component is the first *Beyrichoceras*, *B*. aff. *obtusum*.

LOCALITY INDEX

Clitheroe Limestone Formation

Locality 1. Leagram Knott, north-west corner of quarry (SD 6387 4473), 2·2 km north-east of Chipping Parish Church; coarse crinoidal grainstones in the upper part of the Coplow Limestone Member; early Chadian. Locality 2. Sykes Knoll, Quarry (SD 7993 4417), 0·7 km south-east of Hey House Farm, Downham; coarse packstones near the base of the Bellman Limestone Member; early Chadian.

Hodder Mudstone Formation

Locality 3. Salthill Quarry, south side (SD 7560 4254), 1.6 km east-north-east of Clitheroe Castle. Locality 3a is in coarse crinoidal packstones and mudstones in the upper 3 m of the Limekiln Wood Limestone Member. Locality 3b is in pyritic mudstones in the lower 4 m of the Leagram Mudstone Member; late Chadian. Locality 4. Bellman Quarry, south side, temporary exposure (1979) in a sewer trench (SD 7585 4260), 1.9 km east-north-east of Clitheroe Castle. Locality 4a is a boulder bed at the base of the Leagram Mudstone Member and locality 4b is in the overlying 3 m of these mudstones; late Chadian.

Locality 5. River Hodder, east bank (SD 6803 4310), 1.5 km west-south-west of Bashall Eaves; coarse bioclastic packstone (limestone turbidite) about 30 m above the base of the Chaigley Limestone Member; Arundian.

Locality 6. River Hodder, north bank (SD 6828 4275), 1.9 km south-west of Bashall Eaves; coarse erosive-based bioclastic packstone (limestone turbidite) about 100 m above the base of the Chaigley Limestone Member; Arundian.

Locality 7. Ashnott, Bonstone Brook, north bank (SD 6917 4876), 1.7 km south-west of Newton; nodules 2–3 m below the Dunbarella Bed; Arundian.

Locality 8. Ashnott, Crag Beck, west bank (SD 6912 4862), 1.95 km south-west of Newton; wackestone band in the upper 100 m of the Hodder Mudstone Formation; Holkerian.

Hodderense Limestone Formation

Locality 9. Ashnott, Crag Beck, both banks (SD 6896 4817), 2.4 km south-west of Newton; in the upper 3 m of the formation; Holkerian.

Locality 10. Leagram, Hell Clough, east bank (SD 6358 4630), 3·25 km north-north-east of Chipping Parish Church; wackestones and mudstones in lowest 2 m of the section; Holkerian.

Locality 11. Leagram Brook, both banks (SD 6322 4515), 2.05 km north-north-east of Chipping Parish Church; collections made throughout the section; Holkerian.

Locality 12. River Hodder, Great Falls, both banks (SD 7035 3999), 4·2 km south-west of Clitheroe Castle; collections made throughout the section; Holkerian.

Pendleside Limestone Formation

Locality 13. River Hodder, Great Falls, east bank, 50 m upstream from locality 12; coarse packstone in the basal leaf of the Mearley Limestone Member; ?Holkerian or earliest Asbian.

SYSTEMATIC PALAEONTOLOGY

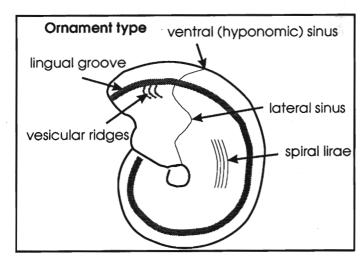
Unless stated otherwise all material is held with the British Geological Survey, Keyworth. Repositories of specimens referred to below are abbreviated as follows; BMNH, The Natural History Museum, London; PIN, Russian Academy of Sciences, Palaeontological Institute, Moscow, Russia; TSNIGR, All Russian Geological Research Institute, St Petersburg, Russia. Dimensions are given in millimetres. Dimensions marked with an asterisk are estimated. Scale bars on figures represent 1 mm. Terminology of the conch form follows Ruzhencev (1962); ornament terminology is summarized in Text-figure 4. The abbreviations D, H, U and W are standard for conch diameter, whorl height, umbilical diameter and conch width, respectively. Lithostratigraphical information on previously described species is given where known.

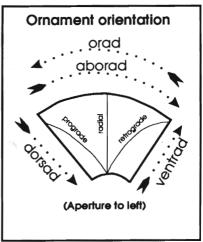
Because of the difficulties in recognition of the base of the Chadian Stage, discussed in Riley (1990b, 1993, 1995), stratigraphical intervals which may be of early Chadian or late Courceyan age are referred to as late Tournaisian. For the purposes of this account the base of the late Tournaisian is taken to correspond with the base of the Scaliognathus anchoralis Conodont Zone (Groessens 1975) or its equivalents. The base of the late Chadian corresponds to the base of the Viséan Series and follows Riley (1990b, 1993, 1995).

Order AGONIATIDA Ruzhencev, 1957 Suborder PROLECANITINA Miller and Furnish, 1954 Superfamily PROLECANITACEA Hyatt, 1884 Family PROLECANITIDAE Hyatt, 1884

Diagnosis (emended). The conch is evolute with moderate whorl expansion and rounded to rectangular whorls in axial section. The suture has 8–14 or more lobes, with a simple ventral lobe, except in gerontic *Dombaracanites* Ruzhencev, 1949. The other lobes are simple with acute or rounded bases. The dorsal lobe is simple and is bounded by an internal lobe, which may protrude ventrad of the umbilical margin.

Range. Tournaisian to Arnsbergian





TEXT-FIG. 4. Conch ornament terminology.

Remarks. The prolecanitids, in comparison with many contemporary ammonoids, show slow species turn-over and, because the conch form is conservative, classification is based on sutural characters. The scheme proposed by Weyer (1972b) is adopted here, with minor modification.

Subfamily PROLECANITINAE Hyatt, 1884

Diagnosis. Prolecanitids with a bell-shaped ventral lobe which is constricted orad. A siphuncular notch is present.

Range. Late Chadian to Brigantian.

Genus MICHIGANITES Ruzhencev, 1962

Type species. Goniatites marshallensis Winchell, 1862, p. 362.

Diagnosis. Prolecanitins with ten-lobed sutures.

Range. Late Chadian to late Asbian.

Remarks. The following species are included:

M. asiaticus (Karpinsky, 1896); Tien-Shan, Kirgizya, CIS, late Tournaisian.

M. discus (M'Coy, 1844); Clane, Ireland, late Tournaisian.

M. europeus (Kullmann, 1963); 'Marbre Griotte, CuIIy', Cantabria, northern Spain, late Tournaisian or early Viséan.

M. greeni (Miller in Miller and Faber, 1892); Indiana, USA, late Tournaisian or early Viséan.

M. hesteri (Moore, 1952); Pendleside Limestone Formation, Barnoldswick, north-west England, late Asbian.

M. marshallensis (Winchell, 1862); Marshall Sandstone, USA, mid-Dinantian.

M. rhipaeus Kusina, 1980; Nortnichsky Horizon, north Urals, Russian Federation, CIS, ?Arundian.

M. serpentinus (Phillips, 1836); Lower Bowland Shale Formation, Black Hall, Chipping, northwest England, late Asbian.

Michiganites sp. juv.

Plate 2, figure 3; Text-figure 5

Material. BGS Ro 5186; septate, adolescent, pyritic, internal mould; locality 3b.

Range. Late Chadian, upper FA Zone, Hodder Mudstone Formation, Leagram Mudstone Member.

Description. The conch is sub-platyconic with moderate whorl expansion and a wide umbilicus. The whorls are subquadrate in transverse section with flat flanks, a sharply rounded ventro-lateral region, and a feebly convex venter. The external ornament is not preserved. Internally, the conch surface is smooth. The ventral lobe is bell-shaped. Lobes U and U1 are acute and spatulate. Lobe U2 is feeble, acute, narrow and divergent-sided (orad).

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5186	18.0	6.7	7.6	4.5	0.25	0.42	0.37

Remarks. This specimen is closest to M. marshallensis and was the only Michiganites found during the study.

Subfamily PROTOCANITINAE Weyer, 1972a

Diagnosis. Prolecanitids with a divergent-sided (orad), V-shaped ventral lobe.

Range. Tournaisian to Arnsbergian.

Genus MEROCANITES Schindewolf, 1922

Type species. Ellipsolites compressus Sowerby, 1813, p. 84, plate 38.

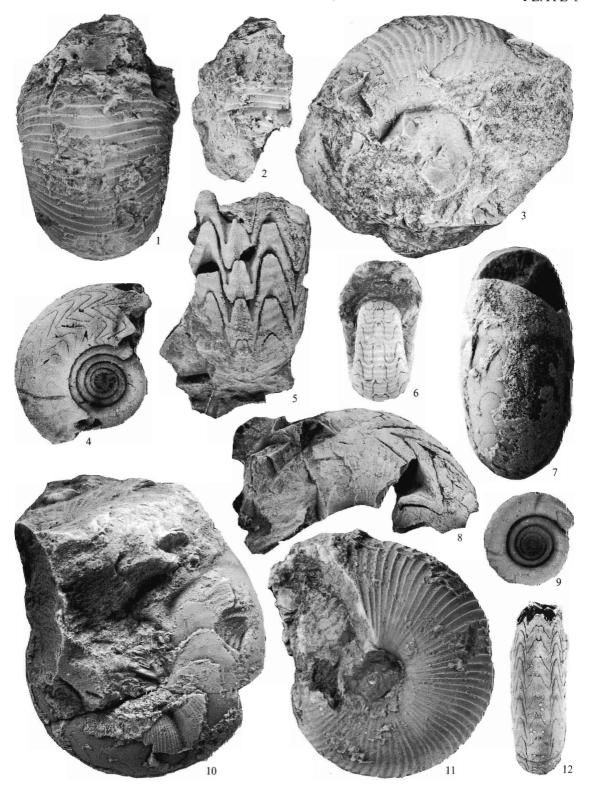
Remarks. Erdbachites Weyer, 1965, is a junior synonym of Merocanites because the type species of Merocanites has a divergent-sided ventral lobe (this can be seen on the type if immersed in silicone oil and illuminated with a strong light). Many species are inadequately described, or based on such

EXPLANATION OF PLATE 1

Figs 1–3, 5–6, 8, 11. Ammonellipsites clitheroensis sp. nov. (see also Pl. 2, fig. 4). 1, 11, holotype, Ro 864; locality 4a; ventral and lateral views of exfoliated adolescent conch; × 3. 2, paratype, Ro 867; locality 4a; ventral view of fragmentary, testiferous adult conch showing constriction; × 1. 3, paratype, Ro 863; locality 4a; lateral view of exfoliated adult conch; note loss of ribs low on flanks; × 1.5. 5, 8, paratype, Ro 865; locality 4a; ventral and lateral views of fragmentary, exfoliated adult conch; × 2. 6, paratype, Ro 866; locality 3b; oral view of exfoliated early adolescent conch; × 5.

Figs 4, 9, 12. Heliocyclus divergens sp. nov.; locality 3b. 4, 12, holotype, LZA 9422; lateral and ventral views of fragmentary, exfoliated adult conch; × 3. 9, paratype, LZA 9445; lateral view of exfoliated adolescent conch; × 3.25.

Figs 7, 10. Rotopericyclus postremus sp. nov.; locality 4a; holotype, Ro 933; ventral and lateral views of partly testiferous adolescent conch; × 3.



RILEY, Ammonellipsites, Helicocyclus, Rotopericyclus

poorly preserved material that it is impossible to apply the name to anything other than the type. Little is known about their variation, hence many species names may be synonyms. The following species are included:

- M. applanatus (Frech, 1899); Erdbach Limestone, Germany, mid-Dinantian.
- M. compressus (Sowerby, 1813); 'Blackrock Limestone', Cork, Ireland, mid-Dinantian.
- M. djaprakensis (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- M. drostei (Collinson, 1955); Broadhead Formation, Kentucky, USA, late Tournaisian.
- M. henslowi (Sowerby, 1821); Carboniferous Limestone, Isle of Man, UK, Asbian.
- M. houghtoni (Winchell, 1862); Marshall Sandstone, USA, mid-Dinantian.
- M. ogivalis (Pareyn, 1961); Banc de Mazzer, Sahara, north Africa, late Asbian.
- M. planorbis Delépine, 1941; Carboniferous slates, Morocco, N. Africa, Dinantian.
- M. similis (Crick, 1895); Haw Bank Limestones, Skipton, north-west England, late Tournaisian.
- M. subapplanatus (Smyth, 1951); Tober Colleen Formation, Rush, near Dublin, Ireland, late Chadian.
- M. subhenslowi (Wagner-Gentis, 1964); 'griotte limestone', Cantabria, northern Spain, late Tournaisian or early Viséan.
- M. tenuis (Sheng, 1984); middle of Donggulubasitao Formation, north-west China, Viséan.

Merocanites cf. applanatus (Frech, 1899)

Plate 2, figure 1, Text-figures 6-8

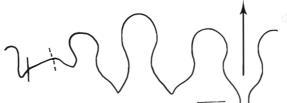
1889	Prolecanites ceratitoides (von Buch); Holzapfel, p. 240, pl. 3, figs 13–14.
1889	Prolecanites Henslowi (Sowerby); Holzapfel, p. 43, pl. 3, fig. 5.
1897	Prolecanites ceratitoides (von Buch); Foord and Crick, p. 251, fig. 120.
1899	Prolecanites applanatus Frech, pl. 46a, fig. 9.
1922	Merocanites Holzapfeli Frech; Schindewolf, p. 15, fig. 2b.
1925	Merocanites applanatus (Frech); Schmidt, p. 539, pl. 19, fig. 15.
1934	Merocanites applanatus (Holzapfel); Bisat, p. 305, pl. 24, fig. 1; text-fig. 22.
1935	Prolecanites (Merocanites) applanatus Frech; Böhm, pp. 126, 136, pl. 7, figs 1-4, 15; text-figs
	17–18.
1941	Merocanites applanatus (Frech); Delépine, p. 45, pl. 3, figs 5-6; text-fig. 6B.
1941	Prolecanites (Merocanites) applanatus (Frech); Schmidt, p. 150, pl. 19, fig. 1.
1951 <i>a</i>	Merocanites applanatus (Frech); Schindewolf, p. 49, pl. 3, figs 3–5.

EXPLANATION OF PLATE 2

- Fig. 1. Merocanites cf. applanatus (Frech, 1899); Ro 5196; locality 9; lateral view of exfoliated, slightly crushed adult conch; ×1.
- Fig. 2. Nomismoceras rotiforme (Phillips, 1836) (see also Pl. 6, fig. 15); Ro 2028; locality 13; lateral view of exfoliated adult conch; × 3.5.
- Fig. 3. Michiganites sp.; Ro 5182; locality 4b; lateral view of exfoliated early adolescent conch; ×2.
- Fig. 4. Ammonellipsites clitheroensis sp. nov. (see also Pl. 1, figs 1–3, 5–6, 8, 11); holotype, Ro 864; locality 4a; oral view of exfoliated adolescent conch; ×2.
- Figs 5–8. *Eonomismoceras planum* sp. nov. 5, paratype, Ro 876; locality 3b; lateral view of exfoliated juvenile conch; × 3·25. 6, paratype, Ro 894; locality 4b; lateral view of exfoliated, slightly crushed adult conch; × 2. 7–8, holotype, Ro 874; locality 3b; lateral and ventral views of exfoliated adult conch; × 3·5.
- Fig. 9. Merocanites quadrilobus sp. nov. (see also Pl. 3, fig. 5); paratype, Ro 5174; locality 4a; fragmentary, exfoliated late adolescent conch; ×1.
- Figs 10–11. *Dimorphoceras leagramense* sp. nov.; holotype, Ro 895; locality 11; lateral and ventral views of fragmentary, exfoliated adult conch; × 5.



RILEY, mid-Dinantian ammonoids



TEXT-FIG. 5. Michiganites sp.; Ro 5186; ventral and dorsal suture (part), D = 17.1, H = 6.2, U = 7.5, W = 4.5.

non 1961 Merocanites applanatus (Frech) var. bicarinatus Pareyn, pl. 2, figs 1–3, pl. 3, figs 6–9.

1963 Merocanites applanatus (Frech); Kullmann, p. 274, pl. 17, figs 1–2; text-fig. 2b.

Type specimen. Prolecanites applanatus, figured by Frech (1899, pl. 46a, fig. 9).

Material. More than 30, mainly distorted internal moulds preserved in micrite; from localities 4a, 9-12.

Range. Holkerian, Upper BB Zone; Hodderense Limestone Formation. Known also from the Erdbach Limestone, Germany and from the upper part of the Milldale Limestones, Staffordshire (Riley 1991).

Diagnosis. A Merocanites with rapid whorl expansion, early development of a subquadrate whorl section (trans.) and a deep, broad umbilical lobe (U2).

Description. The conch is serpenticonic with rapid whorl expansion. The flanks are flat, which together with the narrow ventrolateral shoulder and narrow, feebly convex venter, give a subquadrate appearance to the whorl in axial section after H=10 mm. The external ornament is not preserved and the internal ornament is smooth. The ventral lobe is narrow and V-shaped with sides that diverge orad. It is deep in juveniles, but becomes shallower in the adult, when it is not as deep as the umbilical lobe, U2, or the lateral lobe. The lateral lobe and the umbilical lobes, U3 and U2, are spatulate; U2 is comparatively broad and deep for this genus. All the saddles are rounded.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5263	81.0	26.5	33.0	17:0	0.21	0.41	0.33
Ro 5293	70.0	26.5	27.0	15.5	0.22	0.39	0.39
Ro 5257	25.5	8.7	9.9	6.8	0.27	0.39	0.34
Ro 5270	18.5	7.5	8.6	5.5	0.3	0.46	0.41
Ro 5281	12.2	4.5	5.1	3.7	0.3	0.42	0.37

Remarks. The types of *M. applanatus* were destroyed during World War Two, hence no adequate modern description is available. The Craven material appears to belong to this species on account of the early deepening of lobe U2, the relatively rapid whorl expansion and the early presence of a subquadrate axial whorl section.

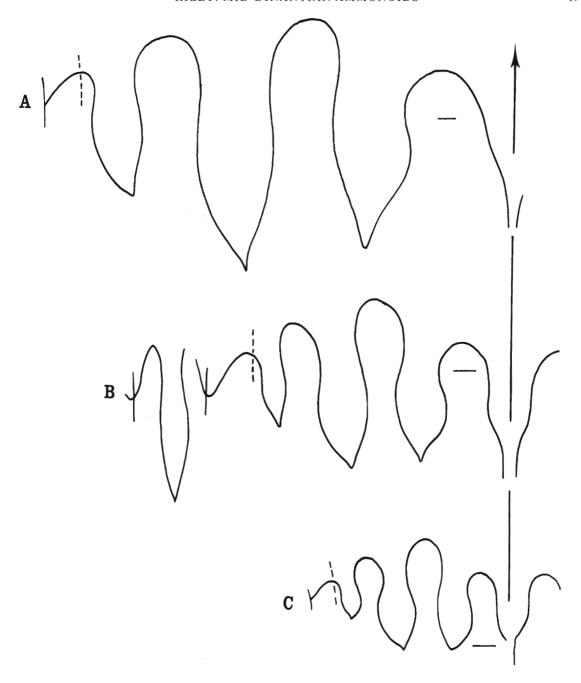
Merocanites quadrilobus sp. nov.

Plate 2, figure 9; Plate 3, figure 5; Text-figures 7–8

Derivation of name. For the four exposed lobes on the conch sides.

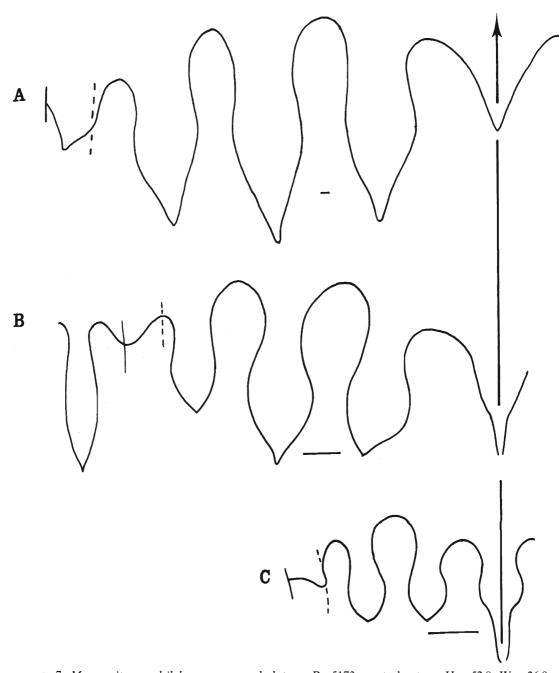
Material. Holotype Ro 5173 and paratypes Ro 5174–5175, all mainly exfoliated, fragmentary, septate conches preserved in micrite; from locality 4a.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation, Leagram Mudstone Formation.



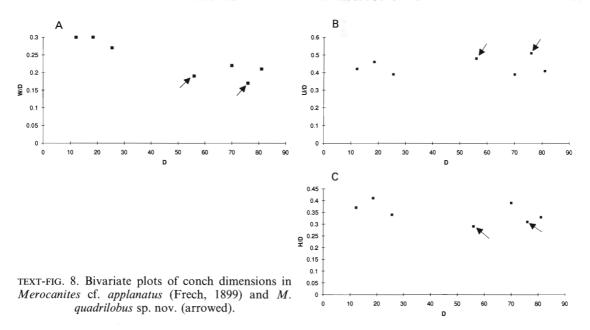
TEXT-FIG. 6. Merocanites cf. applanatus (Frech, 1899). A, Ro 5196; ventral suture, H = 23.0 (reversed image). B, Ro 5224; ventral suture, H = 9.5, W = 6.7. c, Ro 5218; ventral suture, H = 6.9, W = 5.0.

Description. The conch is serpenticonic with moderate whorl expansion. Initially the flanks are rounded, but they are nearly flat after H=15 mm. The venter is narrow and convex. The external ornament is not preserved and the internal ornament is smooth, without constrictions. The external suture shows a broad ventral lobe



TEXT-FIG. 7. Merocanites quadrilobus sp. nov. A, holotype, Ro 5173; ventral suture, H = 52.8, W = 26.9. B, paratype, LZA 9378; ventral and dorsal suture, H = 7.9 (reversed image). C, paratype, LZA 9373; ventral suture, D = 9.1, H = 3.4, U = 4.2, W = 2.5.

which is relatively shallow with sides that diverge orad. Both the lateral lobe and the umbilical lobes, U2 and U3, are deep, spatulate and broad. Characteristically, the umbilical lobe, U1, is exposed ventrad of the whorl suture during late ontogeny and has a concave dorsad wall.



Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5173		52.0		30.2			
Ro 5174	76.0	23.5	39.0	13.0	0.17	0.51	0.31
Ro 5175	56.0	16.5	27.0	10.5	0.19	0.48	0.29

Remarks. It is possible that forms refered to Prolecanites compressus by Parkinson (1926), from the lower part of the Hodder Mudstone Formation, belong to this species; however, none of his material is available for comparison. Crushed merocanitids abound in the lower part of the Hodder Mudstone Formation, but they do not show sutural details. The type of Merocanites compressus is distorted and the nature of the suture in the umbilical region is obscure. Only M. djaprakensis, M. drostei and M. houghtoni have an umbilical lobe, U1, with a concave ventrad wall. The possibility that these species are immature M. quadrilobus sp. nov. is discounted because they are described from conches of similar diameter to the new species. The presence of an exposed umbilical lobe, U1, is unique and converges with the eocanitid genus Cantabricanites Weyer, 1965, which is known from Brigantian and Pendleian strata. In Cantabricanites, however, the four exposed lobes are produced by insertion of a new umbilical lobe, U4, rather than ventrad migration of umbilical lobe, U1.

Order Goniatitida Hyatt, 1884 Suborder Goniatitina Hyatt, 1884 Superfamily Dimorphocerataceae Hyatt, 1884

Diagnosis (emended). The conch is involute, with a small umbilicus, and varies from suboxyconic to pachyconic. The early whorls lack a serpenticonic stage. The external ornament is feeble and consists of bisinuous transverse striae with well developed linguae. Spiral ornament is sometimes present. The ventral lobe is wide and divergent-sided (orad) and usually has subdivided prongs. The lateral lobe is often subdivided, and is sometimes serrate.

Range. Holkerian to Langsettian.

Remarks. The following families are included:

Anthracoceratidae Plummer and Scott, 1937; Pendleian to Langsettian.

Berkoceratidae Librovitch, 1957; Brigantian to Arnsbergian.

Dimorphoceratidae Hyatt, 1884; Holkerian to Kinderscoutian.

The Girtyoceratidae Wedekind, 1918 and Gonioloboceratidae Spath, 1934, formerly placed in the Dimorphocerataceae, are transferred herein to the newly erected superfamily Girtyocerataceae; hence the diagnosis has been emended.

Family DIMORPHOCERATIDAE Hyatt, 1884 Subfamily DIMORPHOCERATINAE Hyatt, 1884 Genus DIMORPHOCERAS Hyatt, 1884

Type species. Glyphioceras gilbertsoni Phillips, 1836.

Diagnosis. Dimorphoceratins with bipartite ventral lobe prongs.

