TWO CONODONT FAUNAS FROM THE LOWER CARBONIFEROUS OF CHUDLEIGH, SOUTH DEVON

by S. C. MATTHEWS

ABSTRACT. Two conodont faunas are reported from the Lower Carboniferous of Chudleigh, south Devon. The lower includes Siphonodella cooperi, S. isosticha and Gnathodus aff. semiglaber. The higher has these siphonodellids in association with Scaliognathus anchoralis. Comparisons are made with other European and with North American records.

Two conodont faunas collected at Chudleigh by Professor M. R. House and Mr. N. E. Butcher provide proof of the Dinantian age of part of the stratigraphic succession there. Only some details of the fauna are given here. The occurrences of conodonts will be discussed in their wider geological context in House and Butcher's forthcoming paper on the Chudleigh district.

The lower of the two faunas comes from the southern side of a track 110 ft. WSW. of Winstow Cottages and immediately below a quarry in the Winstow Cherts. The horizon is almost at the base of the Winstow Cherts of House and Butcher. The conodont material exists in this case as empty moulds of specimens in the siliceous shale. The higher fauna was found in siliceous shales exposed in a temporary trench cut beside a hedge at a point 180 ft. to the south-west of Winstow Cottages. House and Butcher place this horizon in the lithological transition from Winstow Cherts to Posidonia Shales (see their paper). This second collection of material showed light-coloured conodont specimens (presumably degraded remnants of the originally phosphatic microfossils) set in now slightly coarser siliceous rocks. It proved useful to remove the substance of these specimens with concentrated hydrochloric acid in order to produce again a set of moulds. Rubber casts could then be prepared and studied.

The following have been identified:

Lower fauna

Falcodus? sp.
Gnathodus aff. semiglaber (Bischoff)
Gnathodus sp.
Hindeodella sp. indet.
Polygnathus communis Branson & Mehl
Siphonodella cooperi Hass
Siphonodella isosticha (Cooper)
Chondrites sp. and other ichnofossils

Higher fauna

Gnathodus delicatus Branson and Mehl Gnathodus semiglaber (Bischoff)
Gnathodus texanus (sensu germanico)
Gnathodus ef. punctatus (Cooper)
Hindeodella ef. segaformis Bischoff
Polygnathus sp.
Pseudopolygnathus triangula pinnata Voges
Scaliognathus anchoralis Branson and Mehl Siphonodella cooperi Hass
Siphonodella sexplicata (Branson and Mehl)
Siphonodella isosticha (Cooper)
Ostracods
Chitinophosphatic brachiopod

[Palaeontology, Vol. 12, Part 2, 1969, pp. 276-280, pl. 51.]

The conodonts, ostracods, and the chitinophosphatic brachiopod may be taken to be representatives of the pelagic fauna of the times. The ichnofossils, on the other hand, record the activity of members of an endofauna. It may be observed that there is no systematic relationship between the distribution of the conodonts and the courses taken by the endobionts.

The matter of the ages of the two faunas demands careful discussion. The lower fauna may be treated as resembling some mentioned by Voges (1960) and regarded by him as having their place in the upper part of his Siphonodella crenulata-Zone. He identified, in these cases, relatively early occurrence of Gnathodus semiglaber in association with siphonodellids. Ziegler (1960) reported a gnathodid-siphonodellid association from a pre-anchoralis-Zone situation and Böger (1962) recorded a similar coincidence of form. All of these authors, at their times of writing, would have suggested that a Siphonodella crenulata-Zone conodont fauna implies cuIIα age by the cephalopod-based Dinantian standard. Likewise, a first statement on the Chudleigh Dinantian conodont faunas (reported in House and Butcher 1962) gave the age of the lower fauna as cuIIα. But this equation can now be seen to be invalid (Matthews, in press). The only permissible translation into cephalopod terms would suggest that the Siphonodella crenulata-Zone post-dates the Hangenbergkalk (cuI) and pre-dates any western European cuII cephalopod fauna recorded in the literature.

The higher fauna may be referred with confidence to Voges's anchoralis-Zone. Scaliognathus anchoralis and Hindeodella cf. segaformis have been identified, and nothing in the remaining elements of the fauna would contradict their indication of age. Conodont faunas of this character have been obtained from cuII horizons.