Range. Holkerian to mid-Brigantian.

Remarks. The following species are included:

- D. algens Gordon, 1957; Alapah Limestone, Kiruktagiak River Alaska, Asbian or Brigantian.
- D. brancoi Holzapfel, 1889; Erdbach Limestone, Germany, mid-Dinantian.
- D. dnieperense Kusina in Kusina and Poletaev, 1991; Zone XIIa, Solochovskaya Suite, Donets Basin, Ukraine, CIS, ?Asbian.
- D. gilbertsoni (Phillips, 1836); Bolland, north-west England, Viséan.
- D. leagramense sp. nov.; Great Falls Limestone Formation, Leagram, north-west England, late Holkerian.
- D. leitriemense Moore, 1958; Dough Mountain, Co. Leitrim, Ireland, Brigantian.
- D. kathleenae Moore, 1936; Lower Bowland Shale Formation, Dinckley, north-west England, Brigantian.

Dimorphoceras leagramense sp. nov.

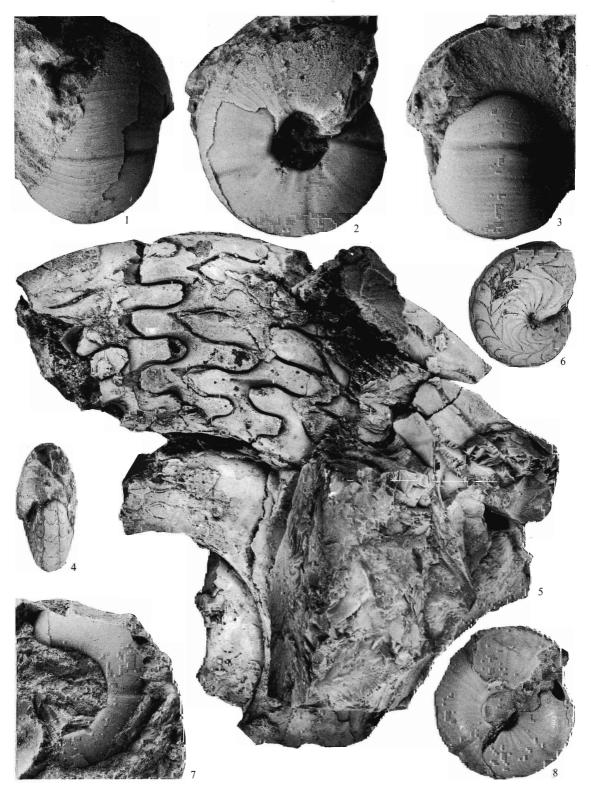
Plate 2, figures 10-11; Text-figures 9-10

Derivation of name. From Leagram, the type locality.

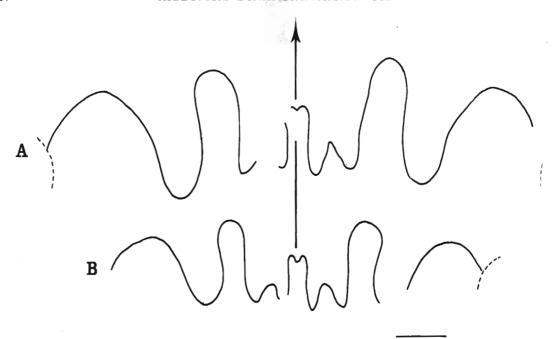
Material. Holotype, Ro 895, paratypes Ro 896–905, internal micritic casts; from locality 11.

EXPLANATION OF PLATE 3

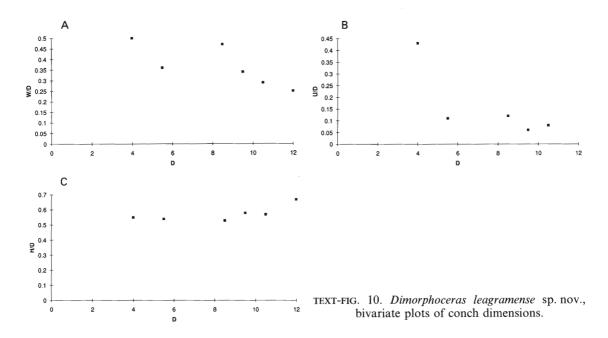
- Figs 1–3, 8. Parahammatocyclus chaigleyensis gen. et sp. nov. (see also Pl. 6, figs 1, 4, 12; Pl. 7, fig. 8). 1–3, holotype, Ro 5127; locality 6; lateral, oral and ventral views of partly testiferous, late adolescent conch; × 4. 8, paratype, Ro 3723; locality 7; exfoliated, partly crushed adolescent conch; × 2.75.
- Figs 4, 6. *Dzhaprakoceras levis* sp. nov. (see also Pl. 4, figs 2, 4); holotype, LZA 9399; oral and lateral views of exfoliated adolescent conch; × 2·75.
- Fig. 5. Merocanites quadrilobus sp. nov. (see also Pl. 2, fig. 9); holotype, Ro 5173; locality 4a; lateral view of fragmentary exfoliated adult conch; note exposure of umbilical lobe U1; ×1.
- Fig. 7. Bollandites? sp.; Ro 5076a; locality 6; lateral view of fragmentary, exfoliated late adolescent conch; × 4.



RILEY, mid-Dinantian ammonoids



TEXT-FIG. 9. Dimorphoceras leagramense sp. nov.; holotype, Ro 895. A, ventral suture, D = 8.2, H = 4.4, U = 1.8, W = 3.9. B, ventral suture, H = 2.9, W = 3.1.



Range. Holkerian, upper BB Zone; upper part of the Hodder Mudstone Formation and Hodderense Limestone Formation.

Diagnosis. A Dimorphoceras with rounded lobes and a high, narrow median saddle.

Description. The conch is suboxyconic and laterally compressed, but pachyconic below the third whorl. The umbilicus is very small. The conch flanks are broad and slightly convex, converging into a narrow, rounded venter. The external ornament is not preserved, but the internal ornament shows feeble, bisinuous undulations with well developed lateral and hyponomic sinuses and linguae, sometimes forming a shallow lingual groove high on the flanks. The external suture shows a ventral lobe which is bipartite, with rounded lobes and subacute subsidiary saddles. There is a narrow, high median saddle, about half the ventral lobe depth and twice the subsidiary saddle height. The lateral saddle is rounded and inflated, especially in juveniles (H = 3). The lateral lobe is deep and rounded with steep walls. The umbilical saddle is high and broad.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 889	12.0	8.0	0.0	3.0	0.25		0.67
Ro 898	10.5	6.0	0.8	3.0	0.29	0.08	0.57
Ro 902	9.5	5.5	0.6	3.2	0.34	0.06	0.58
Ro 895	8.5	4.5	1.0	4.0	0.47	0.12	0.53
Ro 903	5.5	3.0	0.6	2.0	0.36	0.11	0.54
Ro 905	4.0	2.2	0.5	2.0	0.50	0.43	0.55

Remarks. D. leagramense sp. nov. is most similar to D. gilbertsoni from which it differs by its more laterally compressed conch and higher median saddle.

Superfamily GIRTYOCERATACEAE superfam. nov.

Diagnosis. Goniatitinids with oxyconic to subsphaeroconic conches and serpenticonic early whorls. The external ornament usually has both spiral and transverse elements; often with a well developed lingua and sometimes a lingual groove. Simple or dichotomous ribs may be present on the flanks and/or umbilical region. The ventral lobe is broad and divergent-sided (orad) and often has a well developed median saddle. The lateral saddle is rounded to subacute. The lateral lobe is acute.

Range. Late Chadian to Arnsbergian.

Remarks. The following families are included:

Girtyoceratidae Wedekind, 1918; Holkerian to Arnsbergian.

Eogonioloboceratidae Ruzhencev and Bogoslovskaya, 1978; late Chadian to Arnsbergian.

Family GIRTYOCERATIDAE Wedekind, 1918

Diagnosis (emended). The conch is oxyconic to subsphaeroconic. The external ornament has transverse and often spiral elements. The transverse ornament is usually strongly bisinuous. Simple or dichotomous ribs, which may fail during late ontogeny, are present in some stratigraphically young species, as is a lingual groove. The ventral lobe is broad and divergent-sided (orad). The lateral saddle is usually sub-acute and may be rounded. The lateral lobe is acute.

Range. Holkerian to Arnsbergian.

Remarks. The following genera are included:

Beyrichoceras Foord, 1903; Holkerian to Asbian.

Cousteauceras Korn, 1988; late Brigantian to early Pendleian.

Edmooroceras Elias, 1956; early Pendleian.

Eumorphoceras Girty, 1909; late Brigantian to Arnsbergian.

Girtyoceras Wedekind, 1918; Asbian to Brigantian.

Peytonoceras Saunders, 1966; Arnsbergian.

Sulcogirtyoceras Ruzhencev, 1960; late Brigantian.

Torulites Kusina and Yatskov, 1987; ?Arundian.

Tumulites McCaleb, Quinn and Furnish, 1964; Pendleian.

Winchelloceras Ruzhencev, 1965 may be a girtyoceratid; however, its taxonomic position is unclear since the early whorls are undescribed. *Pseudogirtyoceras* Wagner-Gentis in Higgins and Wagner-Gentis 1982 is here considered to be a junior synonym of the eogonioloboceratid *Atratoceras* Librovitch, 1957. Kusina (1987) placed *Beyrichoceras* in the Anthracoceratidae Plummer and Scott, 1937, but the presence of spiral ornament, evolute early whorls and sutural characters of *Beyrichoceras* do not support this.

Genus BEYRICHOCHERAS Foord, 1903

Type species. Goniatites obtusus Phillips, 1836, plate 19, figures 10–11.

Diagnosis. Moderate-sized to large girtyoceratids with subsphaeroconic to pachyconic conches and rounded to rectangular whorls in axial section. The external ornament consists of transverse bisinuous striae, often with a well developed lingua. Spiral ornament is usually limited to the lingua, but may extend over the flanks and venter. Internal constrictions are sometimes present. The ventral lobe is divergent-sided (orad), with flexed sides and acute prongs which diverge abroad. The lateral saddle is acute to subacute. The lateral lobe is acute with flexed sides.

Range. Holkerian to Asbian.

Remarks. The following species are included:

- B. araneum Nicolaus, 1963; Goα2, Goniatitenknollen, Eschenberg bei Hillerhausen, Germany, late Asbian.
- B. constans Kusina, 1987; Viséan Complex, Novaya Zemlya, CIS, Asbian.
- B. delicatum Bisat, 1924; 'reef limestone', Treak Cliff, northern England, Asbian.
- B. mempeli Schmidt, 1941; CuIIδ, or CuIIIα, Herborn, Germany, late Asbian.
- B. obtusum (Phillips, 1836); Lower Bowland Shale Formation, Black Hall, Chipping, northwest England, late Asbian.
- B. parkinsoni Bisat, 1934; Pendleside Limestone Formation, Slaidburn, north-west England, Asbian
- B. rectangularum Bisat, 1934; ?Malham Formation, Grassington, north-west England, Asbian.
- B. tardum Ruzhencev, 1966; limestones in Kzyl-Shin Gorge, south Urals, Russian Federation, late Asbian.
- B. tenue Kusina, 1987; Viséan Complex, Novaya Zemlya, CIS, Asbian.
- B. vesiculiferum (de Koninck, 1880); Visé, Belgium, Asbian.

Beyrichoceras bashachense Popov, 1965, B. furiense Padget, 1952 and B. mackellari Campbell, Brown and Coleman, 1983 are here placed in Bollandoceras. Beyrichoceras trevallynense Brown, Campbell and Roberts 1964, is here transferred to Bollandites. The generic status of Beyrichoceras bootibootiense Campbell, Brown and Coleman, 1983 is unclear; however, the lack of serpenticonic



TEXT-FIG. 11. Beyrichoceras aff. obtusum (Phillips, 1836); Ro 2160; ventral suture, H = 14·4.

early whorls excludes it from *Beyrichoceras*. The conch and suture suggest affinity with *Dzhaprakoceras* or *Beyrichoceratoides*.

Beyrichoceras aff. obtusum (Phillips, 1836)

Plate 8, figure 12: Text-figure 11

	1836	Goniatites obtusus Phillips, p. 233, pl. 19, figs 10-11, ?13 (non fig. 12, fidé Foord 1903).
	1880	Goniatites obtusus Phillips; de Koninck, p. 104, pl. 47, fig. 10-10a (non pl. 46, fig. 3, fidé
		Delépine 1940).
	1897	Glyphioceras obtusum (Phillips); Foord and Crick, p. 169, text-fig. 79.
	1899	?Glyphioceras obtusum (Phillips); Frech, pl. 46a, fig. 3.
	1903	Glyphioceras (Beyrichoceras) obtusum (Phillips); Foord, p. 163, pl. 42, fig. 7a-c (non figs
		8a-b, 9a-b).
	1924	Beyrichoceras obtusum (Phillips), Bisat, p. 86, pl. 9, fig. 3.
	1940a	Beyrichoceras obtusum (Phillips); Delépine, p. 69, pl. 4, figs 3-8; text-fig. 16.
non?	1941	Beyrichoceras obtusum (Phillips); Delépine, p. 64, pl. 3, figs 33–35.
	1952	Beyrichoceras obtusum (Phillips); Bisat, p. 170, fig. 2d (non 2h)

Lectotype. Designated by Bisat (1952), BMNH C247a, figured by Phillips (1836, pl. 19, figs 10–11).

Material. Specimen Ro 2160, fragmentary exfoliated, septate late adolescent preserved in packstone; from locality 13.

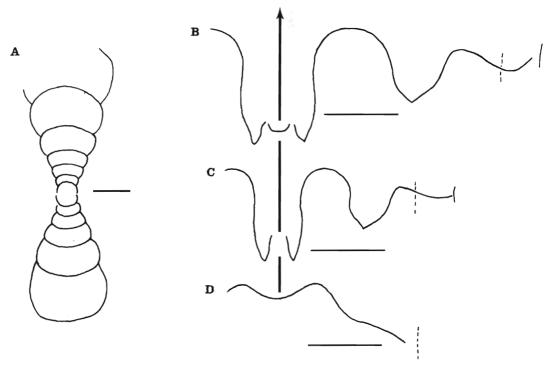
Range. B. obtusum sensu stricto occurs only in the B2a and B2b subzones (late Asbian). In the Craven Basin this interval lies in the Lower Bowland Shale Formation. Elsewhere in Europe this species is known from the same subzones.

Description. The conch is pachyconic with feebly convex sides, a sharply rounded ventro-lateral area and a narrow convex venter, giving an almost subquadrate axial whorl section. The external ornament is not preserved. The internal ornament consists of feeble vescicular ridges in the ventrolateral region. Two feeble bisinuous constrictions are present in the quarter of the whorl preserved. The ventral lobe has slightly divergent acute prongs, flexed divergent sides (orad). The lateral saddle is narrow and sharply rounded. The lateral lobe is broad, with flexed sides and is the same depth as the ventral lobe.

Remarks. This specimen is close to the type in the development of constrictions, widely spaced septa and conch form; however, the median saddle is higher than in the type.

Superfamily NOMISMOCERATACEAE Librovitch, 1957

Diagnosis (emended). The conch is small, laterally compressed, serpenticonic to discoconic with slow whorl expansion. The whorls are often arched at maturity, when the venter may be carinate, acute of flattened. Coiling is usually regular, but is quadrangular in the Entongonitidae. The external ornament is transverse and often strongly bisinuous; ribs may be present (Entogonitidae). The ventral lob is bipartite with simple, rounded to acute prongs and parallel to divergent (orad) sides. The lateral saddle is rounded. The lateral lobe is often narrow and may be rounded or acute.



Range. Late Chadian to Arnsbergian.

Remarks. Two families are included:

Nomismoceratidae Librovitch, 1957; late Chadian to Yeadonian. Entogonitidae Ruzhencev and Bogoslovskaya, 1971; Asbian.

Family NOMISMOCERATIDAE Librovitch, 1957

Diagnosis. Nomismocerataceans with regular coiling.

Range. Late Chadian to Yeadonian.

Remarks. The following genera are included:

Baschkirites Librovitch, 1957; Marsdenian to Yeadonian.

Beleutoceras Rushencev and Bogoslovskaya 1971; Pendleian.

Cavilentia Ruzhencev and Bogoslovskaya, 1971; Arnsbergian.

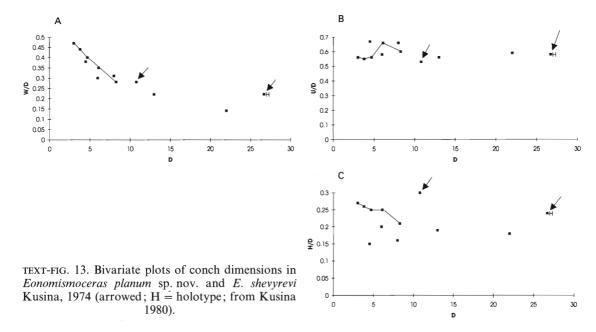
Eonomismoceras Kusina, 1974; late Chadian.

Hudsonoceras Moore, 1946; Alportian to Kinderscoutian.

Nomismoceras Hyatt, 1884; Holkerian to Brigantian.

Pseudonomismoceras Frech, 1899; ?early Viséan.

Simmonoceras Kusina, 1974; Arundian.



Genus EONOMISMOCERAS Kusina, 1974

Type species. Eonomismoceras shevyrevi Kusina, 1974, plate 4, figures 1–2.

Diagnosis (emended). Nomismoceratides with serpenticonic, non-carinate conches and strongly bisinuous transverse ornament. Constructions may be present in early ontogeny. The ventral lobe is narrow with sub-parallel sides, often constricted orad. The lateral saddle is broad and well rounded. The lateral lobe is narrow and acute.

Range. Late Chadian.

Remarks. The following species are included:

E. shevyrevi Kusina, 1974; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.

E. planum sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.

Eonomismoceras planum sp. nov.

Plate 2, figures 5-8; Text-figures 12-13

Derivation of name. From the Latin planus (= flat), with reference to the dorso-ventral flattening of the whorls.

Material. Holotype Ro 874, paratypes Ro 875–894, pyritic internal moulds; from locality 3b.

Range. Late Chadian, Upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. An Eonomismoceras with a steep-sided, broad-based, acute lateral lobe.

Description. The conch is serpenticonic with slow whorl expansion, and dorso-ventrally compressed whorls which have narrow, convex flanks and a broad convex venter. The external ornament is unknown. The internal ornament consists of feeble, bisinuous ridges which show a shallow lateral sinus and a deep, V-shaped hyponomic sinus. Constructions, which follow the external ornament, may be present in juveniles. The ventral lobe is deep, with sub-parallel sides which constrict slightly orad. The lateral saddle is broad and well

rounded. The lateral lobe is shallow, about three-quarters of the ventral lobe depth, broad-based and acute with sub-parallel sides.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 894	22.0	4.0	13.0	3.0	0.14	0.59	0.18
Ro 874	13.0	2.5	7.3	2.8	0.22	0.56	0.19
Ro 876	8.0	1.3	5.3	2.5	0.31	0.66	0.16
Ro 880	6.0	1.2	3.5	1.8	0.30	0.58	0.20
Ro 878	4.5	0.7	3.0	1.7	0.38	0.67	0.15
Ro 893	8.3	1.8	5.0	2.33	0.28	0.60	0.21
Ro 893	6.1	1.5	4.0	2.16	0.35	0.66	0.25
Ro 893	4.7	1.2	2.6	1.89	0.40	0.56	0.25
Ro 893	3.8	1.0	2.1	1.66	0.44	0.55	0.26
Ro 893	3.0	0.8	1.7	1.4	0.47	0.56	0.27

Remarks. The most obvious feature which distinguishes this species from E. shevyrevi is the sub-parallel-sided, broad-based, acute lateral lobe. E. shevyrevi is further distinguished by possessing a higher whorl section at comparable diameters.

Genus NOMISMOCERAS Hyatt, 1884

Type species. Goniatites spirorbis Phillips, 1836, plate 20, figures 51-55.

Diagnosis. (based on Ruzhencev and Bogoslovskaya 1971). The conch is small discoconic and laterally compressed. The whorls are moderately arched at maturity, when the venter may become narrow and flat. The external ornament is transverse and bisinuous with a strong lingua, which may form a lingual groove. The ventral lobe has divergent sides (orad) and rounded prongs. The lateral saddle is broad and well rounded. The lateral lobe is shallow, narrow and rounded.

Range. Holkerian to Brigantian.

Remarks. The following species are included:

- N. frechi Schmidt, 1925; Weitengrund, Germany, Viséan.
- N. germanicum Schmidt, 1925; Bredelar, Germany, late Asbian.
- N. rotiforme (Phillips, 1836); Bolland, north-west England, Viséan.
- N. spirorbis (Phillips, 1836); Lower Bowland Shale Formation, Black Hall, Chipping, northwest England, late Asbian.
- N. vittigerum (Phillip, 1836); Bolland, north-west England, Viséan.

Nomismoceras? spiratissimum Holzapfel, 1889, N.? infracostatum Schindewolf, 1951a and N.? pseudocyclus Campbell, Brown and Coleman, 1983 are too poorly known to be certain of their inclusion in this genus.

Nomismoceras rotiforme (Phillips, 1836)

Plate 2, figure 2; Plate 6, figure 15

1836	Goniatites rotiformis Phillips, p. 237, pl. 2, fig. 56–58.
?1844	Ammonites vittiger Phillips; de Koninck, p. 582, pl. 50, fig. 5a-c.
1880	Goniatites rotiformis Phillips; de Koninck, p. 114, pl. 50, fig. 16.
1897	Nomismoceras rotiforme (Phillips); Foord and Crick, p. 215, text-fig. 103.
on 1901	Nomismoceras rotiforme (Phillips): Frech. pl. 46a. fig. 6a-f.

Holotype. BMNH C.271a, Phillips (1836, pl. 20, fig. 56).

Material. More than 30 exfoliated, mainly fragmentary, distorted, micritic conches representing all growth stages; localities 9–13.

Range. Holkerian, upper BB Zone to B1 Subzone; Hodder Mudstone Formation to Pendleside Limestone Formation.

Description. The conch is small, discoconic and laterally compressed, with a carinate venter at maturity. Whorl expansion is slow. The external ornament is not preserved. The internal ornament consists of feeble, regularly spaced, transverse undulations on the middle of the flanks. The details of the suture are unclear; however, the ventral lobe is broad and strongly divergent-sided (orad). The lateral saddle is broad and rounded. The lateral lobe is shallow, narrow and rounded.

Remarks. This species is distinguished from the similar N. vittigerum by the presence of radial undulations on the flanks,

Superfamily Pericyclaceae Hyatt, 1900 Family MUENSTEROCERATIDAE Librovitch, 1957

Diagnosis. The conch is of medium to large size, subsphaeroconic to discoconic, sometimes suboxyconic, with a medium-sized to closed umbilicus. The early whorls are serpenticonic to cadiconic, never involute. The usually feeble external ornament consists of transverse striae, which may lack or have one or more sinuses; weak spiral ornament is rarely present. The ventral lobe has parallel, divergent (orad) or convergent (orad) sides. The lateral saddle is rounded. The lateral lobe is acute.

Range. Courceyan to Pendleian.

Remarks. The following genera are included:

Beyrichoceratoides Bisat, 1924; late Chadian to Arnsbergian.

Bollandoceras Bisat, 1952; Arundian to Brigantian.

Dzhaprakoceras Popov, 1965; late Tournaisian to Arundian.

Eurites Kusina, 1973; late Tournaisian.

Kozhimites Kusina, 1971; late Chadian.

Muensteroceras Hyatt, 1884; Tournaisian.

Muensteroceratoides Popov, 1965; late Tournaisian.

Terektytes Librovitch, 1957; late Tournaisian or early Viséan.

Kozhimites, originally of the monospecific family Kozhimitidae Kusina, 1971, is close to *Dzhaprakoceras* and therefore separation into a separate family is here considered unjustified. The following genera require further comment.

- (1) Beyrichoceras Foord, 1903. The combination of well developed spiral ornament, strongly bisinuous transverse ornament, subacute lateral saddles and a prolonged juvenile serpenticonic stage places this genus within the Girtyoceratidae.
- (2) *Bollandites* Bisat, 1952. Internal ribbing of the early whorls and in some cases the adolescent whorls, together with sutural characters places *Bollandites* within the Pericyclidae.
- (3) Cluthoceras Currie, 1954. Examination of type and paratype material reveals that Cluthoceras falls within Beyrichoceratoides as now defined. This extends the range of Beyrichoceratoides beyond the Pendleian (the horizon of B. invaginatum) into the Arnsbergian (Calmy Limestone).
- (4) Eoglyphioceras Brüning, 1923. This genus was given priority over Beyrichoceratoides by Korn (1988); however, when originally erected, no types species was given (unlike Beyrichoceratoides) and it appears that the genus was defined by amalgamation of Beyrichoceras and Muensteroceras. This lack of clarity in the original description renders Eoglyphioceras invalid.

Genus BEYRICHOCERATOIDES Bisat, 1924

Type species. Goniatites implicatus Phillips, 1836.