Comparisons with American findings are instructive. The lower fauna, with Siphonodella cooperi, S. isosticha, Polygnathus communis, and the gnathodids, matches in many respects the association of form reported by Collinson, Scott, and Rexroad (1962) from the late Kinderhookian of the Mississippi Valley and treated by them as determining a Siphonodella n. sp. A-Siphonodella cooperi Zone. Later, Rexroad and Scott (1964) advised that Siphonodella n. sp. A should have the name Siphonodella isosticha. These American investigations produced no account of the anchoralis-Zone fauna which European experience (now repeated again at Chudleigh) had shown to succeed the faunas dominated by siphonodellids. Instead, Collinson, Scott, and Rexroad (1962) referred to a major unconformity as fixing the local upper limit of occurrence of Siphonodella and other forms in the Mississippi Valley. The fact of interruption was not, however, made plain in their offer of a sequence of zones, nor in their set of proposals on European correlatives. Rexroad and Scott's (1964) paper made more explicit reference (fig. 1 of their paper) to a gap in the Mississippi Valley standard. Burton (1964) has recorded the presence in the Mississippian of New Mexico of Scaliognathus, Doliognathus and other forms familiar in Europe. He emphasizes (personal communication) the fact that faunal breaks are sharp and coincide with member or formation boundaries throughout the Mississippian of the Sacramento Mountains. From these observations he reasonably draws the conclusion that the ranges of particular forms might appear to be quite different in fuller rock-successions elsewhere. Thompson (1967) encountered Scaliognathus and Doliognathus in his study of certain conodont faunas from southwestern Missouri. It was in fact in the Mississippian of that region that these two genera were first found (Branson and Mehl 1941).

The two faunas from Chudleigh are from a single section in basinal stratigraphy, and the higher fauna shows coincidence of *Scaliognathus* and siphonodellids. No such coincidence is encountered in North America. One possible interpretation of the Chudleigh evidence is that there is recorded, here, a set of associations of conodonts the pattern of whose occurrence in the Mississippian is distorted by interruptions of the stratigraphic succession. If the Mississippian evidence is not in this way distorted, it may be necessary to recognize a differentiation of faunas by which the *anchoralis-Zone* associations recorded in Europe were made distinct from their Mississippian contemporaries. But in that case, there would remain the fact that no such differentiation is evident in the faunas of time immediately preceding.

Acknowledgements. I am grateful to Professor House and Mr. Butcher for allowing me the opportunity to study and make comment on their collection of Lower Carboniferous conodonts from Chudleigh. The photographs are the work of Mr. E. W. Seavill to whom also I offer my thanks.

SYSTEMATIC NOTES

Numbers given, with prefix LZ, refer to the collections of the Institute of Geological Sciences (Geological Survey and Museum). Each four-figure number identifies a rock specimen. A fifth figure, added as a suffix, identifies one particular conodont mould among the number available on the surface of that rock specimen. Thus in Plate 51, figs. 1 and 2 (oral and aboral views of *Scaliognathus anchoralis*) are both taken from LZ 5660 and cannot refer to one and the same conodont. Conversely, figs. 3 and 4 on that plate (aboral and oral views of *Polygnathus* sp.) although they refer to rock-specimens LZ 5664 and LZ 5665, do represent two different aspects of a single individual, revealed as moulds on a rock surface (on 5664) and its counterpart (on 5665).

Falcodus? sp.

Plate 51, fig. 10

Material. LZ 5649.

Remarks. Hass (1959) illustrated a Falcodus from the Chappell limestone of Texas and Klapper (1966), too, has recorded a Mississippian falcodid. Krebs (1960) reported Falcodus? sp. in a triangula inaequalis-Zone fauna and Dvořák and Freyer (1961) listed Falcodus sp. A and Falcodus sp. B in a fauna which they suggested to be of an age near

EXPLANATION OF PLATE 51

All figures \times 30.

Figs. 1-10 refer to conodonts in the higher fauna, and 11-18 to conodonts in the lower fauna.

Figs. 1, 2. Scaliognathus anchoralis. 1, Aboral view of LZ 5660/2. 2, Oral view of LZ 5660/3.

Figs. 3, 4. Polygnathus sp. Oral (LZ 5664, fig. 3) and aboral (LZ 5665, fig. 4) views of one conodont.

Figs. 5, 16. Siphonodella isosticha. LZ 5662/1, LZ 5652/1 Oral.

Fig. 6. Siphonodella sexplicata LZ 5660/1, Oral.

Fig. 7. Gnathodus delicatus. LZ 5661/2. Oral.

Fig. 8. Pseudopolygnathus triangula pinnata. LZ 5663. Oral.

Figs. 9, 14, 15. Siphonodella cooperi. LZ 5661/1, LZ 5653, LZ 5650. Oral.

Fig. 10. Falcodus? sp. LZ 5649. Lateral.