Diagnosis. The conch is involute, small to medium-sized and varies from discoconic to subsphaeroconic and lacks a juvenile serpenticonic stage. The umbilicus is of small to moderate size. The venter is rounded, but may become subquadrate late in ontogeny. The external ornament consists of simple transverse striae which become bisinuous during late adolescence. The internal surface of the conch may develop strong undulations (and more rarely ribs) particularly on the flanks, during late adolescence. Constructions are present; they run parallel with the external ornament but often fail during adolescence. The ventral lobe is narrow, parallel to sub-parallel sided, often constricted orad, sometimes lyre-shaped and may be asymmetrical. It has acute prongs and a low median saddle. The lateral saddle is rounded and often broad, never inflated. The lateral lobe is narrow and acute.

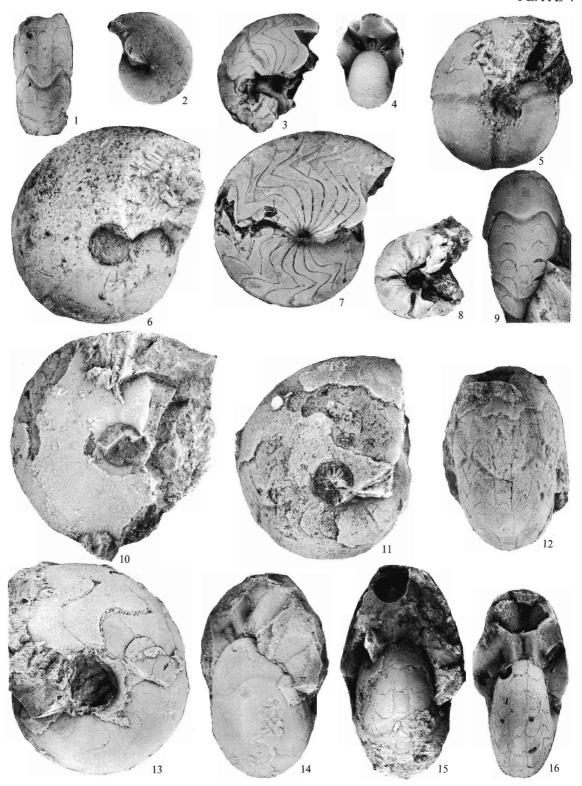
Range. Late Chadian to Pendleian.

Remarks. The following species are included:

- B. fournieri (Delépine, 1940a); black limestones, near Dinant, Belgium, late Asbian.
- B. glicki (Gordon, 1964); Fayetteville Shale (upper part), Elkins, Washington County, Arkansas, USA, Pendleian.
- B. implicatus (Phillips, 1836); Lower Bowland Shale Formation, Black Hall, Chipping, northwest England, late Asbian (B2b).
- B. inflatus sp. nov.; Clitheroe Limestone Formation, Clitheroe, north-west England, late Chadian.
- B. invaginatus Bisat, 1924; Upper Bowland Shale Formation, Malham, north-west England, late Pendleian (E1c).
- B. involutus (Pareyn, 1961); Djenien Series, Djebel Arlal (locality 22), Algeria, late Brigantian or Pendleian.
- B. komiorum Kusina, 1980; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.

EXPLANATION OF PLATE 4

- Figs 1, 3, 7, 16. *Dzhaprakoceras catena* sp. nov.; locality 3b. 1, 3, paratype, Ro 916; ventral and lateral views of exfoliated, fragmentary, adolescent conch; × 2·25. 7, 16, holotype, Ro 5192; lateral and oral views of slightly distorted, exfoliated late adolescent conch; × 2·5.
- Figs 2, 4. *Dzhaprakoceras levis* sp. nov. (see also Pl. 3, figs 4, 6); paratype, Ro 913; locality 3b; lateral and oral views of exfoliated juvenile conch; × 6.
- Fig. 5. Beyrichoceratoides sp. 1; Ro 5161; locality 8; lateral view of partly testiferous, exfoliated adolescent conch; × 2·5.
- Figs 6, 15. Dzhaprakoceras flexum sp. nov. (see also Pl. 5, figs 3, 5); paratype, LZA 9521; locality 3a; partly testiferous, adolescent conch: ×2.
- Figs 8–9. *Dzhaprakoceras undulatum* sp. nov. (see also Pl. 5, fig. 1; Pl. 7, fig. 2). 8, paratype, Ro 928; locality 3b; lateral view of fragmentary, exfoliated juvenile conch; × 2. 9, paratype, Ro 923; locality 4a; ventral view of exfoliated adolescent conch; × 2·25.
- Fig. 10. Dzhaprakoceras deflexum Kusina, 1980 (see also Pl. 6, figs 7–8; Pl. 7, fig. 15); LZA 9522; locality 4a; lateral view of partly testiferous late adolescent conch; × 2.
- Figs 11–12. *Dzhaprakoceras subglobosum* (Librovitch, 1927); Ro 5178; locality 2, lateral and ventral views of partly testiferous adolescent conch; × 3.
- Figs 13–14. Beyrichoceratoides inflatus sp. nov. (see also Pl. 7, figs 10, 12–14); holotype, Ro 5170; locality 4a; lateral and oral views of largely exfoliated late adolescent conch; ×2.



RILEY, Dzhaprakoceras, Beyrichoceratoides

- B. neilsoni (Currie, 1954); Index Limestone, Bishopbriggs, Scotland, Pendleian.
- B. redesdalensis (Hind, 1918); Redesdale Ironstone, Ridsdale, Northumberland, north-east England, Asbian.
- B. stenolobus (Phillips, 1836); Lower Bowland Shale Formation, Black Hall, Chipping, northwest England, late Asbian (B2b).
- B. truemani (Currie, 1954); Index Limestone Shales, Bishopbriggs, Scotland, Pendleian.
- B. truncatus (Phillips, 1836); Bolland, north-west England, early Brigantian.

Beyrichoceratoides can develop similar ventral lobe shapes to those of Dzhaprakoceras and Bollandoceras; the latter has a serpenticonic juvenile stage and is therefore easily distinguished from Beyrichoceratoides. Dzhaprakoceras, like Beyrichoceratoides, lacks a juvenile serpenticonic stage, but develops a bisinuous transverse ornament much earlier in ontogeny than Beyrichoceratoides.

Beyrichoceratoides inflatus sp. nov.

Plate 4, figures 13-14; Plate 7, figures 10, 12-14; Text-figures 14-15

Derivation of name. From the inflated adult whorls.

Material. Holotype, Ro 5170, paratypes Ro 5164, 5168 (late adolescent and mature conches, partly exfoliated and septate), preserved in packstone; from locality 4a.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

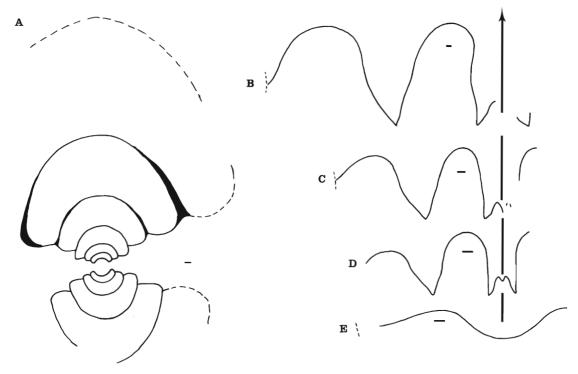
Diagnosis. A large Beyrichocheratoides with cadiconic early whorls which become inflated and subsphaeroconic at maturity.

Description. The conch is large and subsphaeroconic with inflated whorls when mature. The earlier whorls are cadiconic. The umbilicus is moderately wide with convex walls. The conch flanks are weakly convex and merge gently with the broad convex venter. The external ornament consists of delicate, bisinuous, transverse striae with a weak lateral sinus and a broad shallow hyponomic sinus. The internal ornament consists of feeble undulations which are parallel to the external striae. About four constrictions per whorl are present until late adolescence ($D=34\,\mathrm{mm}$); they are parallel with the external ornament. The ventral lobe is narrow with sub-parallel sides that become increasingly divergent orad. The lateral saddle is broad and rounded. The lateral lobe is narrow and acute with flexed sides.

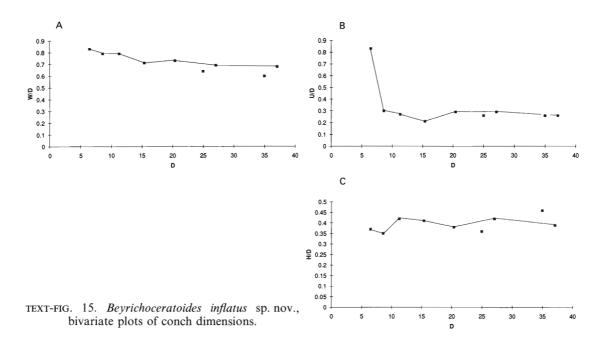
Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5164		28.5	18.1	44.4	_		
Ro 5170	35.0	16.0	9.0	21.0	0.60	0.26	0.46
Ro 5168	25.0	9.0	8.0	16.0	0.64	0.32	0.36
Ro 5164	37.1	14.6	9.7	25.4	0.68	0.26	0.39
Ro 5164	27.1	11.4	7.8	18.8	0.69	0.29	0.42
Ro 5164	20.4	7.8	5.9	14.8	0.73	0.29	0.38
Ro 5164	15.4	6.3	3.2	11.0	0.71	0.21	0.41
Ro 5164	11.3	4.8	3.0	8.9	0.79	0.27	0.42
Ro 5164	8.6	3.0	2.5	6.8	0.79	0.30	0.35
Ro 5164	6.5	2.4	2.1	5.4	0.83	0.32	0.37

Remarks. This species resembles superficially Eurites Kusina in conch form; however, Eurites has retrograde transverse ornament which lacks a lateral sinus.



TEXT-FIG. 14. Beyrichoceratoides inflatus sp. nov. A-B paratype, Ro 5164; A, axial section. B, ventral suture, H = 32·8; C, holotype, Ro 5170; ventral suture, H = 14·6; D-E, paratype Ro 5168; D, ventral suture, H = 8·9; E, constriction, H = 9·8.



Beyrichoceratoides sp. aff. komiorum Kusina, 1980.

Plate 5, figures 2, 6

1980 Beyrichoceratoides komiorum Kusina, p. 62, pl. 5, fig. 3.

Holotype. Figured by Kusina (1980, p. 62, pl. 5, fig. 3); specimen PIN 2775/347.

Material. Ro 5128, exfoliated, non-septate, mid-adolescent preserved in sparry calcite; from locality 2.

Range. Late Tournaisian, middle FA Zone; Clitheroe Limestone Formation (Bellman Limestone Member). The holotype comes from the base of the BB Zone, near the top of the Kosvinsky Horizon, north Urals, Russian Federation.

Description. The conch is small and subdiscoconic with a well rounded venter and a small umbilicus. The external and internal ornament are not preserved. Weak, radial constrictions are present. The suture is not visible.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Holotype (Kusina)	44.1	20.4	6.4	21.4	0.48	0.15	0.51
Ro 5182	17.0	8.0	2.0	10.0	0.59	0.12	0.47

Remarks. The radial constrictions exclude this ammonoid from being assigned to Dzhaprakoceras.

Beyrichoceratoides sp. 1

Plate 4, figure 5; Text-figure 16

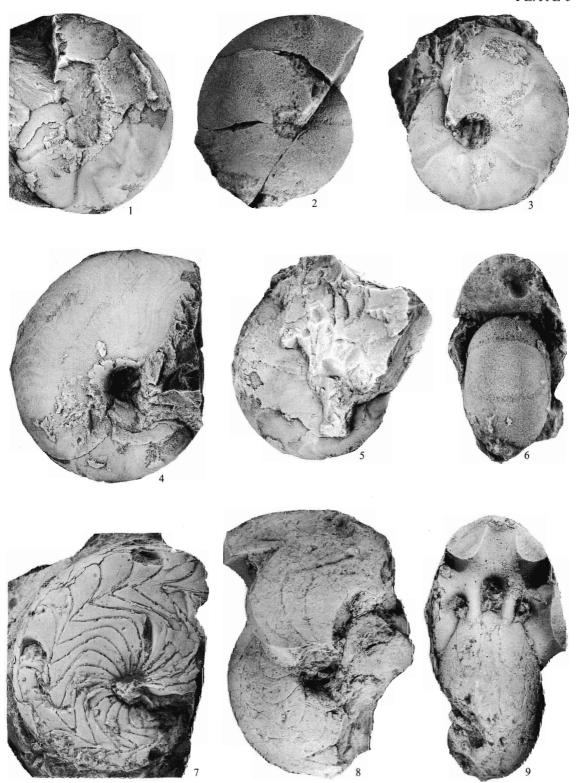
Material. Ro 5160-5162, fragmentary, mid-adolescent conches preserved in sparry calcite; from locality 8.

Range. Holkerian, upper BB Zone; Hodder Mudstone Formation.

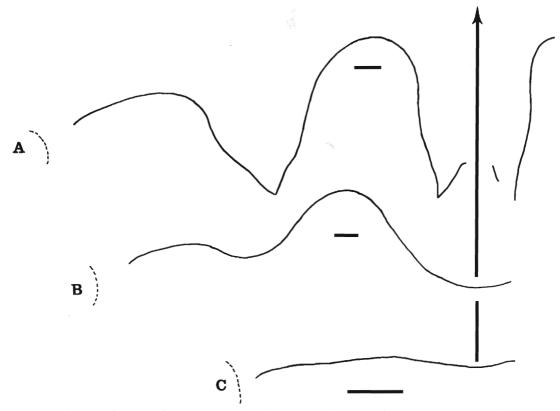
Description. The conch is subdiscoconic and relatively evolute with a well rounded venter. The external ornament is transverse and delicate, radial in juvenile whorls but, in later whorls, slightly bisinuous with a lateral and hyponomic sinus. An internal ornament is not present. There are four constrictions per whorl, which run parallel to the ornament. The ventral lobe is divergent-sided in juveniles but becomes increasingly lyrate with age. The lateral saddle is rounded. The lateral lobe is narrow and acute.

EXPLANATION OF PLATE 5

- Figs 1, 4. *Dzhaprakoceras undulatum* sp. nov. (see also Pl. 4, figs 8–9; Pl. 7 fig. 2); holotype, Ro 925; locality 4a; lateral views of fragmentary, partly testiferous late adolescent conch; ×2.
- Figs 2, 6. Beyrichoceratoides sp. aff. komiorum Kusina, 1980; Ro 5182; locality 2; lateral and ventral views of exfoliated adolescent conch; ×3.
- Figs 3, 5. *Dzhaprakoceras flexum* sp. nov. (see also Pl. 4, figs 6, 15); holotype, Ro 1009; locality 4a; lateral views of partly exfoliated late adolescent conch; × 2.
- Fig. 7. Dzhaprakoceras hispanicum (Foord and Crick, 1897); Ro 3720; locality 5; lateral view of exfoliated adult conch: ×1.5.
- Figs 8–9. *Dzhaprakoceras bellmanense* sp. nov. (see also Pl. 7, figs 12, 14); holotype, Ro 915; locality 4a; lateral and oral views of fragmentary, exfoliated adult conch; ×1.



RILEY, Dzhaprakoceras, Beyrichoceratoides



TEXT-FIG. 16. Beyrichoceratoides sp. A, C, Ro 5161; A, ventral suture, H = 13.5. C, constriction, H = 2.9. B, Ro 5160; constriction, H = 12.0;

EXPLANATION OF PLATE 6

Figs 1, 4, 12. Parahammatocyclus chaigleyensis gen. et sp. nov. (see also Pl. 3, figs 1–3, 8; Pl. 7, fig. 8). 1, paratype, Ro 5137; locality 6; lateral view of mainly exfoliated adolescent conch; note failure of tubercles toward aperture; ×4. 4, paratype, Ro 5105; locality 5; lateral view of fragmentary, partly testiferous adolescent conch; note failure of tubercles in final whorl; ×4. 12, paratype, Ro 3721; locality 5; lateral view of fragmentary, testiferous adolescent conch; ×3.5.

Figs 2–3. *Dzhaprakoceras westheadi* sp. nov. (see also Pl. 7, figs 4–6); holotype, LZA 9407; locality 3b; lateral and oral views of fragmentary, exfoliated adolescent conch; × 2.

Figs 5–6. Bollandoceras hodderense (Bisat, 1924) (see also Pl. 8, figs 2–3, 8–9, 11); lectotype (here designated), BMNH C25732; locality 12; oral and lateral views of fragmentary, exfoliated late adolescent specimen; ×2.

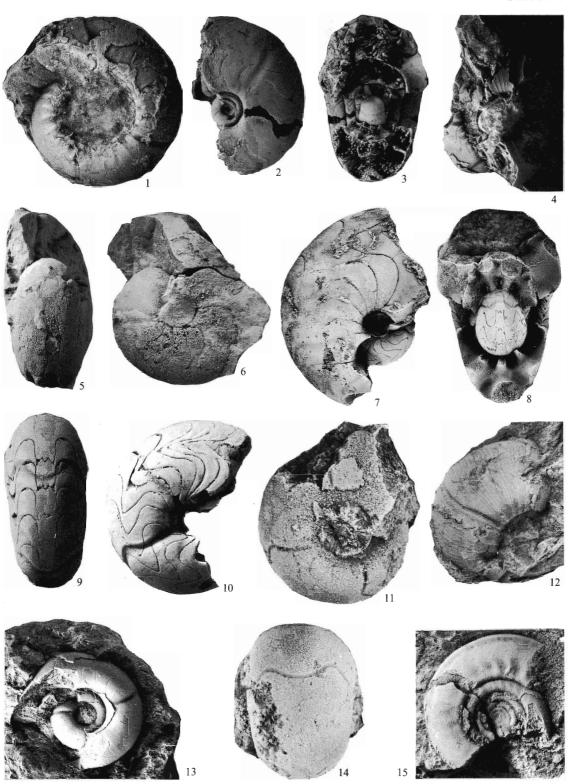
Figs 7–8. Dzhaprakoceras deflexum Kusina, 1980 (see also Pl. 4, fig. 10); Ro 1007; locality 4a; lateral and oral views of fragmentary, exfoliated late adolescent conch; × 2.

Figs 9–10. *Dzhaprakoceras paracatena* sp. nov.; holotype, LZA 9391; locality 3b; ventral and lateral views of fragmentary, exfoliated adolescent conch; ×4.

Figs 11, 14. *Dzhaprakoceras djaprakense* (Librovitch, 1927); Ro 5193; locality 1; lateral and ventral views of exfoliated adolescent conch; × 3.

Fig. 13. Beyrichoceratoides sp. 2 (see also Pl. 18, figs 13, 15); Ro 5076b; locality 6; lateral view of fragmentary, partly testiferous juvenile conch; ×4.

Fig. 15. Nomismoceras rotiforme (Phillips, 1836) (see also Pl. 2, fig. 2); holotype, BMNH C271a; precise locality unknown, 'Bolland'; lateral view of fragmentary, exfoliated conch; ×4.



RILEY, mid-Dinantian ammonoids

\mathbf{r}		
D	ım	ensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5161 Ro 5162	18.5	13·0 9·0	<u> </u>	12·0* 9·5	— 0·51	0.32	0.49

Remarks. This form probably represents a new species; further description awaits the discovery of more material.

Beyrichoceratoides sp. 2

Plate 6, figure 13; Plate 8, figures 13, 15

Material. Specimen Ro 5076b, juvenile, partly testiferous conch, embedded in packstone; Ro 5056, exfoliated slightly crushed testiferous adolescent conch; RH 1638, fragmentary, testiferous, slightly crushed adult conch. All the material is from locality 6.

Range. Arundian, lower BB Zone; Hodder Mudstone Formation (Chaigley Limestone Member).

Description. The adult conch is subdiscoconic, but juveniles are cadiconic. The external ornament consists of straight, transverse striae in juveniles, which become progressively bisinuous during growth. In the adult the striae are prosiradiate from the umbilical edge, which is raised slightly, but then deflect into a gentle lateral sinus on the mid-flank. There is a broad lingua and, apparently, a well developed hyponomic sinus, which is incompletely preserved. The venter is still rounded. The internal ornament shows feeble undulations which follow the path of the external ornament. Constrictions are present, at least until mid-adolescence; there are three per whorl and they also follow the external ornament. The suture is incompletely preserved but shows a rounded lateral saddle and an acute lateral lobe.

Remarks. These specimens are placed in *Beyrichoceratoides* on account of the cadiconic juvenile stage and the late development of strongly bisinuous ornament. *Beyrichoceratoides* sp. 1 is more laterally compressed with a less steep umbilical wall.

Genus Bollandoceras Bisat, 1952

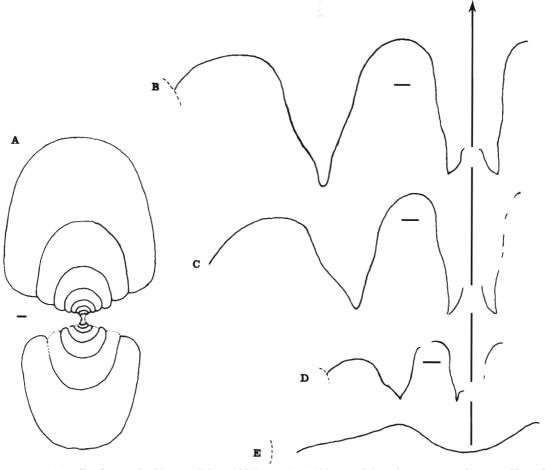
Type species. Beyrichoceras submicronotum Bisat, 1934.

Diagnosis. The conch is of small to moderate size, subsphaeroconic to discoconic with a small umbilicus. Juveniles are serpenticonic, but by early adolescence the conch is cadiconic. The venter may be rounded or subquadrate. The external ornament is transverse, radial in juvenile specimens but later bisinuous with a lateral and hyponomic sinus. Constrictions, if present, are parallel to the ornament. The ventral lobe has flexed, divergent (orad) sides. The lateral saddle is rounded. The lateral lobe is acute, often with flexed sides.

Range. Arundian to Asbian (possibly early Brigantian).

Remarks. The following species are included:

- B. bashachense (Popov, 1965); Kara-Kiya Suite, Tien-Shan, Kirgizya, CIS, Arundian.
- B. excavatum (Phillip, 1836); Carboniferous Limestone, Bowes, north-west England, Asbian.
- B. furiense (Padget, 1952); Fury River, Co. Tyrone, Northern Ireland, Asbian.
- B. globosum Bisat 1952; Lower Bowland Shale Formation, Barnoldswick, north-west England, late Asbian.
- B. hodderense (Bisat, 1924); Great Falls Limestone Formation, Stonyhurst, north-west England, late Holkerian.



TEXT-FIG. 17. Bollandoceras hodderense (Bisat, 1924) A-B, Ro 5190; A, axial section; B, ventral suture, H = 12·2, W = 13·6; C, E, Ro 5191; C, ventral suture, D = 24·4, H = 11·2, U = 4·9. W = 13·0; E, constriction, H = 8·2, W = 10·1 (same scale as C). D, lectotype, BMNH c25732; ventral suture, H = 6·8.

- B. librovitchi (Popov, 1965); base of Kara-Kiya Suite, Tien-Shan, Kirgizya, CIS, Arundian.
- B. mackellari (Campbell, Brown and Coleman, 1983); Caswell Creek Group, Queensland, Australia, Asbian or Brigantian.
- B. micronotoides (Bisat, 1934); 'reef limestone', Swinden Quarry, Cracoe, northern England, Asbian.
- B. micronotum (Phillips, 1836); Bolland, north-west England, Asbian (preserved in reef limestone).
- B. minusculum (Kusina, 1980); Kosvinsky Suite, Kozhim River, north Urals, Russian Federation, CIS, ?Arundian.
- B. stylense Kusina, 1991; zone C_1V_e base of Stylsky Horizon, Donets Basin, Ukraine, CIS, ?Arundian.
- B. submicronotum (Bisat, 1934); 'reef limestone', Cow Low Nick, Castleton, northern England, Asbian.

Dzhaprakoceras? minisculum Kusina, 1980, belongs here on account of its serpenticonic early whorls and divergent-sided (orad), flexed ventral lobe. Beyrichoceras (Bollandites) bashachensis

Popov, 1965 should also be placed in *Bollandoceras* since its constrictions are bisinuous, a feature not present in *Bollandites* as redefined herein (p. 75).

Bollandoceras hodderense (Bisat, 1924)

Plate 6, figures 5-6; Plate 8, figures 2-3, 8-9, 11; Text-figures 17-18

1924	Beyrichoceras hodderense Bisat, p. 84, pl. 9, figs 6-7
non 1940a	Beyrichoceras hodderense Bisat var. vallense nov. var. Delépine, p. 63, pl. 3, figs 13-18.
1961	Bollandoceras hodderense (Bisat); Ramsbottom in Earp et al., pl. 11, fig. 9a-b.
non 1961	Beyrichoceras hodderense Bisat; Pareyn p. 114, pl. 9, figs 8–16; text-fig. 14.

Type specimen. Lectotype here designated, BMNH C. 25732, no original figure.

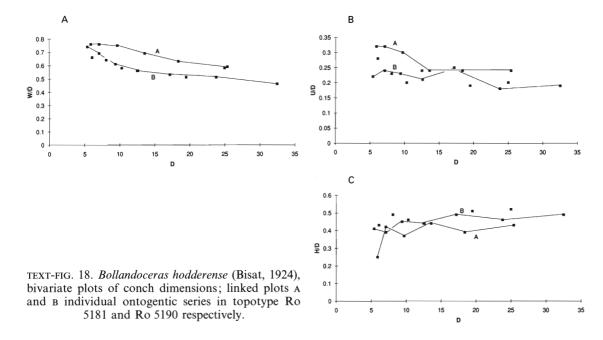
Material. Ro 5188–91, RH 1536–1566, adolescent and mature conches, exfoliated and preserved in micrite; from localities 9 (Ro 5189–91), 11 (RH 1536–1566) and 12 (topotype, Ro 5188).

Range. Holkerian, upper BB Zone; Hodderense Limestone Formation. Known also from the Milldale Limestones in Staffordshire. Records from Belgium and the Sahara are incorrect.

Diagnosis. A *Bollandoceras* with a stout subdiscoconic conch which maintains constrictions and an expanding umbilicus throughout ontogeny.

Description. The conch is variable from discoconic to subpachyconic, involute, with moderate whorl expansion. The umbilicus expands throughout growth. The external ornament consists of weak, gently bisinuous transverse striae, three to four per mm at $H=7\,\mathrm{mm}$. The internal ornament shows feeble undulations which follow the external ornament. There are up to six constrictions per whorl, these run parallel to the external ornament and are retained throughout growth. The suture consists of a deep, narrow ventral lobe with flexed sides which diverge orad, a rounded lateral saddle and a relatively narrow lateral lobe with an acuminate termination.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5190	32.5	16.0	6.0	15.0	0.46	0.19	0.49
Ro 5191	25.0	13.0	5.0	14.5	0.58	0.20	0.52
Ro 1536	19.5	9.9	3.7	10.0	0.51	0.19	0.51
Ro 5188	12.5	5.5	3.0	7.0	0.56	0.24	0.44
Ro 1550	10.3	4.7	2.1	6.0	0.58	0.20	0.46
Ro 1555	8.1	4.0	1.9	5.2	0.64	0.23	0.49
Ro 1564	6.1	2.6	1.7	4.0	0.66	0.28	0.43
Ro 5181	25.4	11.0	6.0	14.9	0.59	0.24	0.43
Ro 5181	18.4	7·1	4.4	11.5	0.63	0.24	0.39
Ro 5181	13.6	6.0	3.3	9.4	0.69	0.24	0.44
Ro 5181	9.7	3.6	2.9	7.3	0.75	0.30	0.37
Ro 5181	7.1	3.0	2.3	5.4	0.76	0.32	0.42
Ro 5181	5.9	1.5	1.9	4.5	0.76	0.32	0.25
Ro 5190	23.8	10.9	4.3	12.2	0.51	0.18	0.46
Ro 5190	17.2	8.4	3.6	9.2	0.53	0.25	0.49
Ro 5190	12.6	5.6	2.7	7.1	0.56	0.21	0.44
Ro 5190	9.4	4.2	2.2	5.7	0.61	0.23	0.45
Ro 5190	7.1	2.8	1.7	4.9	0.69	0.24	0.39
Ro 5190	5.4	2.2	1.2	4.0	0.74	0.22	0.41



Remarks. Bol. hodderense is distinguishable from other subdiscoconic to discoconic Bollandoceras since the umbilicus does not contract and constrictions are retained during ontogeny.

Bollandoceras micronotum Phillips (1836)

Plate 8, figures 5, 7; Text-figures 19-20

1836	Goniatites micronotus Phillips, p. 234, pl. 19, figs 22–23.
1889	Glyphioceras micronotum (Phillips); Holzapfel, pl. 2, fig. 1.
1897	Glyphioceras micronotum (Phillips); Foord and Crick, p. 173, text-fig. 81.
1924	Beyrichoceras micronotum (Phillips); Bisat, p. 46, pl. 9, figs 4-5.
1934	Beyrichoceras micronotum (Phillips); Bisat, pp. 289, 291, pl. 24, fig. 2, text-fig. 18.
1940 <i>a</i>	Beyrichoceras micronotum (Phillips); Delépine, p. 68, pl. 4, figs 9-12.
1941	Beyrichoceras micronotum (Phillips); Delépine, p. 62, pl. 3, figs 25–32.

Holotype. BMNH C262a, figured by Phillips (1836, pl. 19, figs 22-23).

Material. Specimens Ro 4316, 4320, 4322–3, 4328–31, 4333, 4336, 4346, 4350, 4351–3, 4357 and 4362–80, solid partly exfoliated conches, some septate and fragmentary, representing all growth stages, preserved in micrite and silicified packstone; from localities 11 (uppermost beds) and 13.

Range. Asbian, B Zone; Pendleside Limestone to Lower Bowland Shale Formation. Widespread in this zone in western Europe.

Diagnosis. A laterally compressed, discoconic Bollandoceras in which the umbilicus contracts during late ontogeny.

Description. The conch is discoconic with moderate whorl expansion. The umbilicus expands until late ontogeny when it reduces in diameter. The external ornament consists of bisinuous transverse striae. The internal ornament shows feeble undulations which correspond with the external ornament. There are four constrictions per whorl; the constrictions also follow the path of the external ornament, but become weak during ontogeny.

\mathbf{r}					
"	im	PI	1.81	\boldsymbol{n}	ns

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 4322	19.0	10.6	3.1	10.4	0.55	0.16	0.56
Ro 4353	12.4	5.7		6.7	0.54	_	0.50
Ro 4330	9.8	5.4	1.2	5.4	0.55	0.12	0.55
Ro 4330	6.8	3.2	1.3	4.3	0.63	0.19	0.47
Ro 4330	4.8	2.2	1.2	3.4	0.71	0.25	0.46
Ro 4330	3.5	1.4	1.2	2.6	0.74	0.34	0.40
Holotype	24.5	12.3	4.0	12.0	0.49	0.16	0.50

Remarks. Bol. micronotum is close to Bol. minusculum (Kusina, 1980) which, however, lacks constrictions. Bol. micronotum differs from other discoconic Bollandoceras with laterally compressed conches and progressively contracting umbilici, in that the venter is not sub-quadrate at maturity.

Bollandoceras sp. 1

Plate 8, figure 4; Text-figure 21

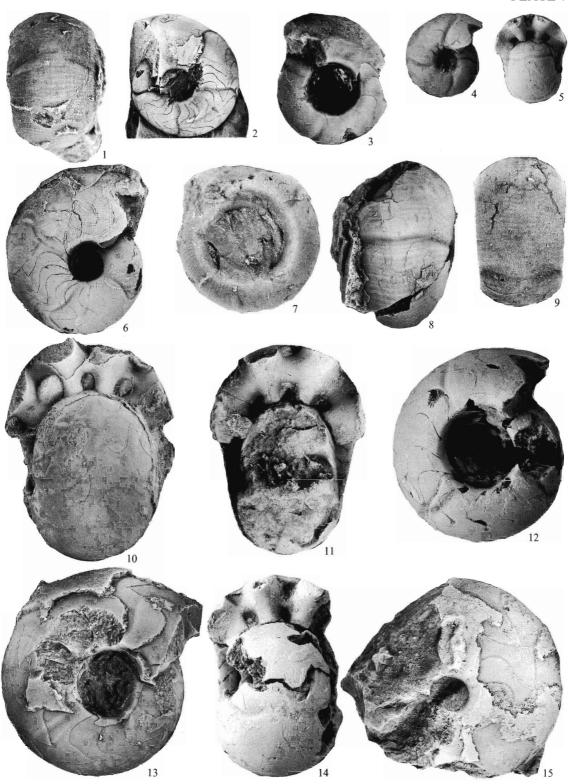
Material. One, exfoliated, septate early adolescent specimen, Ro 5163, preserved in sparry calcite; from locality 8.

Range. Holkerian, upper BB Zone; Hodder Mudstone Formation.

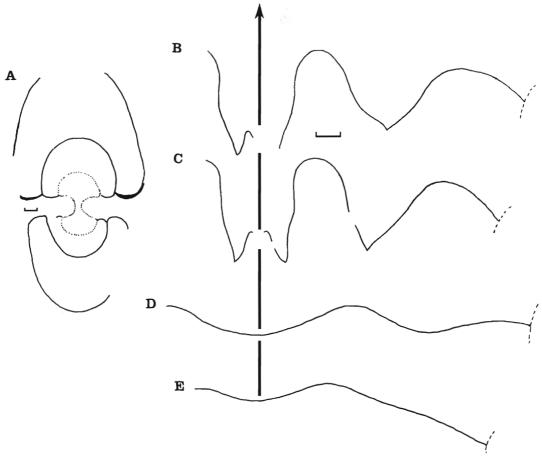
Description. Conch compressed, involute, but with an open umbilicus and rounded venter. The ornament, which is only partly preserved, consists of frail transverse bisinuous striae, with a lateral sinus and a strong hyponomic sinus, separated by a feeble lingua. Three strong constrictions occur, and these apparently follow the external ornament; the internal ornament is apparently smooth. The suture, although not well preserved, shows a flexed, divergent sided (orad) ventral lobe with flexed walls and a rounded lateral saddle.

EXPLANATION OF PLATE 7

- Figs 1, 3, 7, 9, 11. *Bollandites varians* sp. nov. (see also Pl. 8, fig. 14). 1, 7; paratype, Ro 951; locality 11; ventral and lateral views of testiferous juvenile; × 8. 3, paratype, Ro 5187; locality 11; lateral view of exfoliated adolescent conch; × 2. 9, paratype, Ro 953; locality 11; ventral view of exfoliated juvenile conch; note ribbing; × 8. 11, holotype, LZA 7865; locality 12; oral view of fragmentary, exfoliated adult conch; × 1·75. Fig. 2. *Dzhaprakoceras undulatum* sp. nov. (see also Pl. 4, figs 8–9; Pl. 5, fig. 1); paratype, Ro 923; locality 4a;
- lateral view of partly testiferous, adolescent conch; × 2·25.
- Figs 4–6. *Dzhaprakoceras westheadi* sp. nov. (see also Pl. 6, figs 2–3); locality 3b. 4–5, paratype, LZA 9407; lateral and oral views of exfoliated juvenile conch (dissected); ×4. 6, paratype, Ro 920; lateral view of exfoliated, partly crushed adolescent conch; ×2·5.
- Fig. 8. Parahammatocyclus chaigleyensis gen. et sp. nov. (see also Pl. 3, figs 1–3, 8; Pl. 6, figs 1, 4, 12); paratype, Ro 3723; locality 7; ventral view of fragmentary exfoliated adolescent conch; × 2.75.
- Figs 10, 13. Beyrichoceratoides inflatus sp. nov. (see also Pl. 4, figs 13–14); paratype, Ro 5168; locality 4a; oral and lateral views of largely exfoliated, late adolescent conch; × 2.
- Figs 12, 14. *Dzhaprakoceras bellmanense* sp. nov. (see also Pl. 5, figs 8–9); paratype, LZA 9419; locality 3b; fragmentary, exfoliated adolescent conch; × 4.
- Fig. 15. Dzhaprakoceras deflexum Kusina, 1980 (see also Pl. 4, fig. 10; Pl. 6, figs 7–8); LZA 9522; locality 4a; lateral view of fragmentary, partly testiferous adolescent conch; × 2.



RILEY, mid-Dinantian ammonoids



TEXT-FIG. 19. Bollandoceras micronotum (Phillips, 1836). A, Ro 4330; axial section. B, D, holotype, BMNH C262a; B, ventral suture, H = 9.7; constriction at D = 19.7, H = 9.4; C, E, Ro 4322; C, ventral suture, H = 8.4; E, constriction at H = 8.4.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5163	12.0	5.5	2.5	6.5	0.54	0.21	0.46

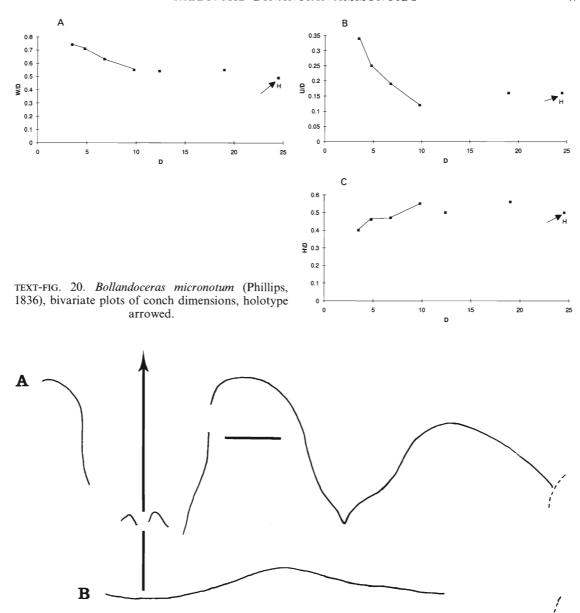
Remarks. Because the ontogenetic development is taxonomically important in the recognition of Bollandoceras species, this specimen has been left in open nomenclature pending discovery of further material.

Bollandoceras sp. 2

Plate 8, figures 6, 10; Text-figure 22

Material. Ro 5136, solid, mostly exfoliated, partly septate adolescent conch; from locality 6.

Range. Arundian, lower BB Zone; Hodder Mudstone Formation (Chaigley Limestone Member).



TEXT-FIG. 21. Bollandoceras sp. 1; A, Ro 5163; ventral suture, D = 11.0, H = 5.5; B, constriction, D = 11.0, H = 5.5.

Description. The conch is small, subdiscoconic, laterally compressed and relatively involute with rapid whorl expansion. The internal ornament consists of feeble transverse striae (five per mm on the ventro-lateral shoulder at H = 3.4), which are radial over the flanks, but curve aborad over the venter to form a broad hyponomic sinus. There are three constrictions per whorl, which have a slight lateral and a deeper hyponomic sinus (the external ornament is not preserved at comparable diameters). The ventral lobe is divergent-sided (orad) with slightly flexed sides. The lateral saddle is broad and rounded. The lateral lobe is shallow, broad and acute, with a flexed dorsal side.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5136	11.0	6.8	2·1	6.2	0.56	0.19	0.62

Remarks. The ornament style, broad nature of the lateral lobe and the divergent-sided (orad) ventral lobe are consistent with *Bollandoceras*. Specific identification awaits discovery of more material, so that the ontogeny can be elucidated. The adolescent conch of this species is more compressed than *Bol.* sp. 1, and the transverse ornament develops a lateral and hyponomic sinus much later in ontogeny.

Genus DZHAPRAKOCERAS Popov, 1965

Type species. Muensteroceras tianshanicum Librovitch, 1927.

Diagnosis (emended). The conch varies from mainly pachyconic to discoconic, but cadiconic to subsphaeroconic forms are also represented. A juvenile serpenticonic stage is absent. The external ornament consists of weak, bisinuous, transverse striae which are present from the late juvenile stage. Constrictions occur commonly and follow the external ornament, but may fail during adolescence. The internal ornament, if present, follows the external ornament and consists of feeble undulations, particularly in the lingual region; weak spiral undulations may also occur. The ventral lobe is lyrate. The lateral saddle is rounded. The lateral lobe is acute.

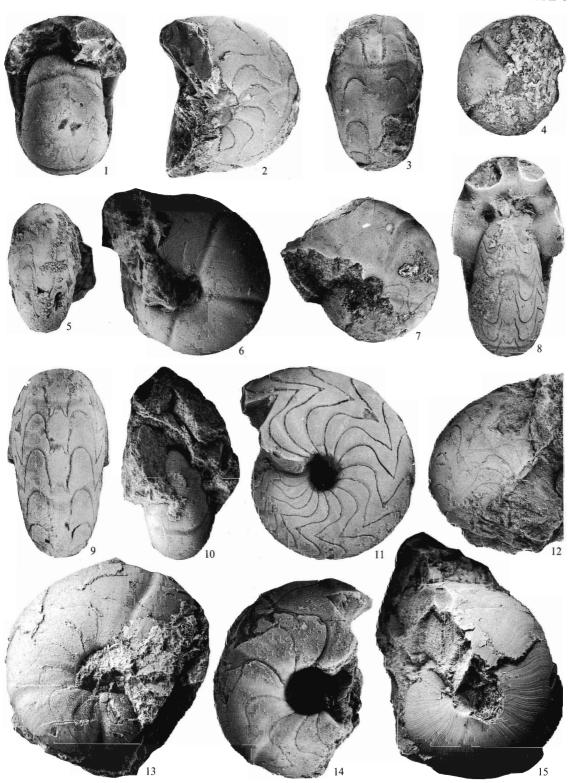
Range. (?Courceyan) Chadian to Arundian.

Remarks. The following species are included:

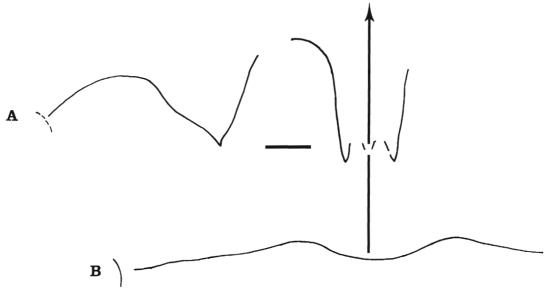
- D. bellmanense sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- D. catena sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.

EXPLANATION OF PLATE 8

- Figs 1, 14. *Bollandites varians* sp. nov. (see also Pl. 7, figs 1, 7, 9). 1, paratype, Ro 5187; locality 11; oral view of exfoliated adolescent conch; × 2. 14, holotype, LZA 7865; locality 12; lateral view of fragmentary, exfoliated adult conch; × 1.75.
- Figs 2–3, 8–9, 11. Bollandoceras hodderense (Bisat, 1924) (see also Pl. 6, figs 5–6). 2–3, topotype, Ro 5181; lateral and ventral views of fragmentary, exfoliated adult conch; locality 12; stout variant, ×1·6. 8–9, 11, Ro 5191; locality 13; oral, lateral and ventral views of exfoliated adult conch; ×2.
- Fig. 4. Bollandoceras sp. 1; Ro 5163; locality 8; lateral view of exfoliated, early adolescent conch; ×2.
- Figs 5, 7. Bollandoceras micronotum (Phillips, 1836); Ro 4322; locality 13; ventral and lateral views of exfoliated adolescent conch; × 2.
- Figs 6, 10. *Bollandoceras* sp. 2; Ro 5136; locality 6; lateral and oral views of exfoliated, early adolescent conch; × 4.
- Figs 13, 15. Beyrichoceratoides sp. 2 (see also Pl. 6, fig. 13); locality 6. 13, Ro 5056; lateral view of exfoliated, slightly crushed adolescent conch; ×4. 15, RH 1638; lateral view of fragmentary, testiferous, slightly crushed adult conch; ×2.
- Fig. 12. Beyrichoceras aff. obtusum (Phillips, 1836); Ro 2160; locality 13; lateral view of fragmentary exfoliated conch; ×2.



RILEY, mid-Dinantian ammonoids



TEXT-FIG. 22. Bollandoceras sp. 2; Ro 5136; A, ventral suture, D = 10·3, H = 5·2, U = 1·8, W = 6·1; B, constriction, same conch dimensions as A.

- D. chermnykhi Kusina, 1980; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.
- D. deflexum Kusina, 1980; ?Nortnichsky Horizon, Chernaya River, north Urals, Russian Federation, CIS, ?Arundian.
- D. djaprakense (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. duponti (Delépine, 1940a); Waulsortian, Drehance, Belgium, late Tournaisian.
 D. flexum sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late
- Chadian.

 D. gracile Kusina, 1980; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation,
- CIS, late Chadian.
- D. grande Kusina, 1980; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.
- D. hibernicum (Delépine, 1940b); Waulsortian, St Doughlas, Dublin, Ireland, late Tournaisian.
- D. hispanicum (Foord and Crick, 1897); Carboniferous, Asturias, Spain, Dinantian.
- D. improcerum Kusina, 1980; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, Chadian.
- D. kokdzharense (Popov, 1965); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. koninckianum (Schindewolf 1951a); Waulsortian, Pauquays, Belgium, late Tournaisian.
- D. latilobatum Kusina, 1983; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.
- D. latiumbilicatum (Kullmann, 1961); 'Marbre Griotte', Cantabria, north Spain, late Tournaisian or 'early Viséan.
- D. levis sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- D. narynense (Popov, 1965); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. paracatena sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- D. planum Kusina and Yatskov, 1987; Milin Suite, Novaya Zemlya, CIS, 'mid-Viséan'.

- D. pseudoparallelum (Delépine, 1941); 'Green Slates', Erfoud, Morocco, mid-Dinantian.
- D. quantulum Kusina and Yatskov, 1987; Milin Suite, Novava Zemlya, CIS, 'mid-Viséan'.
- D. shabyrense (Popov, 1968); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. sonkulica (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. subglobosum (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. subtile Kusina, 1980; Nortnichsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, Arundian.
- D. tianshanicum (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.
- D. westheadi sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- D. undulatum sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.

D. grande, D. latiumbilicatum and D. undulatum sp. nov. may belong to Beyrichoceratoides, since their early whorls are unknown. D.? minusculum Kusina, 1980 is transferred to Bollandoceras on account of its serpenticonic early whorls and divergent-sided (orad) ventral lobe. Dzhaprakoceras was originally intended to include muensteroceratids with lyrate ventral lobes; however, this ventral lobe form is not restricted to Dzhaprakoceras, occurring also in Beyrichoceratoides and Kozhimites. In this study Dzhaprakoceras is interpreted as including muensteroceratids with a combination of bisinuous ornament (present from the late juvenile stage), lyrate ventral lobes, and absence of a juvenile serpenticonic stage. It is unfortunate that the ontogeny of the type species of Dzhaprakoceras is unknown. Muensteroceras pseudaganides George and Ponsford, 1935 is too poorly known for definite inclusion in Dzhaprakoceras; the holotype is probably an Irinoceras, but the paratypes are Dzhaprakoceras sp.

Dzahprakoceras (Akeshakoceras) longilobatum Liang and Wang, 1991 does not belong to Dzhaprakoceras. The holotype (Liang and Wang 1991, pl. 14, figs 11–12) shows spiral ornament and an acute lateral saddle. It is clearly a goniatitid and not a muensteroceratid.

Dzhaprakoceras bellmanense sp. nov.

Plate 5, figures 8–9; Plate 7, figures 12, 14; Text-figures 23–24

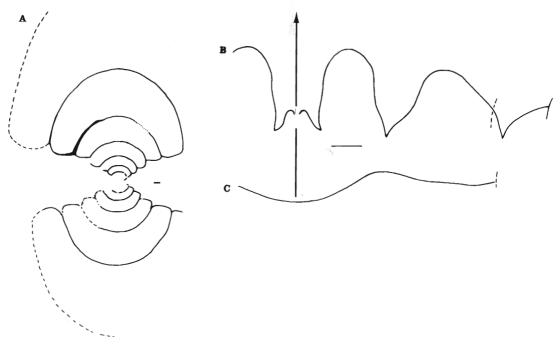
Deviation of name. From the type locality, Bellman Quarry.

Material. Holotype, Ro 915, fragmentary, exfoliated, late adolescent, septate conch preserved in micrite; from locality 4a; and paratype, LZA 9419, exfoliated, septate, pyritic conch; from locality 3b.

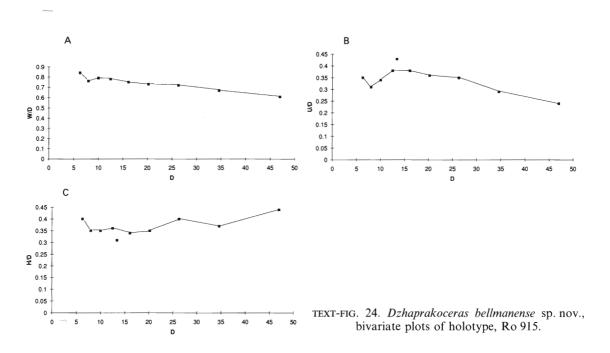
Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A large subdiscoconic Dzhaprakoceras which develops a lyrate ventral lobe during late adolescence.

Description. The conch is large (D=70+), subdiscoconic and has an umbilicus of moderate width. Whorl expansion is moderate, but slow below D=35, when the conch is cadoconic. The external ornament is unknown. The internal ornament consists of very weak transverse ridges with a shallow broad lateral and hyponomic sinus. Constrictions are present below D=13 and run parallel with the external ornament. There are four per whorl. The suture is only clearly seen in the adolescent specimen, where the ventral lobe is slightly divergent-sided (orad), narrow and deep. In the holotype it is lyrate, but poorly preserved. The lateral saddle is narrow and rounded. The lateral lobe is deep and narrow and has flexed, subparallel sides, with an acute termination.



TEXT-FIG. 23. Dzhaprakoceras bellmanense sp. nov. A, holotype, Ro 915; axial section. B-C, paratype, LZA 9419; B, ventral suture, D = 13·4, H = 4·2, U = 5·8; C, constriction, same dimensions as B.