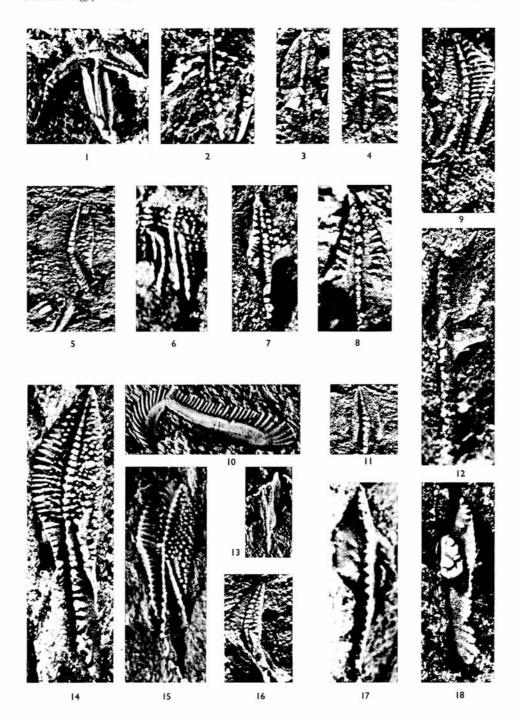
Fig. 11. Polygnathus communis LZ 5651/1. Oral.

Fig. 12. Gnathodus cf. punctatus. LZ 5666. Oral.

Fig. 13. Siphonodella sp. LZ 5651/2. Aboral.

Figs. 17, 18. Gnathodus aff. semiglaber. Oral views of LZ 5659 and LZ 5658.

Palaeontology, Vol. 12 PLATE 51



MATTHEWS, Carboniferous conodonts

to the cuI-cuII limit. Krebs interpreted his specimen as having been reworked. Neither in his nor in Dvořák and Freyer's case was there illustration or description of specimens. There is therefore no possibility of attempting comparisons with the form encountered at Chudleigh.

There must be considerable uncertainty as to the proper generic placing of this specimen. In some respects there is similarity to *Angulodus*. Or, again, there is the form to which Rexroad and Furnish (1964) gave the name *Hindeodus* and which is, by their interpretation a *Hindeodella* deviant. A similar interpretation may apply here.

Gnathodus aff. semiglaber Bischoff 1957

Plate 51, figs. 17, 18

Material. LZ 5659, LZ 5658.

Remarks. The holotype has a cup which is developed to the posterior end on either side of the carina, and the Chudleigh form differs from that condition in having no cupdevelopment on the outer side of the posterior part of the carina. Gnathodus sp. C of Thompson (1967) is closely similar.

Siphonodella cooperi Hass 1959

Plate 51, figs. 9, 14, 15

Material. LZ 5661/1, LZ 5653, LZ 5650.

Remarks. Hass introduced this species of the genus Siphonodella and also S. obsoleta, which is distinguished by the ornament (lacking transverse ridges) on the outer side of the platform and by the notably long rostral ridge which merges posteriorly with the outer margin of the platform. Voges (1959), referring to German occurrences of S. obsoleta, offered the observation that some specimens from the anchoralis-Zone resemble S. cooperi in the details of their platform ornament and suggested that such characteristics do not provide a means of separating the two species (see Voges, 1959, pl. 35, figs. 48–50 for illustration of forms of S. obsoleta with 'atypical upper surfaces of platforms'). It is possible that certain existing records of S. obsoleta may deserve revision. The Chudleigh specimens are consistent with S. cooperi as given by Hass. It can be seen that the carina is slim in its anterior part, then more powerfully developed along that part of the platform which is posterior to the termination of the inner rostral ridge.

Siphonodella isosticha Cooper 1939

Plate 51, figs. 5, 16

Material. LZ 5651/1, LZ 5652/1, LZ 5654/1, LZ 5655/1, LZ 5662/1.

Remarks. Rexroad and Scott (1964) advised that certain siphonodellids earlier treated by Collinson, Scott, and Rexroad (1962) as Siphonodella n. sp. A should be referred to Cooper's species S. isosticha. In the diagnosis offered by Rexroad and Scott a smooth platform (except for the carina and rostral ridges) is specified. A slight development of nodes is nevertheless evident on the specimen which they illustrate on pl. 3, fig. 23 of their paper. Some nodes remain on certain of the Chudleigh specimens also, but they

are not of an abundance to suggest reference to another species. In all cases, the development of rostral ridges is quite other than that which distinguishes *S. obsoleta*. A specimen illustrated by van Adrichem Boogaert (1967, pl. 3, fig. 15) is in some ways comparable with *S. isosticha*.