\mathbf{r}						
"	im	PV	25	വ	ns	1

Specimen	D	H	U	W	W/D	U/D	H/D
Ro 915	47.0	20.7	11.4	28.6	0.61	0.24	0.44
Ro 915	34.6	12.8	9.9	23.2	0.67	0.29	0.37
Ro 915	26.3	10.4	9.2	18.9	0.72	0.35	0.40
Ro 915	20.2	7.0	7.2	14.7	0.73	0.36	0.35
Ro 915	16.1	5.5	6.1	12.0	0.75	0.38	0.34
Ro 915	12.5	4.5	4.8	9.7	0.78	0.38	0.36
Ro 915	10.0	3.5	3.4	7.9	0.79	0.34	0.35
Ro 915	8.0	2.8	2.5	6.1	0.76	0.31	0.35
Ro 915	6.3	2.5	2.2	5.3	0.84	0.35	0.40

Remarks. This species is closest to a form incorrectly attributed to D. hispanicum, by Kusina (1980, pl. 4, fig. 5), but differs in its more inflated whorls and larger umbilicus.

Dzhaprakoceras catena sp. nov.

Plate 4, figures 1, 3, 7, 16; Text-figures 25-26

Derivation of name. From the Latin catena (= chain), with reference to the chain-like appearance of the coalesced ventral lobes.

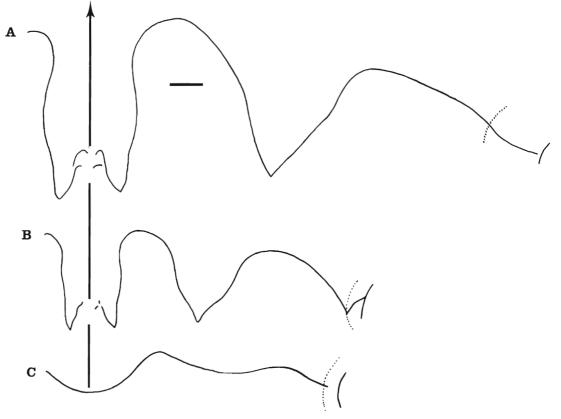
Material. Holotype, Ro 5192, exfoliated, septate, pyritic late adolescent; from locality 3b; paratypes, LZA 9396, 9389, exfoliated, pyritic septate juveniles; from locality 3b; LZA 9395, pyritic, septate early adolescent; from locality 4b; Ro 916, exfoliated, septate adolescent preserved in micrite; from locality 4a.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

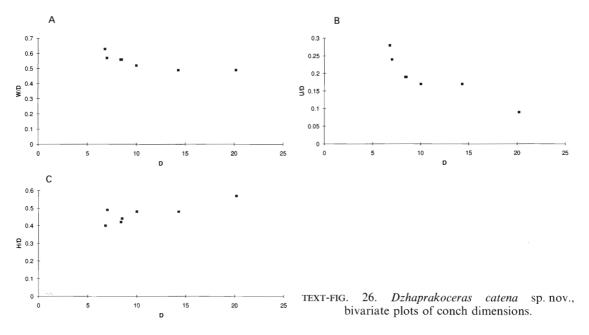
Diagnosis. A small involute, discoconic *Dzhaprakoceras* with coalescent ventral lobes present by early adolescence and a relatively broad lateral lobe. Constrictions occur in the juvenile whorls, but fail by mid-adolescence.

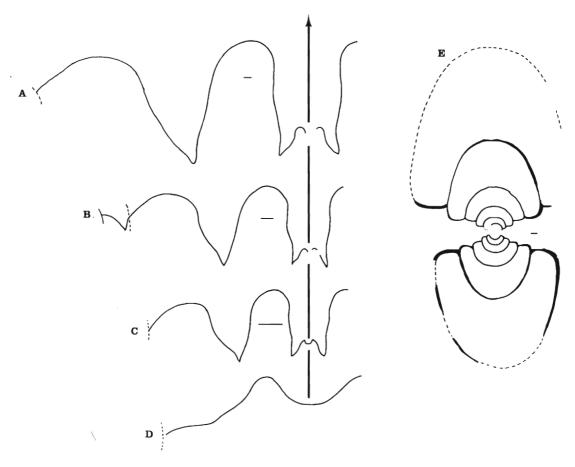
Description. The conch is small, discoconic, laterally compressed and involute with pachyconic early whorls, rapid whorl expansion and a small umbilicus. The external ornament is unknown. The internal ornament consists of weak transverse ridges on the flanks, corresponding with the position of the lateral sinus and lingua; these characters appear to be absent below H = 8. Constrictions are present below H = 5 (about four per whorl), they are bisinuous with a nearly V-shaped hyponomic sinus and a shallow, broad lateral sinus. The ventral lobe is lyrate and deep; it becomes coalescent with adjacent lobes by early adolescence. The lateral saddle is broad and rounded. The relatively broad lateral lobe is nearly as deep as the ventral lobe and is acute, with flexed sides.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5192	20.2	11.5	1.9	9.9	0.49	0.09	0.57
LZA 9405	14.3	6.8	2.5	7.0	0.49	0.17	0.48
LZA 9409	10.0	4.8	1.7	5.2	0.52	0.17	0.48
LZA 9398	8.5	3.7	1.6	4.8	0.56	0.19	0.44
LZA 9394	8.4	3.5	1.6	4.7	0.56	0.19	0.42
LZA 9384	7.0	3.4	1.7	4.0	0.57	0.24	0.49
Ro 929	6.8	2.7	1.9	4.3	0.63	0.28	0.40



TEXT-FIG. 25. Dzhaprakoceras catena sp. nov. A, holotype, Ro 5192; ventral suture, $D=18\cdot4$, $H=9\cdot7$, $U=2\cdot3$, $W=9\cdot1$. B-C, paratype, LZA 9405; B, ventral suture, $D=12\cdot9$, $H=6\cdot2$, $U=2\cdot8$, $W=6\cdot4$; C, constriction, $D=12\cdot6$, $H=6\cdot0$, $U=2\cdot5$, $W=5\cdot0$.





TEXT-FIG. 27. Dzhaprakoceras deflexum Kusina, 1980. A, Ro 1008; ventral suture, H = 26·0, W = 28·4. B, Ro 1007; ventral suture, H = 12·6, W = 13·2. C-D, LZA 9390; c, ventral suture, H = 5·5, W = 7·4; D, constriction, same dimensions as C. E, Ro 1025; axial section.

Remarks. D. gracile has a stouter conch than D. catena sp. nov. D. chermnykhi has a convex (dorsad) ventral side to the lateral lobe whereas in D. catena sp. nov. it is concave. D. latilobatum, unlike D. catena sp. nov., has divergent prongs (aborad) to the ventral lobe. The closest species to D. catena sp. nov. is D. paracatena sp. nov.; however, in the latter species the lateral lobe is narrower and has less divergent sides.

Dzhaprakoceras deflexum Kusina, 1980

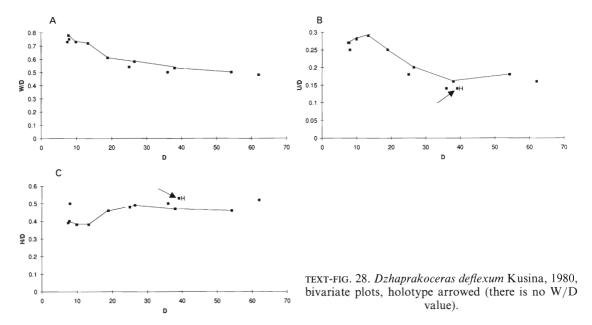
Plate 4, figure 10; Plate 6, figures 7–8; Plate 7, figure 15; Text-figures 27–28

1980 Dzhaprakoceras deflexum, Kusina, p. 58, pl. 4, fig. 7.

Holotype. PIN 2775/365, figured by Kusina (1980, pl. 4, fig. 7).

Material. Seven conches, LZA 9522, Ro 1007–8, and Ro 1023–5, partly exfoliated, some septate, representing all growth stages, preserved in micrite; from locality 4a; and Ro 1006, an early adolescent, pyritic internal mould; from locality 3b.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member). The type material is thought to be from the Arundian lower BB Zone, Nortnichsky Horizon, north Urals.

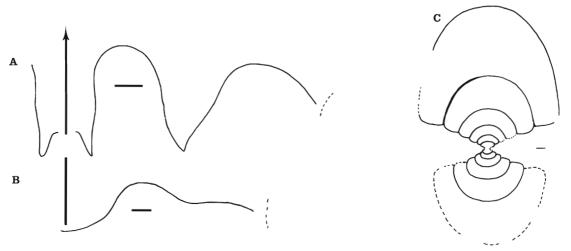


Diagnosis. A subdiscoconic *Dzhaprakoceras* in which the ventral lobe is slightly divergent-sided orad).

Description. The conch is subdiscoconic with a small umbilicus which expands constantly until maturity after which the umbilical diameter remains unchanged. The juvenile whorls are cadiconic; whorl expansion is relatively rapid after D = 13. The external ornament is only partly preserved (LZA 9522); it is virtually smooth, apart from very weak striae and growth lamellae. These are grouped together at the umbilical margin, giving the appearance of bands, but become virtually imperceptible over the mid-flank and venter. The internal ornament consists of weak bisinuous ridges, which in adolescence give rise to vesicular ridges over the ventrolateral region corresponding with linguae. Four constrictions, which are bisinuous, with a deep, broad hyponomic sinus, are present up to late adolescence, after which they fail. The suture shows a narrow deep ventral lobe with lyrate, parallel, flexed sides; the lateral saddle is narrow and rounded; the lateral lobe is deep, with divergent sides (orad) and a subacute termination.

\mathbf{r}				
D	im	en	\$1.0	ons.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 1008	62.0	32.0	10.0	30.0	0.48	0.16	0.52
Ro 1025	54.2	24.7	10.0	27.0	0.50	0.18	0.46
Holotype	39.1	20.8	5.3			0.14	0.53
Ro 1007	36.0	18.0	5.0	18.0	0.50	0.14	0.50
LZA 9522	25.0*	12.0	4.5	13.5	0.54	0.18	0.48
Ro 1006	8.0	4.0	2.0	6.0	0.75	0.25	0.50
LZA 9406	7.5	2.9	2.3	5.5	0.73	0.27	0.39
Ro 1025	38.0	17.8	6.4	20.2	0.53	0.16	0.47
Ro 1025	26.5	13.1	5.2	15.3	0.58	0.20	0.49
Ro 1025	18.9	8.7	4.7	11.6	0.61	0.25	0.46
Ro 1025	13.3	5.0	3.8	9.6	0.72	0.29	0.38
Ro 1025	9.9	3.8	2.8	7.2	0.73	0.28	0.38
Ro 1025	7.8	3.1	2.1	6.1	0.78	0.27	0.40



TEXT-FIG. 29. *Dzhaprakoceras flexum* sp. nov. A, C, paratype, LZA 9521; A, ventral suture, H = 7.9, W = 11.1; C, axial section. B, holotype, Ro 1009; constriction, H = 7.9 (reversed image).

Remarks. The slightly divergent-sided (orad) ventral lobe in adolescents superficially resembles that of Bollandoceras, but is not as severe. The lack of a juvenile serpenticonic stage, coupled with the early appearance of bisinuous transverse ornament, clearly places this species in Dzhaprakoceras. Kusina (1980) gave the type horizon as possibly the Nortnichsky Horizon; the presence of this species in the late Chadian in Britain suggests it is more likely that the type horizon is in the Kosvinsky Horizon.

Dzhaprakoceras sp. juv. cf. djaprakense (Librovitch, 1927)

\	Plate 6, figures 11, 14
1927	Muensteroceras djaprakense Librovotch, p. 37, pl. 7, figs 3-4; text-figs 18-19.
1961	Muensteroceras djaprakense Librovitch; Kullmann, p. 262, text-fig. 7b.
1961	Muensteroceras subglobosum Librovitch; Pareyn, p. 104, pl. 8, figs 27-28.
1968	Dzhaprakoceras djaprakense (Librovitch); Popov, p. 89, pl. 3, figs 6–7, pl. 14, fig. 6, text-figs 15g, i-m, 16z.
1980	Dzhaprakoceras djaprakense (Librovitch); Kusina, p. 53, pl. 4, fig. 1; text-fig. 13.

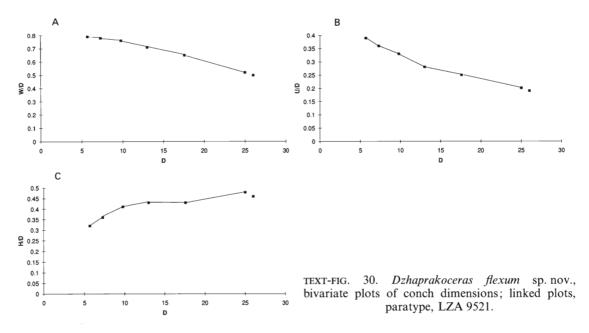
Lectotype. Designated by Kusina 1980, specimen 101/2473, TSNIGR Museum, St Petersburg, figured by Librovitch (1927, pl. 7, figs 3-4).

Material. Ro 5193, mid-adolescent, mostly exfoliated, septate (details of suture unclear), preserved in sparry calcite; from locality 1.

Range. Late Tournaisian, middle FA Zone; Clitheroe Limestone Formation (Coplow Limestone Member). Known also from middle FA Zone assemblages in north Africa, Spain and Tien-Shan (CIS).

Diagnosis. A subglobose Dzhaprakoceras with a relatively wide umbilicus.

Description. The conch is subpachyconic, with a relatively wide umbilicus. The external ornament is poorly preserved and consists of weak bisinuous striae. There are three narrow, deeply incised constrictions per whorl which show through to the exterior of the conch. Both the lateral and hyponomic sinuses are broad and separated by a narrow lingua. The suture is not well preserved, but consists of a rounded lateral saddle with flexed sides and an acute lateral lobe; the rest of the suture is not visible.



Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 5193	16.0	6.5	5.0	12.0	0.75	0.31	0.41

Remarks. The ventral lobe is not preserved; identification is based on the conch shape and the early presence of bisinuous ornament. The only other subglobose *Dzhaprakoceras*, *D. subglobosum*, has a smaller umbilicus and a more laterally compressed conch.

Dzhaprakoceras flexum sp. nov.

Plate 5, figures 3, 5; Text-figures 29-30

Derivation of name. From the Latin flexus (= sinuosity), with reference to the ornament.

Material. Holotype Ro 1009; from locality 4a; and paratype LZA 9521; from locality 3b. Both conches are late adolescents, partly exfoliated with intact body chambers, preserved in micrite.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A small pachyconic Dzhaprakoceras with a wide umbilicus and persistent constrictions.

Description. The conch is small, pachyconic, with cadiconic early whorls, a relatively wide umbilicus and moderate whorl expansion. The external ornament consists of weak, bisinuous, transverse striae with a broad lingua. The internal ornament consists of feeble transverse undulations, more accentuated in the lingual region, giving an impression of a subtle lingual groove. Five constrictions per whorl persist throughout growth and follow the ornament. The ventral lobe is narrow and lyrate. The lateral saddle is broad and rounded. The lateral lobe is acute and broad with flexed sides and is nearly as deep as the ventral lobe.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 1009	26.0	12.0	5.0	13.0	0.50	0.19	0.46
LZA 9521	25.0	12.0	5.0	13.0	0.52	0.20	0.48
LZA 9521	17.6	7.5	4.4	11.5	0.65	0.25	0.43
LZA 9521	13.0	5.6	3.6	9.2	0.71	0.28	0.43
LZA 9521	9.8	4.0	3.2	7.4	0.76	0.33	0.41
LZA 9521	7.3	2.6	2.6	5.7	0.78	0.36	0.36
LZA 9521	5.7	1.8	2.2	4.5	0.79	0.39	0.32

Remarks. This species is closest to *D. subglobosum* from which it differs in its constantly expanding umbilicus and more pachyconic form.

Dzhaprakoceras hispanicum (Foord and Crick, 1897)

Plate 5, figure 7; Text-figures 31–32

1897	Glyphioceras hispanicum, Foord and Crick, p. 190, text-fig. 91.
?1941	Muensteroceras occidentale Foord; Delépine, p. 58, pl. 1, figs 7–8; pl. 2, fig. 1.
1943	Beyrichoceras hispanicum (Foord and Crick); Delépine, p. 47, pl. 1, figs 1–2.
1951	Nautellipsites difficilis (Foord); Smyth, p. 298, pl. 11, figs 1–2; text-figs 1h–1.
1961	?Muensteroceras hispanicum (Foord and Crick); Kullmann, p. 258, pl. 20, fig. 1; text-figs 7a, 8b.
1960	Nautellipsites hispanicus (Foord and Crick); Wagner-Gentis, p. 49, text-figs 1-3, 4a-c.
non 1980	Dzhaprakoceras hispanicum (Foord and Crick); Kusina, p. 56, pl. 4, fig. 5; text-fig. 16.

Type specimen. Glyphioceras hispanicum, BMNH 37492, figured by Foord and Crick, 1897, p. 190, text-fig. 91.

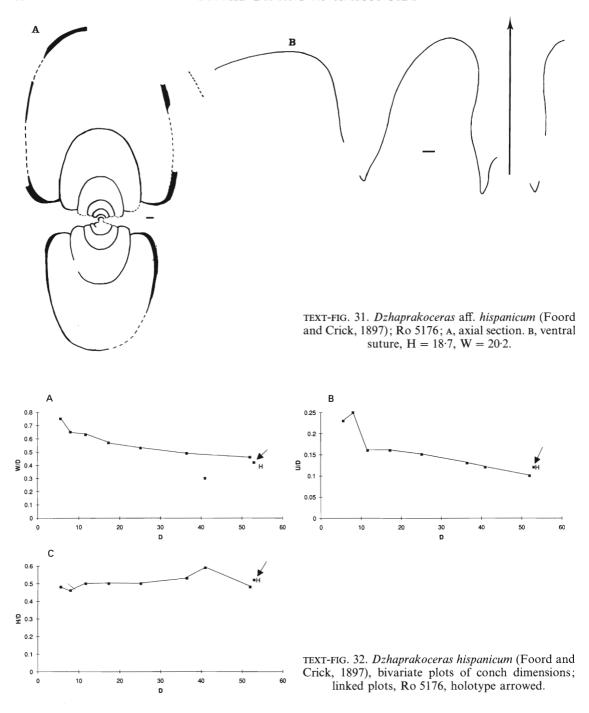
Material. Two septate, mature, mostly exfoliated conches, preserved in micrite; Ro 3720, from locality 5; and Ro 5176, from locality 2.

Range. Late Tournaisian to Arundian, middle FA Zone to lower BB Zone; Clitheroe Limestone Formation (Bellman Limestone Member) to Hodder Mudstone Formation (Chaigley Limestone Member).

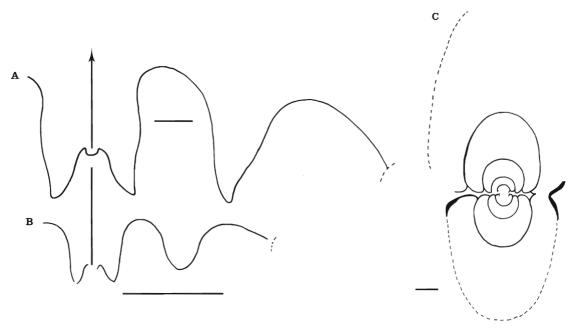
Diagnosis. A large subdiscoconic Dzhaprakoceras with rapid whorl expansion.

Description. The conch is subdiscoconic with rapid whorl expansion and a small umbilicus. The ornament is very weak and consists of simple, bisinuous, transverse striae with a well developed lingua. There are no constrictions and the internal surface of the conch is smooth. The ventral lobe is lyrate, deep and narrow. The lateral lobe is also deep and narrow with a slightly flexed ventrad wall.

D	Н	U	W	W/D	U/D	H/D
53.0	27.5	6.5	22:0	0.42	0.12	0.52
52.0	25.0	5.0	24.0	0.46	0.10	0.48
41.0	24.0	5.0	12.4*	0.30	0.12	0.59
36.5	19.2	4.9	18.0	0.49	0.13	0.53
25.2	12.7	3.9	13.4	0.53	0.15	0.50
17.3	8.7	2.7	9.8	0.57	0.16	0.50
11.7	5.9	1.9	7.4	0.63	0.16	0.50
8.0	3.7	2.0	5.2	0.65	0.25	0.46
5.6	2.7	1.3	4.2	0.75	0.23	0.48
	52·0 41·0 36·5 25·2 17·3 11·7 8·0	52·0 25·0 41·0 24·0 36·5 19·2 25·2 12·7 17·3 8·7 11·7 5·9 8·0 3·7	52·0 25·0 5·0 41·0 24·0 5·0 36·5 19·2 4·9 25·2 12·7 3·9 17·3 8·7 2·7 11·7 5·9 1·9 8·0 3·7 2·0	52·0 25·0 5·0 24·0 41·0 24·0 5·0 12·4* 36·5 19·2 4·9 18·0 25·2 12·7 3·9 13·4 17·3 8·7 2·7 9·8 11·7 5·9 1·9 7·4 8·0 3·7 2·0 5·2	52·0 25·0 5·0 24·0 0·46 41·0 24·0 5·0 12·4* 0·30 36·5 19·2 4·9 18·0 0·49 25·2 12·7 3·9 13·4 0·53 17·3 8·7 2·7 9·8 0·57 11·7 5·9 1·9 7·4 0·63 8·0 3·7 2·0 5·2 0·65	$\begin{array}{cccccccccccccccccccccccccccccccccccc$



Remarks. There is a considerable amount of variation in the suture spacing. Characteristics of this species are the rapid whorl expansion, coupled with a relatively narrow ventral lobe and nearly straight sides to the lateral lobe.



TEXT-FIG. 33. Dzhaprakoceras levis sp. nov. A, C, holotype, LZA 9399; A, ventral suture, H = 7.3; C, axial section. B, Ro 913; ventral suture, H = 1.0, W = 1.9.

Dzhaprakoceras levis sp. nov.

Plate 3, figures 4, 6; Plate 4, figures 2-4; Text-figures 32-33

Derivation of name. From the Latin levis (= smooth), with reference to the smooth internal surface of the conch.

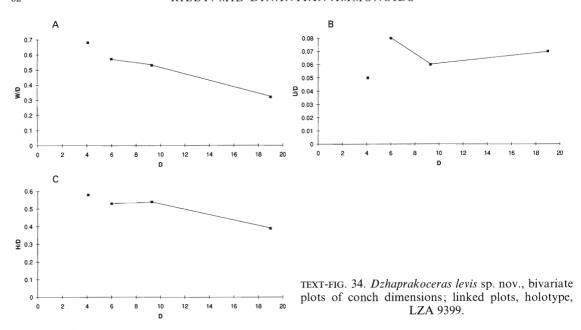
Material. Holotype LZA 9399, mid-adolescent, paratypes Ro 911–4, juveniles; all conches are septate, exfoliated and pyritic; from locality 3b.

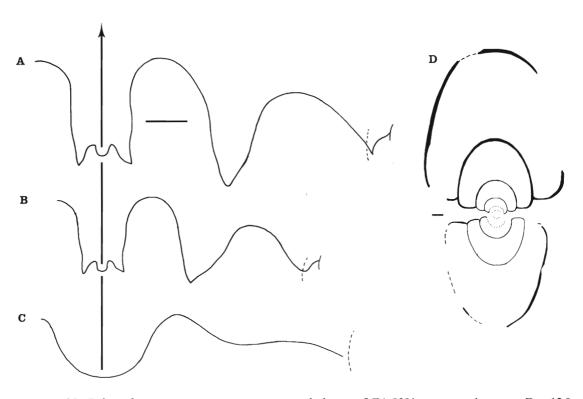
Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A small, involute, discoconic *Dzhaprakoceras* which lacks constrictions, has closely spaced sutures and a broad ventral lobe, with a high median saddle.

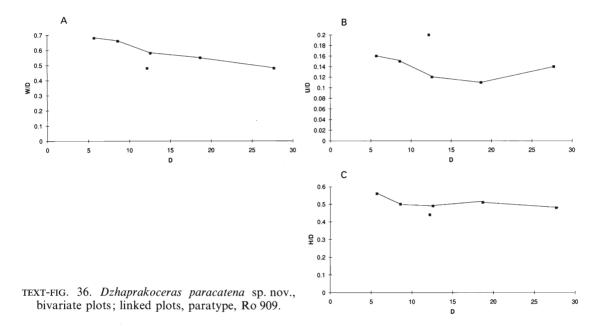
Description. The conch is small, discoconic and involute with rapid whorl expansion and an almost closed umbilicus. Juveniles are sphaeroconic to pachyconic. The external ornament is unknown and the internal surface of the conch is smooth. Constrictions are absent. The ventral lobe is broad and lyrate with a high median saddle (30 per cent. of ventral lobe depth). The lateral saddle is broad and rounded. The lateral lobe is narrow and acute with almost straight sides.

Specimen	D	Н	U	W	W/D	U/D	H/D
LZA 9399	19.0	7.4	1.4	6.0	0.32	0.07	0.39
LZA 9399	9.3	5.0	0.6	5.0	0.53	0.06	0.54
LZA 9399	6.0	3.2	0.5	3.4	0.57	0.08	0.53
Ro 913	4.1	2.4	0.2	2.8	0.68	0.05	0.58





TEXT-FIG. 35. *Dzhaprakoceras paracatena* sp. nov. A-C, holotype, LZA 9391; A, ventral suture, $D = 12\cdot0$, $H = 5\cdot9$, $U = 2\cdot4$, $W = 5\cdot6$; B, ventral suture, $H = 4\cdot0$, $W = 4\cdot7$; c, constriction, $H = 4\cdot9$, $W = 5\cdot3$. D, paratype, Ro 909; axial section.



Remarks. This species has similarities with Kozhimites planus Kusina, 1971, notably in the lateral conch compression and suture form, especially that of the ventral lobe. However, Kozhimites has a distinct ornament style displayed by regular constrictions in which the hyponomic salient is placed well forward of the lateral sinus. None of these features is seen in D. levis sp. nov.

Dzhaprakoceras paracatena sp. nov.

Plate 6, figures 9-10; Text-figures 35-36

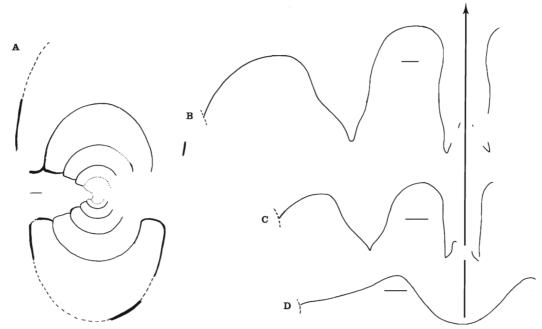
Derivation of name. From the similarity to D. catena sp. nov.

Material. Holotype, LZA 9391, fragmentary, pyritic, exfoliated, adolescent conch; from locality 3b; paratype Ro 913, fragmentary, septate, adolescent conch, preserved in micrite; from locality 4a.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A small discoconic, almost lenticular *Dzhaprakoceras* with early ontogenetic constriction loss and a deep narrow lateral lobe.

Description. The conch is discoconic, almost lenticular, with a small umbilicus, which has gently sloping sides. Juveniles are pachyconic; whorl expansion is rapid. The external ornament is not preserved. The internal ornament consists of regularly spaced, weak bisinuous ridges; these become strongly developed in the ventrolateral region where vesicular ridges occur corresponding with the linguae. The hyponomic sinus is relatively narrow, rounded and deep. Constrictions occur up to D=11; it is estimated that there are four per whorl; they are strongly incised and follow the internal ornament. The suture shows a subparallel-sided ventral lobe which is slightly lyrate; the lateral saddle is broad and rounded, but the lateral lobe is very narrow with flexed sides and a subacute termination.



Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 909	27.7	13.2	3.8	13.3	0.48	0.14	0.48
Ro 909	18.7	9.6	2.0	10.2	0.55	0.11	0.51
Ro 909	12.6	6.2	1.5	7.3	0.58	0.12	0.49
Ro 909	8.6	4.3	1.3	5.7	0.66	0.15	0.50
Ro 909	5.7	3.2	0.9	3.9	0.68	0.16	0.56
LZA 9391	12.2	5.4	2.5	5.9	0.48	0.20	0.44

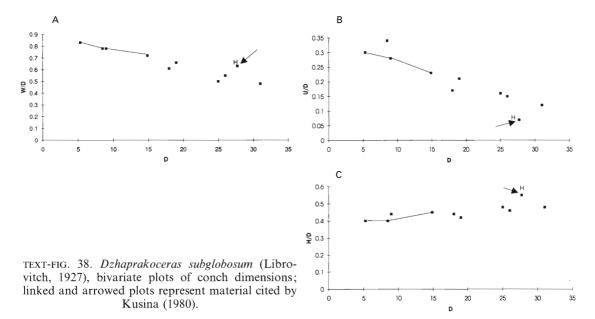
Remarks. This new species is similar to *D. catena* sp. nov. but differs in the early loss of constrictions and the very narrow, deep lateral lobe. Similar also is *D. levis* sp. nov., which has, however, a wider ventral lobe and lacks constrictions.

Dzhaprakoceras subglobosum (Librovitch, 1927)

Plate 4, figures 11–12; Text-figures 37–38

1927	Muensteroceras subglobosum Librovitch, p. 35, pl. 6, fig. 7; pl. 7, figs 1-2; text-fig. 17.
1941	Muensteroceras subglobosum Librovitch; Librovitch, p. 143, pl. 37, fig. 3; text-fig. 25.
non 1961	Muensteroceras subglobosum Librovitch; Pareyn, p. 104, pl. 8, figs 27–28.
1968	Dzhaprakoceras subglobosum (Librovitch); Popov, p. 88, pl. 3, figs 4-5; pl. 6, fig. 5; text-fig. 16.
1980	Dzhaprakoceras subglobosum (Librovitch); Kusina, p. 57–8, pl. 4, fig. 6.

Lectotype. Designated by Kusina (1980), specimen 100/2473, TSNIGRI Museum, St Petersburg; no reference to original figure.



Material. Twelve conches preserved in wackestone, mostly exfoliated and septate, some fragmentary, representing various growth stages. Specimens Ca 1592, 1863, DM 1265, 1269, 1275 and Ro 5177–8, 5180–1, 5183–5; from locality 2.

Range. Late Tournaisian, middle FA Zone; Clitheroe Limestone Formation (Bellman Limestone Member). Known also from the middle FA Zone in Tien-Shan (CIS) and possibly from the BB Zone in the north Urals (CIS).

Diagnosis. A small, pachyconic, subglobose *Dzhaprakoceras* with a narrow ventral lobe and maintaining a nearly constant umbilical diameter after early adolescence.

Description. The conch is pachyconic, subglobose and of moderate size. The umbilicus is open, of nearly constant diameter, but becomes narrower during the sixth or seventh whorl. The ornament consists of very delicate transverse striae (not well preserved). Constrictions are present; they are regular, about five per whorl, and bisinuous; they show a well developed hyponomic sinus, but a shallower lateral sinus and fail after $H=9\,\mathrm{mm}$. The ventral lobe has slightly divergent sides in the immature conch, but becomes narrow deep and lyrate when mature. The lateral saddle is evenly rounded and slightly inflated orad. The lateral lobe is moderately broad with flexed sides.

Specimen	D	Н	U	W	W/D	U/D	H/D
DM 1265	31.0	15.0	4.0	15.0	0.48	0.12	0.48
DM 1269		13.0	4.5	15.0			
DM 1275	26.0	12.0	3.8	14.5	0.55	0.15	0.46
Ro 5177	25.0	12.0	4.0	12.5	0.50	0.16	0.48
Ca 1663	19.0	8.0	4.0	12.5	0.66	0.21	0.42
Ca 1592	18.0	8.0	3.0	11.0	0.61	0.17	0.44
Ro 5183	9.0	4.0	2.5	7.0	0.78	0.28	0.44

Remarks. The species is very similar to *D. djaprakense* which, however, is more dorso-ventrally compressed and has a larger umbilicus. Specimens with the specimen number prefix Ca and DM were incorrectly assigned to *Muensteroceras difficile* by Ramsbottom in Earp et al. (1961, p. 179).

Dzhaprakoceras undulatum sp. nov.

Plate 4, figures 8–9; Plate 5, figure 1; Plate 7, figure 2; Text-figures 39–40

Derivation of name. From the Latin undulatus (= undulated) with reference to the ornament.

Material. Holotype Ro 925, paratypes Ro 921–3, 926–7, partly exfoliated conches at various growth stages preserved in wackestone; from locality 4; Ro 928–9, pyritic septate adolescent moulds; from locality 3.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A small, relatively evolute, subdiscoconic *Dzhaprakoceras* with ontogentic constriction loss and the progressive development of a strongly bisinuous transverse ornament.

Description. The conch is small, subdiscoconic and relatively evolute with slow whorl expansion. Juveniles are initially cadiconic but later pachyconic. The external ornament consists of weak transverse bisinuous striae with well developed lateral and hyponomic sinuses. The internal ornament is also well developed and consists of vesicular ridges which follow the external ornament. There are five constrictions per whorl; they are deeply incised and follow the external ornament, but fail at maturity. The ventral lobe is deep, narrow and lyrate. The lateral saddle is broad and rounded. The lateral lobe is broad and acute with almost straight sides.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 925	27.5	13.5	5.0	12.0	0.44	0.18	0.49
Ro 928	10.5	4.0	2.5	5.2	0.50	0.24	0.38
Ro 929	6.5	2.5	2.0	3.5	0.54	0.31	0.31
Ro 927	17.5	7.5	3.5	9.0	0.51	0.20	0.43
Ro 927	12.6	5.4	2.8	6.8	0.54	0.22	0.43
Ro 927	9.3	3.7	2.2	5.5	0.59	0.24	0.40
Ro 927	6.9	2.8	1.7	4.5	0.65	0.25	0.41

Remarks. This new species is very similar to D. deflexum which, however, is more involute, retains constrictions and has a flexed ventrad side to the lateral lobe.

Dzhaprakoceras westheadi sp. nov.

Plate 5, figures 8–9; Plate 6, figures 2–3; Plate 7, figures 4–5; Text-figures 41–42

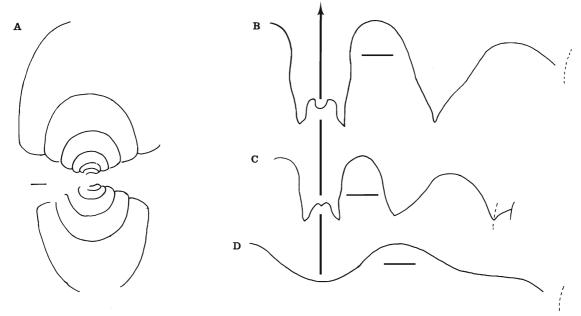
Derivation of name. In honour of the late Mr S. Westhead, a keen amateur geologist from Clitheroe, famous for his excellent collection of local crinoids.

Material. Holotype, LZA 9407, exfoliated, adolescent, pyritic conch; from locality 3b. Paratypes; Ro 917 part exfoliated conch preserved in micrite; from locality 4a; and LZA 9386, 9393, 9401–3, 9407–8, 9410–1, 9414, Ro 919, Ro 5171, exfoliated pyritic conches; from locality 3b.

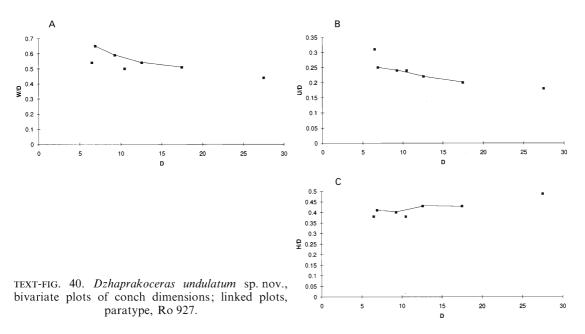
Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation, Leagram Mudstone Member.

Diagnosis. A medium-sized, subdiscoconic Dzhaprakoceras with strongly bisinuous transverse ornament.

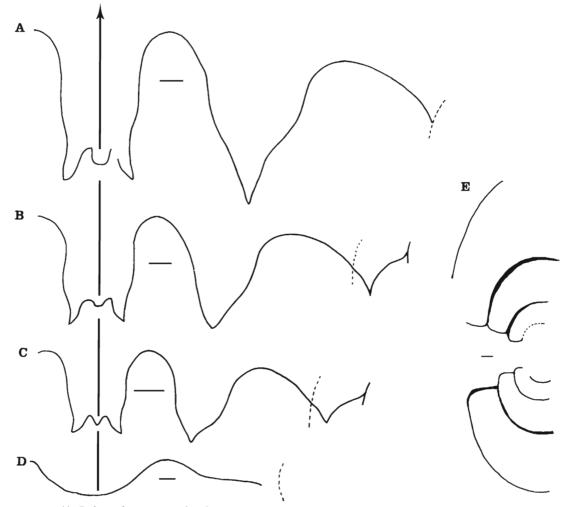
Description. The conch is subdiscoconic with moderate whorl expansion and an average-sized umbilicus with steep sides. The juvenile whorls are cadiconic. The external ornament is not preserved, but the internal



TEXT-FIG. 39. Dzhaprakoceras undulatum sp. nov. A, paratype, Ro 927; axial section, B, D, paratype, Ro 923; B, ventral suture, H = 5.8; D, constriction, H = 5.8. C, paratype, Ro 928; ventral suture, H = 4.4, W = 5.6.

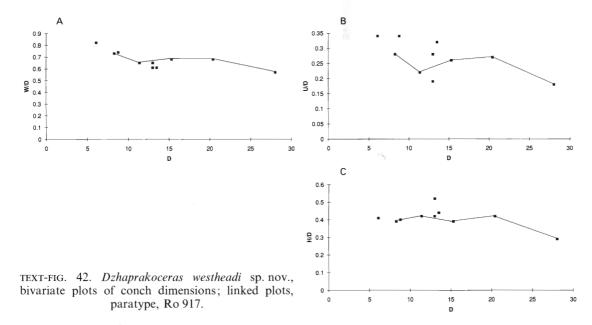


ornament consists of weak, strongly bisinuous transverse ridges which are well developed on the ventrolateral region, giving rise to vesicular ridges which correspond to the linguae. The hyponomic sinus is broad and relatively deep. There are five, well developed constrictions per whorl, which follow the path of the internal ornament. The suture consists of a narrow, lyrate ventral lobe with slightly divergent sides (orad). The lateral saddle is narrow and sharply rounded; the lateral lobe is deeper than the ventral and is narrow with flexed sides and an acute termination.



TEXT-FIG. 41. *Dzhaprakoceras westheadi* sp. nov. A, E, paratype, Ro 917; A, ventral suture, $H=11\cdot4$, $W=15\cdot3$; E, axial section. B, D, holotype, LZA 9407; B, ventral suture, $H=8\cdot1$, $W=12\cdot8$; D, constriction, $H=9\cdot1$, $W=13\cdot4$. C, paratype, LZA 9403; ventral suture, $H=5\cdot7$, $W=8\cdot4$.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 917	28.0	8.0	5.0	16.0	0.57	0.18	0.29
Ro 917	20.4	8.5	5.5	13.8*	0.68	0.27	0.42
Ro 917	15.3	6.0	4.0	10.4*	0.68	0.26	0.39
Ro 917	11.4	4.8	2.5	7.4*	0.65	0.22	0.42
Ro 917	8.3	3.2	2.3	6.1*	0.73	0.28	0.39
LZA 9402	13.5	5.9	4.3	8.3	0.61	0.32	0.44
LZA 9401	13.0	5.4	3.7	8.5	0.65	0.28	0.42
Ro 919	13.0	6.8	2.5	7.9	0.61	0.19	0.52
LZA 9408	8.8	3.5	3.0	6.5	0.74	0.34	0.40
LZA 9410	6.1	2.5	2.1	5.0	0.82	0.34	0.41



Remarks. This species is most similar to D. bellmanense sp. nov., but differs in having greater whorl inflation, more strongly bisinuous internal ornament and a smaller umbilicus. Juveniles and early adolescents are similar to D. flexum, but differ from that species in having a more divergent (orad) ventral lobe at comparable diameters.

Family Pericyclidae Hyatt, 1900

Diagnosis (emended). Small to large ammonoids in which the conch varies from subsphaeroconic to discoconic, rarely oxyconic, and may be involute or evolute. The external ornament is mainly transverse, simple or branched, often strong with ribs and/or tubercles; spiral ornament is sometimes present. The internal ornament is transverse and follows the external. The ventral lobe varies from narrow to moderate width, with parallel, divergent or convergent (orad) sides. The lateral saddle may be rounded, subacute or acute, but the lateral lobe is always acute.

Range. Tournaisian to Asbian.

Remarks. The Pericyclidae can be split into two groups based on the shape of the lateral saddle. The following new subfamilies are introduced below:

Ammonellipsitinae subfam. nov. (lateral saddles acute to subacute); late Tournaisian to late Chadian

Pericyclinae Hyatt, 1900 [nom. trans. Riley] (forms with well rounded lateral saddles); Tournaisian to Asbian

Subfamily AMMONELLIPSITINAE subfam. nov.

Diagnosis. Pericyclidae which have acute or subacute lateral saddles, divergent sides (orad) to the ventral lobe and a serpenticonic juvenile stage.

Range. Late Tournaisian to late Chadian.

Remarks. The following genera are included:

Ammonellipsites Parkinson, 1822; late Tournaisian to late Chadian.

Fascipericyclus Turner, 1948; late Tournaisian to late Chadian.

Helicocyclus Schindewolf, 1951a; late Chadian. Stenocyclus Schindewolf, 1951a; late Tournaisian.

The erection of the new subfamily restricts *Ammonellipsites* to the species group of which *Ellipsolithes funatus* Sowerby, 1813 is the type. Gordon's (1957, 1964) treatment of the above genera as subgenera of *Ammonellipsites* is not adopted here.

Genus AMMONELLIPSITES Parkinson, 1822.

Type species. Ellipsolithes functus Sowerby, 1813, pl. 32.

Diagnosis. The conch is commonly pachyconic, but may be serpenticonic or cadiconic; juveniles are serpenticonic. The external and internal ornament consists of transverse ribs which are usually simple, but may show irregular bifurcation; spiral lirae may occur externally between the ribs. The ventral lobe has strongly divergent sides (orad) and a high median saddle. The lateral saddle and lateral lobe are acute.

Range. Late Tournaisian to late Chadian.

Remarks. The following species are included:

- A. ballardensis Gordon, 1964; Grand Falls Chert Member, Boone Formation, Kansas, USA, late Tournaisian or early Viséan (texanus Conodont Zone).
- A. clitheroensis sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- A. djaprakensis (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS; late Tournaisian.
- A. doohylensis (Foord, 1903); Carboniferous Limestone, Doohyle, Co. Limerick, Ireland, late Tournaisian.
- A. funatus (Sowerby, 1813); Blackrock limestones, near Cork, Ireland, mid-Dinantian.
- A. kayseri (Schmidt, 1925); Erdbach Limestone, CuIIγ, Germany, mid-Dinantian.
- A. kochi (Holzapfel, 1889); Erdbach Limestone, CuII₂, Germany, mid-Dinantian.
- A. multicostatus (Foord, 1903); Carboniferous Limestone, Middleton, Ireland, late Tournaisian.
- A. nikitini (Librovitch, 1927); Dzhaprak Suite, Tien-Shan, Kirgizya, CIS, late Tournaisian.

Ammonellipsites is closest to Stenocyclus, a genus which develops ontogenetically a carinate venter, absent even in gerontic specimens of Ammonellipsites. Mature Fascipericylus may have a ribbing character similar to that of Ammonellipsites with infrequent rib bifurcation; however, the earlier ontogeny of Fascipericyclus shows strong, regular rib bifurcation and this feature, coupled with the low median saddle, serves to distinguish the two genera.

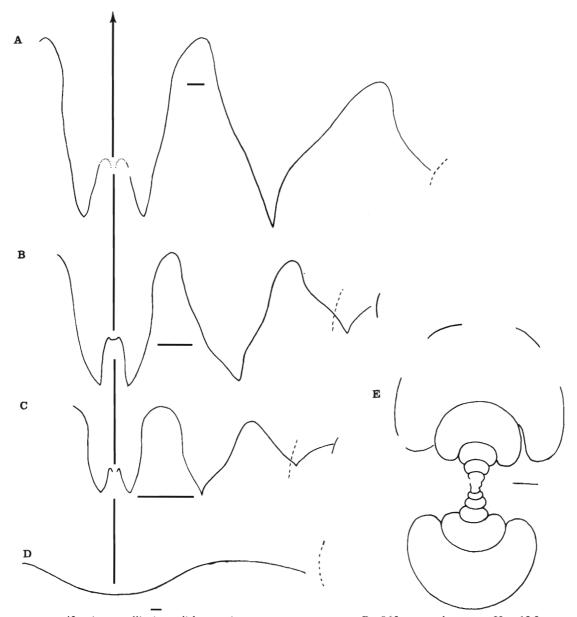
Ammonellipsites clitheroensis sp. nov.

Plate 1, figures 1-3, 5-6, 8, 11; Plate 2, figure 4; Text-figures 43-44

Derivation of name. From Clitheroe, the type locality.

Material. Holotype Ro 864, exfoliated late adolescent, preserved in wackestone; from locality 4a; and paratypes LZA 9404, 9412 and 9452, pyritic, septate, exfoliated adolescents; from locality 3b. Also paratypes Ro 863–9, adolescent and mature conchs, some exfoliated and septate, preserved in wackestone; from locality 4a.

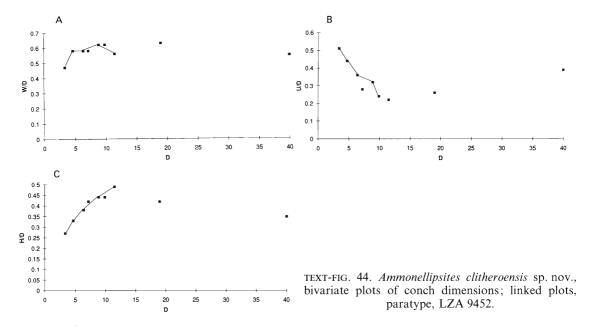
Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Whitemore Limestone and Leagram Mudstone members).



TEXT-FIG. 43. Ammonellipsites clitheroensis sp. nov. A, paratype, Ro 865; ventral suture, $H = 12\cdot 2$. B, E, paratype, LZA 9452, B, ventral suture, $H = 4\cdot 9$, $W = 6\cdot 5$; E, axial section. C, paratype, Ro 866; ventral suture, $D = 5\cdot 7$, $H = 2\cdot 0$, $U = 2\cdot 7$, $W = 3\cdot 6$. D, paratype, Ro 863; transverse ornament, $H = 14\cdot 9$, $W = 5\cdot 1$.

Diagnosis. An Ammonellipsites with a relatively prolonged juvenile serpenticonic stage and ontogenetic development of constrictions.

Description. The conch is pachyconic with a small umbilicus and slow whorl expansion. Juveniles are serpenticonic until the fourth whorl, when the conch becomes cadiconic. The external ornament consists of flat-topped transverse ribs with a broad, shallow hyponomic sinus. The internal ornament consists of asymmetrical ribs which correspond with the external ornament, but fail low on the flanks, especially at maturity.

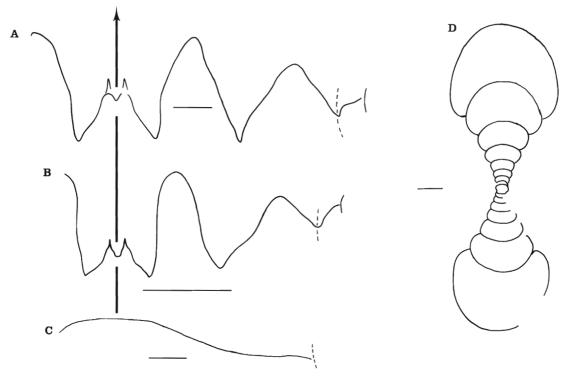


Constrictions develop at maturity and have been seen only on fragmentary specimens; they are deeply incised and follow the ornament; they are visible externally and are enhanced by suppression of the adjacent ribs. The ventral lobe is broad and deep, with strongly divergent (orad) sides and a high median saddle. The lateral saddle and lateral lobe are broad and acute with flexed sides.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 863	40.0	14.0	15.5	22.0	0.55	0.39	0.35
Ro 864	19.0	8.0	5.0	12.0	0.63	0.26	0.42
LZA 9452	11.5	5.5	2.5	6.4	0.56	0.22	0.49
LZA 9452	8.9	3.9	2.8	· 5·5	0.62	0.32	0.44
LZA 9452	6.4	2.5	2.3	3.7	0.58	0.36	0.38
LZA 9452	4 ·7	1.5	2.1	2.7	0.58	0.44	0.33
LZA 9452	3.4	0.9	1.7	1.6	0.47	0.51	0.27
LZA 9404	9.9	4.4	2.4	6.1	0.62	0.24	0.44
Ro 866	7.2	3.0	2.0	4.2	0.58	0.28	0.42

Remarks. This new species is very similar to A. kochi, which, however, from Holzapfel's (1899) and Schmidt's (1924) descriptions, appears to have a shorter juvenile serpenticonic stage, with spiral lirae, and apparently lacks constrictions when mature. Unfortunately, A. kochi has never been adequately described and the type material was destroyed during World War Two, hence comparison is difficult. The form which occurs in the Tober Colleen Formation (Rush Slates) of the Dublin Basin, attributed to A. kochi by Smyth (1951), has a much longer serpenticonic juvenile stage than A. clitheroensis and differs in ornament strength and sutural characters. The Irish form appears to be a new species.



TEXT-FIG. 45. Helicocyclus divergens sp. nov. A, C–D, holotype LZA 9422; A, ventral suture, H = 4·6, W = 4·8; C, constriction; D, axial section. B, paratype, LZA 9388; ventral suture, D = 7·0, H = 1·8, W = 2·9.

Genus HELICOCYCLUS Schindewolf 1951a

Type species. Helicocyclus gracillimus Schindewolf, 1951a.

Diagnosis (emended). Ammonellipsitins with predominantly small (D < 55 mm) laterally compressed, evolute, discoconic conches, which are serpenticonic until late adolescence. Whorl expansion is slow. The external ornament consists of simple or branched, relatively weak, transverse ribs which bend orad over the flanks and venter, with consequential loss of a hyponomic sinus. The internal ornament, where present, consists of weak ribs which follow the external ornament. Constrictions occur at all diameters and follow the external ornament. The ventral lobe has divergent sides (orad), acute prongs and a broad, high median saddle. The lateral saddle is acute to subacute. The lateral lobe is acute with flexed sides.