REFERENCES

- BISCHOFF, G. 1957. Die Conodonten-Stratigraphie des rheno-herzynischen Unterkarbons mit Berücksichtigung der Wocklumeria-Stufe und der Devon/Karbon-Grenze. *Abhandl. hess. Landesamt. Bodenforsch.* 19, 64 pp.
- BÖGER, H. 1962. Zur Stratigraphie des Unterkarbons im Velberter Sattel. Decheniana, 114, 113-70, pl. I. BOOGAERT, H. A. VAN A. 1967. Devonian and Lower Carboniferous conodonts of the Cantabrian Mountains (Spain) and their stratigraphic application. Proefschrift, University of Leiden.
- BRANSON, E. B., and MEHL, M. G. 1941. New and little known Carboniferous conodont genera. J. Paleont. 15, 97–106, pl. 19.
- BURTON, R. C. 1964. A preliminary range chart of Lake Valley formation (Osage) conodonts in the southern Sacramento Mountains, New Mexico. New Mexico Geological Society, 15th Field Conference, 73-5.
- COLLINSON, C., SCOTT, A. J., and REXROAD, C. B. 1962. Six charts showing biostratigraphic zones and correlations based on conodonts from the Devonian and Mississippian rocks of the upper Mississippi Valley. Circ. Illinois geol. Surv. 328, 1–32.
- COOPER, C. L. 1939. Conodonts from a Bushberg-Hannibal horizon in Oklahoma. J. Paleont. 13, 379-422, pl. 39-47.
- DVOŘÁK, J., and FREYER, G. 1961. Die Devon/Karbon-Grenze im Mährischen Karst (Südteil des mährischen Sedimentationsbeckens) auf der Grundlage von Conodontenfaunen. *Geologie*, 10, 881–95.
- HASS, W. H. 1959. Conodonts from the Chappel limestone of Texas. *Prof. Pap. U.S. geol. Surv.* **294–J**, 365–99, pl. 46–50.
- HOUSE, M. R., and BUTCHER, N. E. 1962. Excavations in the Devonian and Carboniferous rocks of the Chudleigh area, south Devon. Proc. Ussh. Soc. 1, 28-9.
- KLAPPER, G. 1966. Upper Devonian and Lower Mississippian conodont zones in Montana, Wyoming, and South Dakota. *Univ. Kansas Paleont. Contr. Pap.* 3, 1–43, pl. 1–6.
- KREBS, W. 1960. Neue Ergebnisse zur Stratigraphie des Oberdevons und Unterkarbons in der südwestlichen Dill-Mulde (Rheinisches Schiefergebirge). Notizbl. hess. Landesamt. Bodenforsch. Wiesbaden, 88, 216-42.
- MATTHEWS, S. C. (in press). Comments on palaeontological standards for the Dinantian C.R. Congr. Avanc. Étud. Stratigr. carb. Sheffield, 1967.
- REXROAD, C. B., and FURNISH, W. M. 1964. Conodonts from the Pella Formation (Mississippian) of south-central Iowa. J. Paleont. 38, 557-76, pl. 111.
- and scott, A. J. 1964. Conodont zones in the Rockford limestone and the lower part of the New Providence Shale (Mississippian) in Indiana. *Bull. Indiana Geol. Surv.* 30, 7–54, pl. 2, 3.
- THOMPSON, T. L. 1967. Conodont zonation of Lower Osagean rocks (Lower Mississippian) of south-western Missouri. *Missouri geol. Surv. wat. Resour.*, Rept. Invest. 39, 88 pp., pl. 1-6.
- VOGES, A. 1959. Conodonten aus dem Unterkarbon I und II (Gattendorfia- und Pericyclus-Stufe) des Sauerlandes. Paläont. Z. 33, 266–314, pl. 33–5.
- —— 1960. Die Bedeutung der Conodonten für die Stratigraphie des Unterkarbons I und II (Gatten-dorfia- und Pericyclus-Stufe) im Sauerland. Fortschr. Geol. Rheinld. Westf. 3 (1), 197–228.
- ZIEGLER, W. 1960. In KRONBERG, P., PILGER, A., SCHERP, A., and ZIEGLER, W., Spuren altvariscischer Bewegungen im nordöstlichen Teil des Rheinischen Schiefergebirges. Fortschr. Geol. Rheinld. Westf. 3 (1), 1-46, pl. 1-7.

S. C. MATTHEWS Geology Department Queen's Building University Walk Bristol. 8

Typescript received 10 May 1968