Range. Late Chadian.

Remarks. The following species are included:

- H. divergens sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- H. gracillimus Schindewolf, 1951a; Winterberg, Harz, Germany, Viséan.
- *H. involutus* Kusina *in* Shumansky and Kusina, 1977; Kosvinsky Horizon, Kozhim River, north Urals, Russian Federation, CIS, late Chadian.

This genus was erected using juvenile material and hence the diagnosis has been modified. *Pericylus*

tianshanicus Librovitch, 1927 and the synonymous *P. evolutus* Librovitch, 1927, although included in *Helicocyclus* by Schindewolf (1951a) and Kusina (1980), have much higher whorl expansion rates than *Helicocyclus*, and have a low median saddle and a subacute lateral saddle. The author therefore follows Popov (1968) in placing *P. tianshanicus* within *Fascipericyclus*.

Helicocyclus divergens sp. nov.

Plate 1, figures 4, 9, 12; Text-figures 45-46

Derivation of name. From the divergent-sided ventral lobe.

Material. Holotype LZA 9422 and paratypes LZA 9383, 9421, 9424, 9426–30, 9433–4, 9445, RH 1525–35 and Ro 5400; all pyritic internal septate moulds representing all growth stages; from locality 3b.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Diagnosis. A Helicocyclus with a broad and strongly divergent-sided (orad) ventral lobe.

Description. the conch is subdiscoconic, but serpenticonic below $D=10\,\mathrm{mm}$ and has a large umbilicus, convex flanks and a narrow, arched, rounded venter. Whorl expansion is slow. The external ornament is unknown. The internal ornament consists of delicate, simple transverse ribs which slope orad over the flanks and venter (spaced at about 0.75 mm on the venter at $D=12\,\mathrm{mm}$); a hyponomic sinus is not present. There are four constrictions per whorl throughout growth, these follow the internal ornament. The ventral lobe is broad with strongly divergent sides (orad) and a broad, high (c. 40 per cent. of ventral lobe depth) median saddle. The lateral saddle is narrow and acute. The lateral lobe is acute with slightly flexed sides.

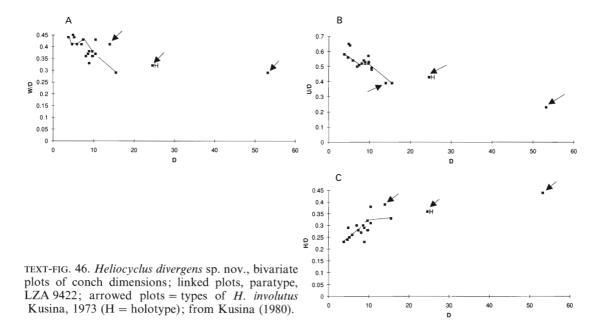
Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
LZA 9426	10.5	4.0	5.0	4.5	0.43	0.48	0.38
LZA 9424	10.5	3.3	5.1	3.9	0.37	0.49	0.31
LZA 9429	9.8	2.7	5.2	3.5	0.36	0.53	0.28
LZA 9428	9.7	2.7	5.5	3.5	0.36	0.57	0.28
RH 1531	8.9	2.6	4.7	3.4	0.38	0.52	0.29
LZA 9445	8.9	2.0	4.7	2.9	0.33	0.53	0.23
LZA 9427	8.6	2.6	4.6	3.2	0.37	0.54	0.30
RH 1534	8.1	2.2	4.2	2.1	0.26	0.52	0.27
LZA 9388	7.0	2.1	3.5	2.9	0.41	0.50	0.30
LZA 9434	5.2	1.3	3.3	2.3	0.44	0.64	0.25
LZA 9383	4.9	1.4	3.2	2.2	0.45	0.65	0.29
LZA 9422	15.5	5.0	6.0	4.5	0.29	0.39	0.33
LZA 9422	9.7	3.1	5.1	3.7	0.38	0.52	0.32
LZA 9422	7-4	2.1	3.8	3.2	0.43	0.51	0.28
LZA 9422	5.9	1.5	3.2	2.4	0.41	0.54	0.26
LZA 9422	4.7	1.2	2.6	2.0	0.41	0.56	0.24
LZA 9422	3.8	0.9	2.2	1.7	0.44	0.58	0.23

Remarks. H. divergens sp. nov. is very similar to H. involutus; however, the broad divergent-sided ventral lobe of the new species is distinctive. H. gracillimus does not show as much whorl arching at comparable diameters and lacks a broad divergent-sided ventral lobe.

Subfamily PERICYCLINAE Hyatt, 1900 [nom. trans. Riley]

Diagnosis. Pericyclids with broad, rounded lateral saddles and a ventral lobe with subparallel, parallel or convergent (orad) sides. The ornament sometimes fails during late ontogeny.



Range. Tournaisian to Asbian.

Remarks. The division of the pericyclids by previous workers, e.g. Turner (1948), Schindewolf (1951b) and Gordon (1957), was based on the divergence of the ventral lobe sides and is not followed here. Species groups previously treated as subgenera, e.g. Goniocyclus, are here treated as genera. The following genera are included:

Asiacyclus Librovitch, 1962; Tournaisian

Bollandites Bisat, 1952; Arundian to Asbian.

Goniocyclus Gordon, 1986; mid-Tournaisian.

Hammatocyclus Schindewolf, 1951b; ?late Tournaisian to Arundian.

Neopericyclus Popov, 1965; late Tournaisian

Parahammatocyclus gen. nov.; Arundian

Pericyclus Mojsisovics von Mojsvár, 1882; Tournaisian.

Polaricyclus Riley, 1991; late Tournaisian.

Rhiphaeocyclus Kusina, 1973; Arundian.

Rotopericyclus Turner, 1948; Tournaisian to late Chadian.

Bollandites, formerly placed in the Muensteroceratidae, is here included with the pericyclins since it has many similarities, particularly in general conch and suture form, with Hammatocyclus.

Genus BOLLANDITES Bisat, 1952

Type species. Beyrichoceratoides castletonense Bisat, 1924.

Diagnosis. The conch is of variable size (D \sim 15–100 mm), cadiconic, platyconic or subsphaeroconic with serpenticonic early whorls. The external ornament consists of simple, weak, transverse striae with a gentle lateral sinus and a broad feeble hyponomic sinus; the ornament is coarser in juveniles; spiral lirae are rarely present. An internal ornament is present in juveniles and consists of simple transverse ribs which follow the external ornament, usually the internal ornament does not persist into adolescence. Strongly incised constrictions are present in juveniles, but these commonly fail

during late adolescence. The ventral lobe is relatively broad and has divergent (orad), flexed sides. The lateral saddle is rounded and the lateral lobe is acute with flexed sides.

Range. Arundian to Asbian.

Remarks. The following species are included:

- B. baoshanensis Liang and Zhu, 1988; Qingshigou Baoshan County, western Yunnan, People's Republic of China, Viséan.
- B. castletonensis (Bisat, 1924); 'reef limestone', Treak Cliff, northern England, Asbian.
- B. donetsensis Kusina, 1991; Zone C₁V_e, base of the Stylsky Horizon, Donets Basin, Ukraine, ?Asbian.
- B. mediocris Kusina, 1991; Zone C₁V_e, base of the Stylsky Horizon, Donets Basin, Ukraine, ?Asbian.
- B. phillipsi (Foord and Crick, 1897); Carboniferous Limestone, Yorkshire, northern England, ?Asbian.
- B. sulcatus Bisat, 1952; Pendleside Limestone Formation, Barnoldswick, north-west England, late Asbian.
- B. tianshanicum (Popov, 1968); base of Kara-Kiya Suite, Tien-Shan, Kirgizya, CIS, Arundian.
- B. trevallynensis (Campbell, Brown and Coleman, 1983); Bonnington Siltstone, New South Wales, Australia, Viséan.
- B. umbilicatus (Bisat, 1934); 'reef limestones' Newton Gill, Hellifield, north-west England, late Asbian.
- B. varians sp. nov.; Hodderense Limestone Formation, Leagram north-west England, late Holkerian.

Bollandites bashatchensis Popov, 1965, should be transferred to Bollandoceras (see p. 40). B. kokdzharensis Popov, 1965 and B. narynensis Popov, 1965, judged by their conch form and suture, fall within Dzhaprakoceras (see p. 50). Bollandites bowsheri Gordon, 1957, cannot be placed in Bollandites since its transverse ornament progrades over the venter; the generic status of this species is at present unclear but it probably represents a juvenile Goniatites sp. Bollandites kiligwae Gordon, 1957, was based on a suite of juveniles; this species lacks a serpenticonic juvenile stage, thus excluding it from Bollandites and it may belong to Beyrichoceratoides or Dzhaprakoceras. Bollandites boreus Kusina, 1980, is similar in conch form to Bollandites, but has retrograde constrictions and a ventral lobe without flexed sides, this species is here placed in Rotopericyclus (see p. 80). Bollandites trevallynensis, originally described under Beyrichoceras, has a transverse ornament with feeble sinuses and a suture typical of Bollandites.

Bollandites varians sp. nov.

Plate 7, figures 1, 3, 7, 9, 11; Plate 8, figure 14; Text-figures 47–48

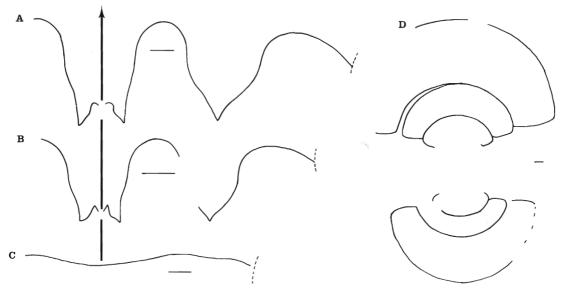
Derivation of name. From the variable shape of the conch.

Material. Holotype Ro 5187, solid septate, exfoliated conch; from locality 11; and paratypes Ro 934–8; from locality 9; Ro 940–53; from locality 10; LZA 7865; from locality 12; solid and fragmentary conches at all growth stages, some juveniles showing ornament.

Range. Late Chadian, upper BB Zone; Hodderense Limestone Formation.

Diagnosis. A subsphaeroconic to pachyconic *Bollandites* with a broad venter and wide umbilicus throughout growth.

Description. The conch is variable, subsphaeroconic to cadiconic at maturity, but always serpenticonic before the fourth whorl; whorl expansion is slow. The venter may broaden considerably above D = 9 mm. The



TEXT-FIG. 47. Bollandites varians sp. nov. A, paratype, LZA 7865; ventral suture, H = 6·7, W = 11·8. B-C, holotype, Ro 5187; B, ventral suture, H = 4·0, W = 8·6; C, constriction, same dimensions as B. D, paratype, Ro 940; axial section.

umbilical walls are steep-sided. The external ornament consists of weak, simple, transverse striae, spaced at four per mm at W=7 mm, but is much coarser in juveniles. The hyponomic sinus is broad and very weak; a lateral sinus is present in some adults. Weak transverse ribs are present on internal casts below D=6 mm. Constrictions are retained throughout growth, but are particularly strong in juveniles (D=<6 mm) and run parallel to the external ornament. The suture has rounded saddles and a divergent (orad) ventral lobe with flexed sides; the lateral lobe is deep and relatively narrow, with flexed sides.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
Ro 940	32.0	13.0	10.0	22.0	0.69	0.31	0.41
LZA 7865	21.0	10.0	4.5	12.0	0.57	0.21	0.48
Ro 941	19.5	6.0	11.0	11.0	0.56	0.56	0.31
Ro 939	16.0	4.5	6.5	12.5	0.78	0.41	0.28
Ro 942	14.0	5.0	6.0	10.0	0.71	0.43	0.36
Ro 951	5.3	1.6	3.3	3.5	0.66	0.62	0.30

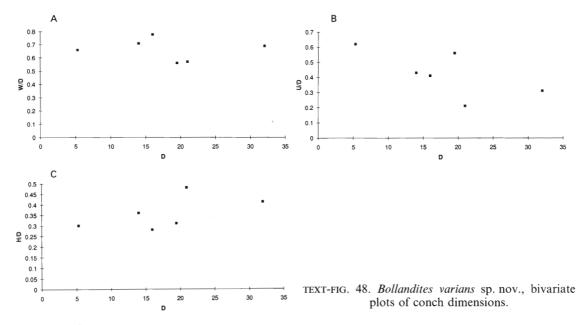
Remarks. The conch form of this stratigraphically early species is in stark contrast to that of late Asbian forms, which show lateral whorl compression, accelerated ontogeny and arched venters.

Bollandites? sp.

Plate 3, figure 7

Material. Ro 5076a, a fragmentary, late adolescent, exfoliated conch; from locality 6.

Range. Arundian, lower BB Zone; Hodder Mudstone Formation (Chaigley Limestone).



Description. The conch is apparently cadiconic with a wide umbilicus and a steep umbilical shoulder. The external ornament is not preserved. The internal ornament, which is preserved only low on the flanks, consists of weak radial undulations. There are no sutures.

Remarks. This specimen is tentatively placed in Bollandites on account of the general conch form.

Genus PARAHAMMATOCYCLUS gen. nov.

Type species. Parahammatocyclus chaigleyensis sp. nov.

Diagnosis. Conch small, D = < 20 mm, subsphaeroconic in the adult, but initially serpenticonic with an intermediate cadiconic stage. During the cadiconic stage, umbilical tubercles are present. The ornament consists of simple, coarse transverse striae; these are nearly radial, with only a weak hyponomic sinus. The external suture has a divergent-sided (orad) ventral lobe, a rounded lateral saddle and a deep, broad lateral lobe, with flexed sides and an acute mammilate termination.

Range. Arundian.

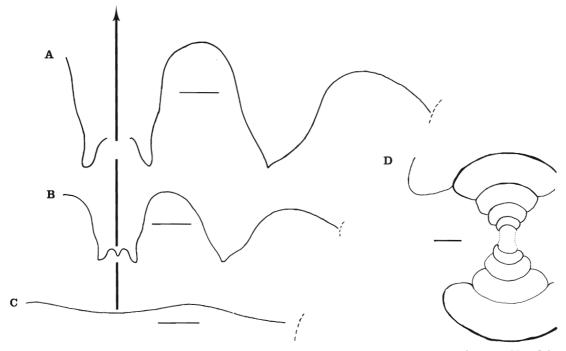
Remarks. This genus is intermediate between Bollandites and Hammatocyclus. It differs from the former in the presence of umbilical tubercles, and from the latter by ontogenetic loss of tubercles and the presence of a divergent-sided (orad) ventral lobe. Bearing in mind these morphological relationships and the stratigraphical position of Parahammatocyclus, it is reasonable to propose that the new genus represents an intermediate in a phyletic lineage between Hammatocyclus and Bollandites. The new genus is at present monospecific.

Parahammatocyclus chaigleyensis gen. et. sp. nov.

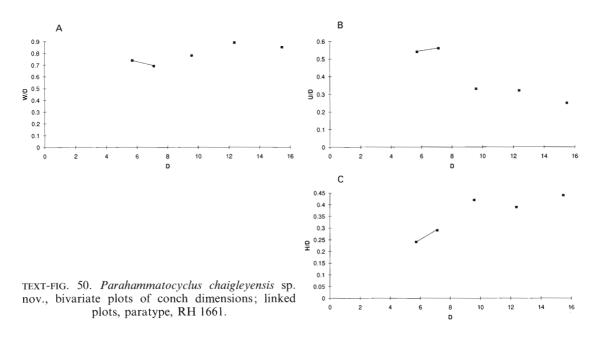
Plate 3, figures 1-3, 8; Plate 6, figures 1, 4, 12; Plate 7, figure 8; Text-figures 49-50

Derivation of name. From Chaigley, the hamlet near where the holotype was found.

Material. Holotype Ro 5127, partly testiferous late adolescent; from locality 6; paratypes, Ro 3721 (fragmentary testiferous adolescent conch; from locality, 5), Ro 3723 (exfoliated, partly crushed adolescent;



TEXT-FIG. 49. Parahammatocyclus chaigleyensis sp. nov. A, C, paratype, Ro 1657; A, ventral suture, H = 5.6; C, constriction, H = 3.2. B, paratype, RH 1631; ventral suture (reversed image), H = 3.5, W = 5.8. D, paratype, RH 1661; transverse section.



from locality, 7), Ro 5105 (fragmentary, partly testiferous adolescent; from locality, 5), Ro 5137 (mainly exfoliated adolescent; from locality 6), RH 1631 and 1657 (exfoliated adolescent conches; from locality 6) and RH 1661, juvenile embedded in packstone and sectioned axially. All specimens are ankeritic.

Range. Arundian, lower BB Zone; Hodder Mudstone Formation (Chaigley Limestone Member).

Diagnosis. As for genus.

Description. The conch is small (D = < 20 mm), serpenticonic in the early whorls, subsphaeroconic when mature, with an intermediate cadiconic stage. The umbilicus is subquadrate at maturity. The external ornament consists of coarse, simple, transverse striae, which show only a feeble hyponomic sinus. Umbilical tubercles are present during adolescence, but these are lost at maturity. The internal ornament consists of weak transverse undulations which follow the external ornament, except near the umbilical tubercles, where the undulations sometimes branch. The umbilical tubercles are more strongly developed internally than externally. At least four constrictions, moderately incised, are present; they follow the external ornament and are preserved also on the exterior of the conch. The external suture consists of a divergent (orad) ventral lobe, rounded lateral saddle and a deep, broad, acute lateral lobe with strongly flexed sides.

Dimensions.

Specimen	D	Н	U	W	W/D	U/D	H/D
RH 1651	15.5	6.8	3.8	13.2	0.85	0.25	0.44
Ro 5127	12.4	4.8	4.0	11.0	0.89	0.32	0.39
Ro 5144	9.6	4.0	3.2	7.5	0.78	0.33	0.42
RH 1661	7.1	2.1	4.0	5.0	0.69	0.56	0.29
RH 1661	5.7	1.4	3.1	4.2	0.74	0.54	0.24

Remarks. The branching of the transverse undulations on the internal ornament further suggests the pericyclid affinities of this new taxon.

Genus ROTOPERICYCLUS Turner, 1948

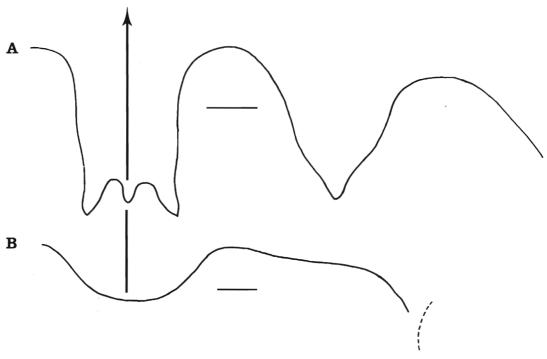
Type species. Pericyclus rotuliformis Crick, 1899.

Diagnosis. Small to large pericyclinae with subdiscoconic conches. The ornament consists of transverse ribs or striae which pass retrograde over the flanks and venter, giving rise to a deep hyponomic sinus. Subordinate spiral lirae may be present, giving rise to crenulation. Constrictions are deeply incised, regular and numerous, visible both internally and externally. The ventral lobe of the suture has subparallel sides which diverge very slightly orad; the lateral saddle is rounded and the lateral lobe acute.

Range. Tournaisian to late Chadian.

Remarks. The following species are included:

- R. boreus (Kusina, 1980); Kosvinsky Suite, Kozhim River, north Urals, Russian Federation, CIS, late Chadian
- R. multiplicatus (Delépine, 1940a); ?Calcaire de Vaux, Belgium, late Tournaisian.
- R. postremus sp. nov.; Hodder Mudstone Formation, Clitheroe, north-west England, late Chadian.
- R. rotuliformis (Crick, 1899); Carboniferous Limestone, Saint Doughlas, Ireland, Courceyan.



TEXT-FIG. 51. Rotopericyclus postremus sp. nov.; holotype, Ro 933. A, ventral suture, H = 7.2, W = 7.2; B, constriction, same dimensions as A.

Rotopericyclus postremus sp. nov.

Plate 1, figures 7, 10; Text-figure 51

Derivation of name. From the Latin, postremus (= last).

Diagnosis. A small Rotopericyclus, lacking ribs, but with weak transverse ornament and subordinate spiral lirae.

Range. Late Chadian, upper FA Zone; Hodder Mudstone Formation (Leagram Mudstone Member).

Material. Holotype, Ro 933, partly testiferous conch, presumably mature; from locality 4a.

Description. The conch is subdiscoconic and compressed with a relatively wide, steep-walled umbilicus. The external ornament consists of coarse crenulate transverse striae, which pass retrograde over the flanks, where they are slightly concave (orad), and the venter, resulting in a deep hyponomic sinus (2 mm deep at D=16 mm). An internal ornament is not developed. Four or five strong constrictions are present; these are deeply incised next to the umbilicus and are clearly visible on the external conch surface. The suture has a slightly divergent-sided (orad) ventral lobe, with slightly flexed sides. The lateral saddle is rounded. The lateral lobe is relatively narrow, nearly as deep as the ventral lobe, and has a subacute, mammilate termination.

Remarks. Despite the abundant and diverse ammonoid fauna present in the lower part of the Hodder Mudstone Formation, there are no other records of Rotopericyclus. The new species is the youngest example known and is so distinct, in lacking ribs, that it warrants description as a new species, despite the paucity of material.

DISCUSSION AND CONCLUSIONS

The Worston Shale Group ammonoids are particularly important in relating mid-Dinantian ammonoid stratigraphy to established conodont and foraminiferal zonal schemes. This is possible because of the extremely thick sequence preserved in the Craven Basin which displays a fortuitous interbedding of limestone turbidites within the hemipelagic ammonoid mudstone facies. Despite these advantages, it has only been possible to improve controls on the absolute ranges of a few species. This is a consequence of the nature of ammonoid preservation, the many stratigraphical gaps, a lack of faunas conspecific between separate regions and poor biostratigraphical controls on correlation. These limitations are the main reason for leaving the FA and BB zones undivided into subzones

The mid-Dinantian assemblages described provide a key insight into this poorly known interval of ammonoid stratigraphy and only the faunas from the CIS surpass them in diversity. Between the CIS and the Craven Basin only a few species are conspecific, namely Dzhaprakoceras deflexum, D. diaprakense, D. subglobosum and Muensteroceratoides aksuensis, but there are some apparent differences in stratigraphical distribution. For example, in the northern Urals, D. deflexum and D. subglobosum were reported to come possibly from the BB Zone, Nortnichsky Horizon (Kusina 1980), whereas in the Craven Basin D. deflexum is recorded only from the upper FA Zone, and D. subglobosum from the middle FA zone, as in Tien-Shan (Popov 1968). These differences may be real, but it must be stressed that the northern Urals occurrences of D. deflexum and D. subglobosum were only tentatively located stratigraphically by Kusina (1980). All the genera known from the Worston Shale Group are recorded from the CIS, apart from Parahammatocyclus gen. nov. This disparity between distribution at the species and generic levels is typical of ammonoid occurrences documented from younger Carboniferous strata. Species conspecific with those from other regions include Merocanites applanatus, recorded from the Erdbach Limestone in Germany, and Dzhaprakoceras hispanicum, described from the Genicera Formation in Spain and the Tober Colleen Formation, Ireland. The faunas from the Craven Basin suggest that D, hispanicum is a longranging form, from middle FA to early BB. Because of the condensed nature of the Erdbach Limestone it is difficult to assess the true stratigraphical range of M. applanatus.

Some proposals on ammonoid evolution can also be made. Bisat (1952), when erecting the genus *Bollandites*, suggested an affinity with the pericyclids, particularly *Hammatocyclus*. This seems to be borne out by the discovery of *Parahammatocyclus* gen. nov., which appears to be the direct ancestor of *Bollandites*, resulting in a lineage originating with *Hammatocyclus*. There is still, however, the problem that *Parahammatocyclus* gen. nov. is recorded from above the first entry of *Bollandites*. *Helicocyclus* is likely to have originated from *Ammonellipsites*. Both these genera have similar sutures and a serpenticonic juvenarium, but in *Helicocyclus* the serpenticonic whorl form is maintained throughout ontogeny.

Amongst the muensteroceratids, Beyrichoceratoides, Dzhaprakoceras and Muensteroceratoides are closely related because a serpenticonic juvenarium is lacking. These genera also show orad constriction of the ventral lobe. If Muensteroceras is the ancestor, then there is a possible lineage from Muensteroceratoides through Beyrichoceratoides to Dzhaprakoceras, but the stratigraphy of these genera suggests that Dzhaprakoceras appears before Beyrichoceratoides. Unfortunately, the precise timing of the appearance of genera in the early part of the FA zone is not well enough known to resolve this problem. Bollandoceras has a serpenticonic juvenarium; it is therefore unlikely that Beyrichoceratoides or Dzhaprakoceras are ancestors; upper FA forms referred to Muensteroceras by Kusina (1980) seem more probable.

The origin of *Dimorphoceras* remains a mystery, but some juvenile involute *Dzhaprakoceras*, such as *D. levis* sp. nov., show similarity to juvenile *Dimorphoceras*. *Kozhimites*, with its strongly bisinuous ornament, laterally compressed conch and broad ventral lobe, is also a possible candidate as ancestor.

In the nomismoceratids, the lineage *Eonomismoceras–Simmonoceras–Nomismoceras* reflects the stratigraphical and morphological progression of these genera, although the stratigraphical position of *Simmonoceras* at the base of the BB Zone is not conclusively proven.

Winchelloceras seems to be close morphologically to Beyrichoceras, but unfortunately the juvenarium of the former is undescribed. There are no clues in the mid-Dinantian assemblages as to the origin of the important genus Goniatites, which appears in the late Dinantian. Some superficial similarity exists between the sutures of Goniatites and the ammonellipsitins; however, the juvenaria differ dramatically and there is a long stratigraphical gap between the extinction of the ammonellipsitins, at the top of the FA Zone, and the appearance of Goniatites, in the upper part of the B zone.

Acknowledgements. Part of this research was undertaken whilst the author was in receipt of a NERC CASE Award to the University of Bristol and the British Geological Survey (1977–1980). I thank my former supervisors Drs W. H. C. Ramsbottom and the late S. C. Matthews; R. F. Grayson who discovered the ammonoid occurrences at Salthill Quarry; staff who gave assistance whilst examining collections, especially, W. Baird, V. I. Davidov, N. Monaghan, R. M. Owens, D. Phillips, A. V. Popov and G. D. Sevastopulo. L. F. Kusina and two anonymous referees provided valuable comments on early drafts of this work. This paper is published with the permission of the Director, British Geological Survey (NERC).

REFERENCES

- AITKENHEAD, N., BRIDGE, D. Mc. C., KIMBELL, S. F. and RILEY, N. J. 1992. *Geology of the country around Garstang*. Memoir of the Geological Survey, 1:50000 sheet 67, England and Wales.
- BISAT, W. S. 1924. The Carboniferous goniatites of the North of England and their zones. *Proceedings of the Yorkshire Geological Society*, **20**, 40–124, pls 1–10.
- —— 1934. The goniatites of the *Beyrichoceras* Zone in the North of England. *Proceedings of the Yorkshire Geological Society*, **22**, 280–309, pls 17–24.
- —— 1950. The junction faunas of the Viséan and Namurian. *Transactions of the Leeds Geological Association*, **6**, 10–21, pls 1–2.
- —— 1952. The goniatite succession at Cowdale Clough, Barnoldswick, Yorkshire. *Transactions of the Leeds Geological Association*, **6**, 155–181, pls 1–3.
- BÖHM, R. 1935. Études sur les faunes du Dévonian supérieur et du Carbonifère inférieur de la Montagne Noire. Imprimé de la Charité, Montpellier, 203 pp., 10 pls.
- BROWN, D. A., CAMPBELL, K. S. W. and ROBERTS, J. 1964. A Viséan cephalopod fauna from New South Wales. *Palaeontology*, 7, 682–694, pls 102–103.
- BRÜNING, K. 1923. Beiträge zur kenntnis des rheinisch-westfälischen Unterkarbons. Geologisches Archiv, 1, 263–266.
- CAMPBELL, K. S. W., BROWN, D. A. and COLEMAN, A. R. 1983. Ammonoids and the correlation of the Lower Carboniferous rocks of eastern Australia. *Alcheringa*, 7, 75–123.
- collinson, c. 1955. Mississippian prolecanitid goniatites from Illinois and adjacent states. *Journal of Paleontology*, **29**, 433–438, pl. 45.
- CRICK, G. C. 1895. On a new species of *Prolecanites* from the Carboniferous Limestone of Haw Bank tunnel, Skipton, Yorkshire. *Transactions of the Manchester Mineralogical Society*, **23**, 80–83.
- crick, G. C. 1899. On some new or little-known goniatites from the Carboniferous Limestone of Ireland. Annals and Magazine of Natural History, 18, 429–454.
- CURRIE, E. D. 1954. Scottish Carboniferous goniatites. *Transactions of the Royal Society*, *Edinburgh*, **62**, 527–602, pls 1–4.
- DELÉPINE, G. 1940a. Les goniatites du Dinantien de la Belgique. Mémoires du Musée Royal d'Histoire Naturelle de Belgique, 91, 1-91, pls 1-5.
- 1940b. Contribution a l'étude des Goniatites du Waulsortian d'Irlande et de Belgique. Annales de la Société Géologique du Nord, 64, 134-149, pl. 4.
- —— 1941. Upper Tournaisian Goniatites from New South Wales, Australia. *Annals and Magazine of Natural History*, Series 2, 7, 386–395, pl. 5.
- —— 1943. Les faunes marines du Carbonifère des Asturies (Espagne). Memoires de l'Academie des Sciences, Institute de France, 66, 1–122, pls 1–6.
- EARP, J. R., MAGRAW, D., POOLE, E. G., LAND, D. H. and WHITEMAN, A. J. 1961. Geology of the country around Clitheroe and Nelson. Memoir of the Geological Survey of Great Britain, England and Wales, Sheet 68, 346 pp.

- ELIAS, M. K. 1956. Upper Mississippian and lower Pennsylvanian formations of south-central Oklahoma. Bulletin of the American Association of Petroleum Geologists, 1, 56–134, pls 1–6.
- FOORD, A. H. 1903. The Carboniferous Cephalopoda of Ireland. Part 5. *Palaeontographical Society Monograph*, 57 (269), 147–234, pls 40–49.
- —— and CRICK, G. C. 1897. Catalogue of fossil Cephalopoda in the British Museum (Natural History), part 3. British Museum (Natural History), London, 303 pp.
- FRECH, F. 1899. Die Steinkohlenformation. Lethaia palaeozoica (Stuttgart), 2, 257–433, pls 46a-b.
- —— 1901. Die Dyas. Lethaea geognostica (Stuttgart), 3, 435–578, pls 1–13.
- GEORGE, T. N. and PONSFORD, D. R. A. 1935. Mid-Avonian goniatites from Gower. *Annals and Magazine of Natural History*, **16**, 354–370, pl. 11.
- GIRTY, G. H. 1909. Fauna of the Caney Shale of Oklahoma. *Bulletin of the United States Geological Survey*, 377, 1–106, pls 1–13.
- GORDON, M. 1957. Mississippian cephalopods of northern and eastern Alaska. *United States Geological Survey*, *Professional Paper*, **283**, 1–61, pls 1–6.
- —— 1964. Carboniferous cephalopods of Arkansas. *United States Geological Survey, Professional Paper*, **460**, 1–322, pls 1–30.
- —— 1983. Late Kinderhookian (Early Mississippian) ammonoids from the Western United States. Bulletin of the United States Geological Survey, P 1375, 1–83.
- —— 1986. Late Kinderhookian (Early Mississippian) ammonoids of the Western United States. *Memoir of the Paleontological Society*, **19**, 1–36.
- GROESSENS, E. 1975. Distribution de conodontes dans le Dinantien de la Belgique. International symposium on Belgian micropaleontological limits from Emsian to Viséan, Namur, September, 1–10, 1974. *Geological Survey of Belgium Publication*, 17, 1–193.
- HIGGINS, A. C. and WAGNER-GENTIS, T. 1982. Conodonts, goniatites and the biostratigraphy of the earlier Carboniferous from the Cantabrian Mountains, Spain. *Palaeontology*, **25**, 313–350, pls 1–3.
- HIND, W. 1918. The distribution of the British Carboniferous goniatites. *Geological Magazine*, (6), **6**, 434–450. HOLZAPFEL, E. 1889. Die Cephalopoden führenden Kalke des unteren Carbon von Erdbach-Breitscheid bei Herborn. *Paläontologische Abhandlungen*, (N.F.1), **1**, 1–74, pls 1–8.
- HYATT, A. 1884. Genera of fossil cephalopods. *Proceedings of the Boston Society of Natural History*, 22, 253–339.
- —— 1900. Cephalopoda. 502–604. *In ZITTEL*, к. а. *Text book of palaeontology*. First English edition, translated by EASTMAN, C. R. London and New York.
- KARPINSKY, A. P. 1896. O nakhozhgenii v Azii *Prolecanites* i o razvitii etogo roda. [On the occurrence of *Prolecanites* in Asia and on the development of this genus]. *Izvestiya Imperatorskoi Akademii Nauk*, *St Petersburg*, 4, 179–194. [In Russian].
- KONINCK, L. G. de 1844. Description des animaux fossils qui se trouvent dans le terrain carbonifère de Belgique. Liège, 636 pp., pls A–H, 1–55.
- 1880. Faune du Calcaire carbonifère de la Belgique. Deuxiéme partie. Genres: Gyroceras, Cyrtoceras, Gomphoceras, Orthoceras, Subclymenia et Goniatites. Annales Musée Royale d'Histoire Naturelle de Belgique, séries paléontologie, 5, 1–133, pls 32–50.
- KORN, D. 1988. Die Goniatiten des Kulmplattenkalkes (Cephalopoda, Ammonoidea; Unterkarbon, Rheinisches Schiefergebirge). Geologie und Paläontologie in Westfalen, 11, 1–293, 60 pls.
- KULLMANN, J. 1961. Die Goniatiten des Unterkarbons im Kantabrischen Gebirge (Nordspanien). 1. Stratigraphie. Paläontologie der U. O. Goniatitina Hyatt. *Neues Jahrbuch für Geologie und Paläontologie*, 113, 219–326, pls 19–23.
- —— 1963. Die Goniatiten des Unterkarbons im Kantabrischen Gebirge (Nordspanien). 2. Stratigraphie. Paläontologie der U. O. Prolecanitina Miller & Furnish. Die Altersstellung der Faunen. Neues Jahrbuch für Geologie und Paläontologie, 116, 269–324, pls 17–20.
- KUSINA, L. F. 1971. O nekotorykh novykh i maloizvestnykh rannevizeyskikh (saurskikh) ammonoideyakh. [New and little-known early Viséan (Saourian) ammonoids]. *Paleontogicheskii Zhurnal*, **1971**, 37–48. [In Russian]. —— 1973. Krevizii roda *Muensteroceras*. [A contribution to revision of the genus *Muensteroceras*].
- Paleontogicheskii Zhurnal, 1973, 14–25. [In Russian].
- —— 1974. Saurskii kompleks rannekamennougol'nykh ammonoidei. [Sauorian Stage, early Viséan ammonoids]. *Paleontogicheskii Zhurnal*, **1974**, 18–31. [In Russian].
- —— 1980. Sauskikh ammonoidei. [Sauorian ammonoids]. Trudy Institutata Paleontologechescogo Academiya Nauk SSSR, 181, 1–108, pls 1–8. [In Russian].

- —— 1983. O nekotorykh rannekamennougol'nykh ammonoideyakh Urala. [Some new Carboniferous ammonoids from the Urals]. *Paleontogicheskii Zhurnal*, **1983**, 91–95. [In Russian].
- —— 1987. Vizeiskii kompleks ammonoidei. [Viséan Stage ammonoids]. *Paleontogicheskii Zhurnal*, **1987**, 52–61, pl. 6. [In Russian].
- and POLETAEV, V. I. 1991. Novie Vizeiskii ammonoidei Donbassa i Dneiprovsko-Donetskay Vpadeeni. [New Viséan ammonoids from the Donbases and the Dnieper-Donets Depression]. *Paleontogicheskii Zhurnal*, 1991, 35–45, pl. 4. [In Russian].
- and YATSKOV, S. V. 1987. Kamennougolnye ammonoidei Milinskoi svity Novoi Zemli. [Ammonoids from Milin Suite (Lower Carboniferous of Novaya Zemlya)]. *Byulletin' Moskovskogo Obshchestva Ispytateley Prirody, Otdel Geologicheskii*, **86**, 101–108. [In Russian].
- LEES, A. and MILLER, J. 1985. Facies variation in Waulsortian buildups, Part 2; mid-Dinantian buildups from Europe and North America. *Geological Journal*, **20**, 134–159.
- LIANG XI-LUO and WANG MINGQIAN 1991. Carboniferous cephalopods of Xinjiang. *Palaeontologica Sinica*, **180**, series B, **27**, 1–111, pls 1–40. [In Chinese with English summary].
- and ZHU KUI-YU 1988. Early Carboniferous of Baoshan, Yunnan. *Acta Palaeontologica Sinica*, 27, 288–301, pls 1–4 [In Chinese with English summary].
- LIBROVITCH, L. S. 1927. Nizhiekamennoygolnye golovonogie iz raiona ozera Son-Kul (Tyan-Shan). [Lower Carboniferous Cephalopoda of the Son-Kul region (Tien-Shan Mountains)]. *Materialy po Obshchei i Prikladnoi Geologii*, 74 Geologicheskii Komitet, 1–55, pls 1–7. [In Russian].
- —— 1941. Klass Cephalopoda, Cefalopody. [Class Cephalopoda]. 5: Atlas rukovodjashtchikh form isopaemykh faun SSSR. *Nizhnij otdel kamennougolnoj sistemy*, 132–153, pls 33–40.
- —— 1957. O nekotorykh novykh gruppakh goniatitov iz kammennougolnykh otlozhenij SSSR. [On some new groups of goniatites from the Carboniferous sediments of the USSR]. *Ezhegodnik Vsesojuznogo Paleontologicheskogo Obshchestva*, **16** (1955–1956), 246–273, pls 1–4. [In Russian].
- —— 1962. In RUZHENCEV, V. E. 1962. 5, Mollusca—Cephalopoda 1. In ORLOV, Y. A. (ed.). Osnovy paleontologii. [Fundamentals of palaeontology]. Izdatel'stvo Akademii Nauk SSSR, Moskva. 409 pp., 32 pls.
- Mc'CALEB, J. A., QUINN, J. H. and FURNISH, W. M. 1964. Girtyoceratidae in the southern Mid-continent. *Circular of the Oklahoma Geological Survey*, **67**, 1–41, pls 1–4.
- M'COY, F. 1844. A synopsis of the characters of the Carboniferous Limestone fossils of Ireland. London, 274 pp., 29 pls.
- MILLER, S. A. and FABER, C. L. 1892. Description of some Carboniferous and Subcarboniferous Cephalopoda. *Cincinnati Natural History Society Journal*, **14**, 164–168, pl. 6.
- MILLER, A. K. and FURNISH, W. M. 1954. The classification of the Paleozoic ammonoids. *Journal of Paleontology*, **28**, 685–692.
- MOJSISOVICS von MOJSVÁR, E. 1882. Die Cephalopoden der mediterranen Triasprovinz. Abhandlungen der Kaiserlich-Koniglichen Geologischen Reichsanstalt, 10, 1–322, pls 1–94.
- MOORE, E. W. J. 1930. Species of the genus *Dimorphoceras* in the Bowland shales. *Geological Magazine*, 67, 162–168.
- —— 1936. The Bowland Shales from Pendle to Dinckley. *Journal of the Manchester Geological Association*, **1**, 167–192, pls 1–3.
- —— 1939. The goniatite genus *Dimorphoceras* and its development in the British Carboniferous. *Proceedings* of the Yorkshire Geological Society, **24**, 103–128, pl. 15.
- —— 1946. The Carboniferous goniatite genera *Girtyoceras* and *Eumorphoceras*. *Proceedings of the Yorkshire Geological Society*, **65**, 387–445, pls 22–27.
- —— 1950. The genus *Sudeticeras* and its distribution in Lancashire and Yorkshire. *Journal of the Manchester Geological Association*, **2**, 31–50, pls 1–3.
- —— 1952. Notes on the Genera *Prolecanites* and *Epicanites* with descriptions of two new species. *Liverpool and Manchester Geological Journal*, 1, 71–76, pl. 7.
- —— 1958. Dimorphoceratidae from the Upper Viséan shales of County Leitrim, Eire. *Proceedings of the Yorkshire Geological Society*, **31**, 219–226, pl. 14.
- NICOLAUS, H. J. 1963. Zur Stratigraphie und Fauna der *crenistria-*zone im kulm des Rheinischen Schiefergebirges. *Beihefte zum Geologischen Jahrbuch*, **53**, 1–246, pls 1–18.
- PADGET, P. 1952. A new species of *Beyrichoceras* from the "Yoredale beds", County Tyrone, Northern Ireland. *Annals and Magazine of Natural History*, new series 12, 5, 864–868, pl. 32.
- PAREYN, C. 1961. Les massifs carbonifères du Sahara Sud-Oranais. Tome 2. *Publications du Centre de Recherches Sahariennes*, *Série Géologie*, 1, 1–244, pls 1–28.

- PARKINSON, D. 1926. The faunal succession in the Carboniferous Limestone and Bowland Shales at Clitheroe and Pendle Hill, Lancashire. *Quarterly Journal of the Geological Society*, *London*, **82**, 188–239.
- PARKINSON, J. 1822. Introduction to the study of fossil organic remains. London, 364 pp., 10 pls.
- PHILLIPS, J. 1836. Illustrations of the geology of Yorkshire, or a description of the strata and organic remains. 2, The Mountain Limestone District. John Murray, London, 253 pp., 25 pls.
- PLUMMER, F. B. and SCOTT, G. 1937. Upper Paleozoic Ammonites in Texas. *In* The Geology of Texas, volume 3, part 1. *Bulletin of the University of Texas*, **3701**, 1–516, pls 1–41.
- POPOV, A. V. 1965. Novye vizeiskii amonoidei Tyan-Shanya. [New Viséan ammonoids of the Tien-Shan]. *Paleontogicheskii Zhurnal*, **1965**, 35–49, pls 3–4. [In Russian].
- —— 1968. Vizeiskii ammonoidei Severnogo Tyan-Shanya i ikh stratigraficheskoe znachenie. [Viséan ammonoids of the northern Tien-Shan and their stratigraphical significance]. Akademiya Nauk Kirgiskoy SSR, Instituta Geologii, 116 pp., 8 pls.
- RILEY, N. J. 1990a. Revision of the *Beyrichoceras* Ammonoid-Biozone (Dinantian), NW Europe. *Newsletters on Stratigraphy*, **21**, 149–156.
- —— 1990b. Stratigraphy of the Worston Shale Group (Dinantian), Craven Basin, north-west England. *Proceedings of the Yorkshire Geological Society*, **48**, 163–187.
- —— 1991. A global review of mid-Dinantian ammonoid biostratigraphy. *Courier Forschungsinstitut Senckenberg*, **130** (for 1990), 133–143 [imprinted 1990].
- —— 1993. Dinantian (Lower Carboniferous) biostratigraphy and chronostratigraphy in the British Isles. Journal of the Geological Society, London, 150, 427–446.
- —— 1995. Foraminiferal biostratigraphy of the Chadian Stage stratotype (Dinantian), Chatburn, northwest England. *Bulletin de la Société Belge de Géologie*, **103**, 13–49.
- RUZHENCEV, V. E. 1949. Sistematika i evolyutsiya semeiskv Pronoritidae Frech i Medlicottiidae Karpinsky. [Systematics and evolution of the families Pronoritidae Frech and Medlicottiidae Karpinsky]. *Trudy Paleontologicheskogo Instituta*, *Akademii Nauk SSSR*, 19, 1–206. [In Russian].
- —— 1957. Filogenetischekaya sistema paleozoiskikh ammonoidei. [Phylogenetic systematics of the Palaeozoic Ammonoidea]. *Byulleten' Moskovskogo Obshchestva Ispytatelei Prirody*, otdel geologicheskii, **32**, 49–64. [In Russian].
- —— 1960. Printsipy sistematiki sistema i filogeniya paleozoiskikh ammonoidei. [Principles of systematics, classification and phylogeny of the Palaeozoic Ammonoidea]. *Trudy Paleontologicheskogo Instituta*, *Akademii Nauk SSSR*, **83**, 1–331. [In Russian].
- —— 1962. 5, Mollusca Cephalopoda 1. *In ORLOV*, Y. A. (ed.). *Osnovy paleontologii*. [Fundamentals of palaeontology]. Izdatel'stvo Akademii Nauk SSSR, Moskva. 409 pp., 32 pls. [In Russian].
- —— 1965. Osnovnye kompleksy ammonoideii kammennougol'nogo perioda. [Basic ammonoid complexes of the Carboniferous Period]. Paleontogicheskii Zhurnal, 1965, 3–17. [In Russian].
- —— 1966. Vizeyskiye ammonoidey na Yuzhnom Urale. [Viséan ammonoids from the south Urals]. *Paleontogicheskii Zhurnal*, **1966**, 47–59, pl. 5. [In Russian].
- —— and BOGOSLOVSKAYA, M. F. 1971. Namyurskii etap v evolyutsii ammonoideii. [Namurian time in ammonoid evolution]. *Trudy Paleontologicheskogo Instituta, Akademii Nauk SSSR*, **113**, 1–382, pls 1–40. [In Russian].
- SAUNDERS, W. B. 1966. New goniatite ammonoid from the Late Mississippian of Arkansas. *Oklahoma Geology Notes*, **26**, 43–48, pl. 1.
- SCHINDEWOLF, O. H. 1922. Über eine Unterkarbofauna aus ostthuringen. Senckenbergiana, 4, 8-20.
- —— 1951a. Über ein neues Vorkommen unterkarbonischer *Pericyclus*-Schichten im Oberharz. *Sonderdruck aus dem Neuen Jahrbuch für Geologie und Paläontologie, Abhandlungen,* **93,** 23–116, pls 3–7.
- —— 1951b. Zur Gleiderung der *Pericyclus*-Gruppe (Cephal., Goniatit.). *Neues Jahrbuch für Geologie und Paläontologie, Monatshefte 1951*, **10**, 305–310.
- SCHMIDT, H. 1924. Zwei Cephalopodenfaunen an der Devon-Carbongrenze im Sauerland. *Jahrbuch der Preussischen Geologischen Landesanstalt*, **44**, 98–171, pls 6–8 [for 1923].
- —— 1925. Die carbonischen Goniatiten Deutschlands. *Jahrbuch der Preussischen Geologischen Landesanstalt*, für 1924, **45**, 489–609, pls 19–26 [for 1924].
- —— 1941. Eine neue Fauna mit *Pericyclus* von Riefensbeek im Harz. *Jahrbuch der Reichsstelle fur Bodenforschung*, **60 (1939)**, 148–156, pls 19–20.
- SHENG HUAIBIN 1984. Lower Carboniferous ammonoid faunule from the Zhifang area Xinjiang. *Acta Geologica Sinica (Peking)*, **58**, 284–291, pl. 1. [In Chinese, with English summary].
- SHIMANSKY, V. N. and KUSINA, L. F. 1977. Rannekamennougol'nyye golovonogiye Pripolyarnogo Urala. [Lower

- Carboniferous Cephalopoda of the Polar Urals]. Byulletin' Moskovskogo Obshchestva Ispytateley Prirody, Otdel Geologicheskii, **52**, 79–90, pls 1–2. [In Russian].
- SMYTH, L. B. 1951. A Viséan cephalopod fauna in the Rush Slates of Co. Dublin. *Proceedings of the Royal Irish Academy*, **53**, 289–309, pl 11.
- SOWERBY, J. 1813. Mineral conchology of Great Britain, Volume 1, London, 234 pp., 102 pls [imprinted 1812].
 —— 1821. Mineral conchology of Great Britain, Volume 3, London, 194 pp., pls 204–306.
- SPATH, L. F. 1934. Catalogue of fossil Cephalopoda in the British Museum (Natural History), Part 4, The Ammonoidea of the Trias. British Museum (Natural History), London, 521 pp.
- TURNER, J. S. 1948. Mid-Dinantian reef limestones of Dublin and Cork. *Transactions of the Leeds Geological Association*, **6**, 44–56.
- WAGNER-GENTIS, C. H. T. 1960. On Nautellipsites hispanicus (Foord and Crick). Estudios Geologicos Instituto "Lucas Mallada", 16, 43–51.
- 1964. 6. Description of Goniatites. *In Higgins*, A. C., Wagner-Gentis, C. H. T. and Wagner, R. H. Basal Carboniferous strata in part of Northern Leon, NW Spain: stratigraphy, conodont and goniatite faunas. *Bulletin de la Société Belge de Géologie, de Paléontologie et d'Hydrologie*, 72, 228–248, pls 1–4.
- WEDEKIND, R. 1918. Die Genera der Paleoammonoidea (Goniatiten). *Palaeontographica*, **62**, 85–184, pls 14–22. WEYER, D. 1965. Zur Ammonoideen-Fauna der *Gattendorfia-*Stufe von Dzikowiec (Ebersdorf) in Dolny Slak (Niederschlesien), Polen. *Berichte der Geologischen Gesellschaft in der Deutschen Demokratischen Republik*, **10**, 443–464, pls 1–8.
- WEYER, D. 1972a. Trilobiten und Ammonoideen aus der *Entogonites nasutus-*Zone (Unterkarbon) des Buchenbergsattels (Elbingeroder Komplex, Harz), Teil 1. *Geologie*, **21**, 166–184.
- —— 1972b. Trilobiten und Ammonoideen aus der *Entogonites nasutus-*Zone (Unterkarbon) des Buchenbergsattels (Elbingeroder Komplex, Harz), Teil 2. *Geologie*, **21**, 318–349.
- WINCHELL, A. 1862. Notice of rocks lying between the Carboniferous Limestone of the Lower Peninsula of Michigan and the limestones of the Hamilton Group; with descriptions of some cephalopods supposed to be new to science. *American Journal of Science and Arts*, Series 2, 33, 352–366.

NICHOLAS J. RILEY
British Geological Survey
Keyworth, Nottingham NG12 5GG, UK

Revised typescript received 7 July 1995